



Variation in stable carbon and oxygen isotopes of individual benthic foraminifera: tracers for quantifying the magnitude of isotopic disequilibrium

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Abstract. Stable carbon and oxygen isotopic compositions ($\delta^{13}\text{C}$ and $\delta^{18}\text{O}$) of benthic foraminiferal carbonate shells have been used to reconstruct past bottom-water environments. However, the details of factors controlling the isotopic disequilibrium between the shells and the surrounding bottom seawater (so-called the “vital effect”) are still ambiguous. In this study, we analyzed the isotopic composition of individual benthic foraminifera of multiple species by using a customized high-precision analytical system, and found that the magnitude of the isotopic disequilibrium between benthic foraminiferal shell and the surrounding bottom seawater ($\delta^{13}\text{C}_{\text{DIC}}$ and $\delta^{18}\text{O}_{\text{water}}$) in different species is correlated with inter-individual isotopic variations. As a result, we can choose suitable species as bottom-water proxies by using the inter-individual isotopic variations. In addition, by using the simplified interpretation of the inter-individual and inter-species isotopic variations established in this study, we could reconstruct the $\delta^{13}\text{C}$ values of dissolved inorganic carbon in bottom water by correcting foraminiferal isotopic compositions for the isotopic shift resulting from the isotopic effects (vital effect, microhabitat effect, and many other reported isotopic effects). Our findings will allow the use of isotope data for benthic foraminifera as more reliable proxies

for reconstructing past bottom-water conditions and evaluating global carbon cycling.

1 Introduction

Variations in the stable carbon and oxygen isotopic composition ($\delta^{13}\text{C}$ and $\delta^{18}\text{O}$) of foraminiferal carbonate shells have been used for over half a century to estimate paleoenvironmental parameters such as temperature, to quantify global changes in sea level and deep-sea circulation, and to document events such as large seafloor methane releases (e.g. Emiliani, 1955; Shackleton and Opdyke, 1973; Kennett et al., 2000; Zachos et al., 2001). In particular, $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ of benthic foraminiferal shells have been used as tracers to reconstruct past bottom-water environments. However, there is still ambiguity concerning the factors controlling the isotopic disequilibrium between the isotopic compositions ($\delta^{13}\text{C}$ and $\delta^{18}\text{O}$) of benthic foraminifera and the environmental factors ($\delta^{13}\text{C}$ of dissolved inorganic carbon (DIC) and $\delta^{18}\text{O}$ of seawater) (Grossman, 1987; McCorkle et al., 1990; Spero et al., 1997; Zeebe et al., 1999; Bijma et al., 1999; Erez, 2003; Schmiiedl et al., 2004).

Major isotopic variations in some species of benthic foraminifera are utilized as paleoindicators of bottom-water conditions, but recently researchers have begun to quantify in detail the relationship between the isotopic composition of benthic foraminifera and environmental factors (e.g. Rathburn et al., 1996; McCorkle et al., 1997; Rathburn et al., 2003; Mackensen and Licari, 2004; Martin et al., 2004; Schmiedl et al., 2004; Fontanier et al., 2006; McCorkle et al., 2008). Precise calibration and validation of isotopic indicators in benthic foraminifera will broaden the range of their application as paleoenvironmental tracers. However, two problems in using the isotopic evidence in foraminiferal shells have limited their range of application and the material available for use.

First, until recently it was not possible to analyze the stable isotopic compositions of carbonate samples smaller than about 20 μg (Revesz and Landwehr, 2002; Ishimura et al., 2004; de Groot, 2008; Velivetskaya et al., 2009) and obtain results with an acceptable error range; thus, each sample analyzed included multiple individuals, resulting in isotopic values averaged across individuals. Among thousands of calcareous foraminifera, the reported isotopic data of individual foraminiferal shells in previous studies are limited to the taxa that have large and thick CaCO_3 shells (e.g. Rathburn et al., 2003; Uchida et al., 2008; Martin et al., 2010; Bernhard et al., 2010a). Therefore, it was difficult to obtain data from sites with only foraminifera of small size or rare occurrence, for example, the deep sea or the high latitudes.

Second, although the isotopic composition of biogenic carbonate is primarily determined by water temperature, $\delta^{18}\text{O}$ of water, and $\delta^{13}\text{C}$ of DIC, the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of deep-sea benthic foraminifera are known to deviate widely from equilibrium with seawater (Grossman, 1987; Rathburn et al., 1996; McCorkle et al., 1997; Fontanier et al., 2006; Mackensen and Licari, 2004), a variation generally known as the “vital effect”. Possible causes of isotopic disequilibrium include respiration, ontogenetic effects, microhabitats, or carbonate ion concentrations, and many other factors have been suggested in previous studies (Rohling and Cooke, 1999; Mackensen, 2008). The complex interactions between these factors and the isotopic composition of biogenic calcite make it difficult to discuss them separately, or to know the original values in calcite at equilibrium with bottom water. To reduce this “isotopic noise”, isotopic analyses in the past had to include several individual foraminifera. However, such average isotopic values do not always indicate the absolute isotopic values at equilibrium with ambient water (e.g. Grossman, 1987), and the details of the isotopic variations of whole benthic foraminifera are not well known. In other words, we could not see the whole image of characteristics of the isotopic disequilibrium of the benthic foraminifera. The stable isotopes in foraminifera could be more effectively used as environmental proxies, for example for modeling global carbon cycling, with a clearer understanding of the characteristics of their isotopic disequi-

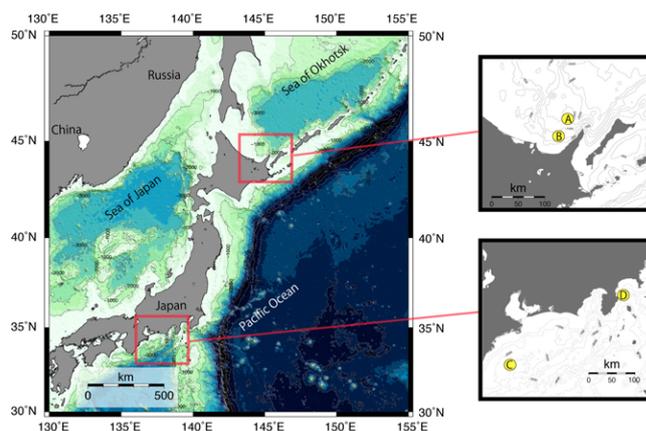


Fig. 1. Map of the study area and sampling sites. Site A ($44^{\circ}10' \text{N}$, $144^{\circ}45' \text{E}$; water depth 870 m) and site B ($44^{\circ}31' \text{N}$, $145^{\circ}00' \text{E}$; water depth 1208 m), Sea of Okhotsk off Abashiri; site C ($33^{\circ}42' \text{N}$, $137^{\circ}05' \text{E}$; water depth 1881 m), Nankai Trough off Kumano; site D ($35^{\circ}00' \text{N}$, $139^{\circ}14' \text{E}$; water depth 1182 m), Sagami Bay.

librium. In this study, we analyzed the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ of individual shells of deep-sea benthic foraminifera in core-top samples from four sites in continental margin of the northwestern Pacific Ocean to characterize the magnitude of inter-individual isotopic variations by using a custom-made high-precision isotope analysis system (Ishimura et al., 2004, 2008). The quantity required by their system is less than 1/100 of that required by conventional analytical methods. Furthermore, the use this technique allows inter-individual and inter-species variability in recorded stable isotope signatures to be determined. We expect the results to be useful for exploring which species are most appropriate to use as paleoindicators in paleoenvironmental studies.

2 Materials and methods

2.1 Foraminifer and water samples

Surface sediment samples were collected with multicorers from two stations close to each other on the same continental slope in the southwestern Sea of Okhotsk off Abashiri, Hokkaido Island, Japan. Samples were collected from a water depth of 870 m at $44^{\circ}10' \text{N}$, $144^{\circ}45' \text{E}$ during cruise HO76 of R/V *Hokusei* (Hokkaido University, Japan) in September 1997 and from 1208 m at $44^{\circ}31' \text{N}$, $145^{\circ}00' \text{E}$ during cruise MR06-04 of R/V *Mirai* (Japan Agency for Marine-Earth Science and Technology (JAMSTEC)) in October 2006 (Fig. 1).

The top layer of the surface sediment (0–10 cm below the seafloor) was subsampled every centimeter and used to determine inter-individual isotopic variation for each foraminiferal species selected for study. The sediment was stained with 0.5 % Rose Bengal solution for at least one week to distinguish living (Rose Bengal stained) foraminifera from

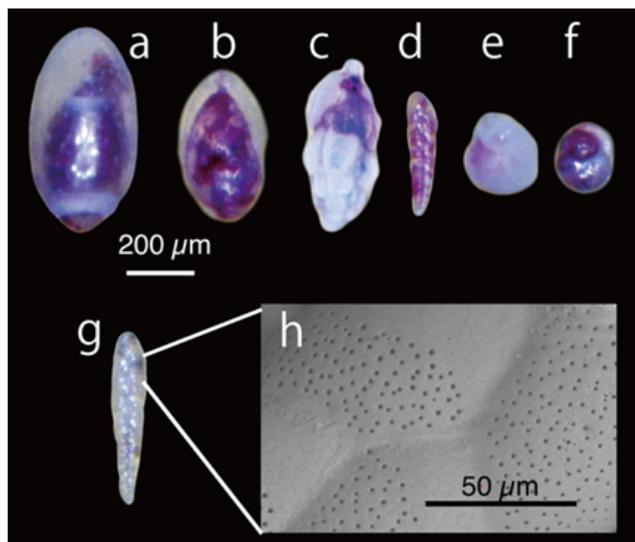


Fig. 2. Photomicrographs of Rose Bengal-stained specimens of analyzed foraminifera (a–g). (a) *Chilostomella ovoidea*, (b) *Globobulimina auriculata*, (c) *Uvigerina akitaensis*, (d) *Brizalina pacifica*, (e) *Islandiella norcrossi*, (f) *Takayanagia delicata*, (g) dead *B. pacifica*. (h) Scanning electron micrograph of *B. pacifica*.

dead (not stained) (Fig. 2). We did not use ethanol and formaldehyde but filtered seawater for the Rose Bengal solution. Previous studies reported that Rose Bengal staining method does not allow accurate identification of living individuals (e.g. Bernhard, 1988; Bernhard et al., 2006). However, during the isotopic analytical procedures of stained individuals in this study, we observed that foraminiferal soft tissues remained completely in the reaction tubes after the acid reaction of the calcite shell. Therefore, we considered the stained individuals as living. Note that our analytical method (Ishimura et al., 2004) allows us to observe the entire reaction process under a microscope. The samples were then washed in a 63- μm -mesh sieve with 50 °C water. The sediment retained in the sieve was oven-dried at 40 °C, and individual foraminifera were picked out for isotopic analysis.

The dominant species in these samples were the hyaline calcareous benthic foraminifera of the *Cassidulina* group (*Cassidulina norvangi*, *Islandiella norcrossi*, *Takayanagia delicata*), *Brizalina pacifica*, and *Stainforthia loeblichii* (Fig. 3). We analyzed the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of the benthic foraminifera *C. norvangi*, *B. pacifica*, and *S. loeblichii* from the 870-m samples, and *Uvigerina akitaensis*, *I. norcrossi*, *T. delicata*, *Chilostomella ovoidea*, *Globobulimina auriculata*, and *B. pacifica* from the 1208-m samples (Fig. 3). All specimens used for isotopic measurements were examined under a stereomicroscope and confirmed to have transparent calcite shells with no authigenic carbonate (Fig. 2). We cleaned all individual foraminifera with Milli-Q water before isotopic analysis, and we did not find any dissolution and remineralization in our analyzed samples.

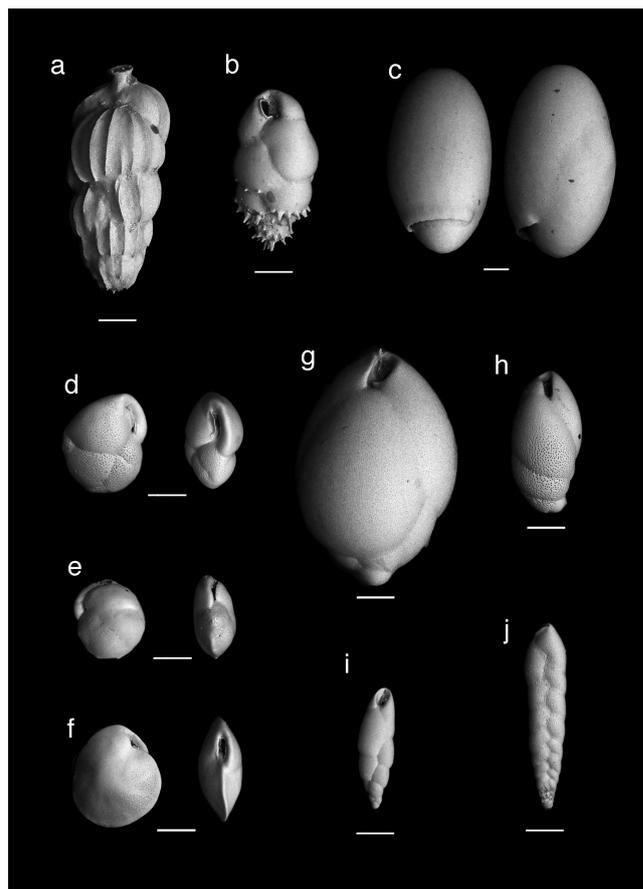


Fig. 3. Scanning electron microscope images of studied species. (a) *Uvigerina akitaensis*, (b) *Bulimina aculeata*, (c) *Chilostomella ovoidea*, (d) *Cassidulina norvangi*, (e) *Takayanagia delicata*, (f) *Islandiella norcrossi*, (g) *Globobulimina auriculata*, (h) *Rutherfordoides cornuta*, (i) *Stainforthia loeblichii*, and (j) *Brizalina pacifica*. Scale bars represent 100 μm . (a), (c), (d), (e), (f), (i) and (j); Sea of Okhotsk; (b), Nankai Trough; (g) and (h), Sagami Bay.

Water samples were also collected from the 1208-m water-depth site where foraminiferal samples were collected during cruise MR06-04 of the R/V *Mirai* in August 2006 (Table 1). Pore-water samples were squeezed from sediments that had been collected with a 50-cm-long multicorer, at depth intervals of 0, 1–3, 4–6, and 7–9 cm below the sea floor. Water-column samples were collected using Niskin bottles mounted on a conductivity-temperature-depth (CTD) rosette sampler and used for isotopic analysis.

For comparison with our samples, we also determined the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of *Bulimina aculeata* (Fig. 3) in surface sediment from the Nankai Trough (33°42' N, 137°05' E; water depth 1881 m, Fig. 1) off Kumano, Japan, collected with a push corer by the submersible *Shinkai6500* during cruise YK02-02 of R/V *Yokosuka* (JAMSTEC) in May 2002. This species has been used previously for the isotopic reconstruction of paleoenvironments (Oba, 1988; McCorkle et al., 1997; Mackensen and Licari, 2004). The sediment

sample from 0 to 7 cm below the sea floor was used for pore-water analysis and foraminiferal examination and then treated in the same manner as the samples from the Okhotsk Sea for analysis of the stable isotopes of individual shells. We also determined the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of *G. auriculata*, *C. ovoidea*, and *Rutherfordoides cornuta* (Fig. 3) in the upper 2 cm of sediment collected with a push corer from Sagami Bay (35°00' N, 139°14' E; water depth 1182 m) during cruise KT05-18 of R/V *Tansei* (JAMSTEC) in August 2005.

2.2 Stable carbon and oxygen isotope analysis

We used a continuous-flow isotope ratio mass spectrometry analytical system (Ishimura et al., 2004, 2008) to determine $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ of individual foraminifera. This system allowed us to determine $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ of as little as 0.2 μg CaCO_3 with a long-term external precision of better than $\pm 0.10\%$ for $\delta^{13}\text{C}$ and $\pm 0.14\%$ for $\delta^{18}\text{O}$. This system can be used for high-precision stable isotope measurements of all foraminifera, including even the smallest species (Kimoto et al., 2009). Isotopic values are reported relative to the Vienna Pee Dee belemnite (VPDB) standard. In addition to determining stable isotopes in both living (Rose Bengal stained) and dead individual foraminifera, we also determined the stable isotopic compositions of samples composed of five individuals of *C. norvangi*, *B. pacifica*, or *S. loeblichii* for comparison. We also include here the analytical results from a previous study (Ishimura et al., 2004) for *C. norvangi* (reported as *Globocassidulina* sp.). The mass of calcite in foraminiferal shells was calculated from the volume of CO_2 gas evolved during their reaction with phosphoric acid (Ishimura et al., 2004).

The $\delta^{13}\text{C}$ values of dissolved inorganic carbon (DIC) in water were determined by the method presented in Miyajima et al. (1995). The $\delta^{18}\text{O}$ in water samples was analyzed by using a wavelength-scanned cavity ring-down spectroscopy isotopic water analyzer (L2120-i; Picarro Inc., Santa Clara, CA, USA) at Hokkaido University. Analytical errors of these methods are within $\pm 0.1\%$ for $\delta^{13}\text{C}$ and $\pm 0.2\%$ for $\delta^{18}\text{O}$. $\delta^{18}\text{O}$ values of water samples are reported relative to Vienna Standard Mean Ocean Water (VSMOW).

2.3 Calculation of isotopic differences between foraminiferal shells and bottom water ($\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$)

In this study, the stable isotopic compositions of foraminiferal shells are discussed as $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$. $\Delta\delta^{13}\text{C}$ is the difference between the $\delta^{13}\text{C}$ of foraminiferal shell and the $\delta^{13}\text{C}_{\text{DIC}}$ of bottom water; $\Delta\delta^{18}\text{O}$ is the difference between the $\delta^{18}\text{O}$ of foraminiferal shell and calcite in equilibrium with bottom-water conditions ($\delta^{18}\text{O}_{\text{e.c.}}$), as discussed in previous studies (McCorkle et al., 1990; Rathburn et al., 1996; McCorkle et al., 1997; Schmiedel et

Table 1. $\delta^{13}\text{C}$ of dissolved inorganic carbon, $\delta^{18}\text{O}$ of water samples, and water temperature in the study area. Sediment intervals are ± 0.5 cm. Bottom water samples are collected with multicorers.

Site	Sea of Okhotsk (MR06-4 St.4) (44°31' N, 145°00' E)		
Water depth	1208 m		
Water depth (m)	$10^3 \times \delta^{13}\text{C}$ (VPDB)	$10^3 \times \delta^{18}\text{O}$ (VSMOW)	Temperature (°C)
30	+0.9	-0.7	-
50	+1.0	-0.6	-
75	+0.8	-0.6	-
100	+0.9	-0.6	-
200	+0.9	-0.6	-
300	+0.6	-0.6	-
500	-0.5	-0.5	-
700	-0.3	-0.4	2.2
1000	-0.3	-0.4	2.3
1158	-0.4	-0.4	2.3
Site	Sea of Okhotsk (MR06-4 St.4) (44°31' N, 145°00' E)		
Water depth	1208 m		
Depth in sediment (cm)	$10^3 \times \delta^{13}\text{C}$ (VPDB)	$10^3 \times \delta^{18}\text{O}$ (VSMOW)	Temperature (°C)
Bottom water	-0.4	-0.3	2.3
2	-0.2	-0.4	-
4	-0.3	-0.6	-
7	-0.6	-0.5	-
60	-	-0.4	-
Site	Nankai Trough (33°40' N, 136°37' E)		
Water depth	2040 m		
Depth in sediment (cm)	$10^3 \times \delta^{13}\text{C}$ (VPDB)	$10^3 \times \delta^{18}\text{O}$ (VSMOW)	Temperature (°C)
Bottom water	-1.0	-0.1	1.9
1	-1.2	-0.1	-
2	-1.3	-0.2	-
3	-2.6	-0.2	-
5	-2.9	0.0	-
7	-2.7	-0.2	-
9	-2.6	-0.3	-
Site	Sagami Bay (35°00' N, 139°14' E)		
Water depth	1182 m		
Depth in sediment (cm)	$10^3 \times \delta^{13}\text{C}$ (VPDB)	$10^3 \times \delta^{18}\text{O}$ (VSMOW)	Temperature (°C)
Bottom water	0.0	-0.2	2.9

The isotopic data from the Nankai Trough are from Toki (2004).

al., 2004; Fontanier et al., 2006; Basak et al., 2009). The $\delta^{18}\text{O}_{\text{e.c.}}$ values were calculated using equations proposed by Friedman and O'Neil (1977), which is the same procedure as in previous studies (Fontanier et al., 2006; Basak et al., 2009). The bottom-water temperatures at the study sites were 2.3 °C in the Sea of Okhotsk (Matsunaga and

Tanaka, 2006), 1.9 °C in the Nankai Trough (Hamamoto et al., 2005), and 2.9 °C in Sagami Bay.

2.4 Reliability of microscale isotopic analysis of carbonate

The reliability of isotope analysis for samples larger than 0.2 µg CaCO₃ has already been demonstrated (Ishimura et al., 2004, 2008). The same analytical procedures can be used to analyze sub-microgram quantities to several hundred micrograms of carbonate. We did not observe any evidence of added errors from foraminiferal sampling (e.g. addition of authigenic carbonate, staining by Rose Bengal, etc.) or any systematic analytical errors (e.g. leakage of air, isotopic fractionation). In addition, to clarify the possibility of CO₂ gas generation through the reaction between acid and organic materials, we checked whether foraminiferal soft tissues react with acid. We did not find the generation of CO₂ gas from soft tissue over several days of observation.

3 Results and discussion

We obtained a dataset that included (1) $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ of foraminifera and water samples, (2) weights of individual shells, (3) isotopic distributions within species (inter-individual isotopic variations), and (4) weighted average isotope values for each species (inter-species isotopic variations). The isotopic values of bottom water are shown in Table 1. All analytical results from foraminiferal shells are shown in Table 2. No systematic difference was observed in isotopic values between living (Rose Bengal stained) and dead individuals, and isotopic differences among shells collected from different depths were within the range of inter-individual isotopic deviations for each species. This means there is no noticeable relationship between individual isotopic compositions within the same species and the sediment depth at which individuals were collected. The analyzed individuals show various inter-species differences in $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$, together with species-specific inter-individual variations. The average single-shell isotopic values approximately corresponded to the average values from five shells analyzed together, confirming that inter-species differences in average isotopic values were not due to the reduced sample size (Table 2).

3.1 Characteristics of isotopic disequilibrium in benthic foraminifera: choosing more reliable species as bottom-water indicator

We calculated the difference between $\delta^{13}\text{C}$ values of foraminifera and DIC in bottom water ($\Delta\delta^{13}\text{C}$) and between $\delta^{18}\text{O}$ values of foraminifera and calcite in equilibrium with bottom water ($\Delta\delta^{18}\text{O}$). The individual foraminifera showed inter-species differences in $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$, as well as variation of inter-individual isotopic dispersions

among species (Fig. 4 and Table 2). The magnitude of inter-individual isotopic dispersion of *Bulimina aculeata* was identical within the analytical precision, and their isotopic values were almost the same as the $\delta^{13}\text{C}$ of DIC in bottom water ($\delta^{13}\text{C}_{\text{DIC}}$) and $\delta^{18}\text{O}$ of calcite in equilibrium with bottom water ($\delta^{18}\text{O}_{\text{e.c.}}$). Although *Uvigerina akitaensis* and the *Cassidulina* group (*Islandiella norcrossi*, *Cassidulina norvangi*, and *Takayanagia delicata*) showed slightly negative average $\Delta\delta^{13}\text{C}$ values, their $\delta^{18}\text{O}$ values were close to $\delta^{18}\text{O}_{\text{e.c.}}$ values, with smaller deviations (SD < 0.4 ‰, Fig. 4). These small inter-individual variations in isotopic composition and mean $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values close to seawater demonstrate the usefulness of these species for estimating the past isotopic composition of deep-sea water from even a limited number of individuals. Other species exhibited inter-individual isotopic variations substantially greater than the analytical precision. Moreover, $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values in *Brizalina pacifica* and *Stainforthia loeblichii* were extremely negative compared with previously reported values for benthic foraminifera (Grossman, 1987; Rathburn et al., 1996; McCorkle et al., 1997; Mackensen and Lincari, 2004; Fontanier et al., 2006).

All species tended to show ¹³C- and ¹⁸O-enriched $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values in proportion to individual weight (Fig. 5a). Also, the $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values were inversely proportional to shell weight, and larger (heavier) individuals tended to have isotopic values closer to $\delta^{13}\text{C}_{\text{DIC}}$ and $\delta^{18}\text{O}_{\text{e.c.}}$ (Fig. 5b, Supplement Fig. 1). The isotopic shift associated with growth stage has been reported as the “ontogenetic isotope effect” for some larger species (Schmiedl et al., 2004; Fontanier et al., 2006; McCorkle et al., 2008; Schumacher et al., 2010). We found that this trend is not limited to certain species but is common among species. Our initial findings show that species with low inter-individual deviations in isotopic composition are more suitable as direct proxies of the bottom-water environment. Moreover, the magnitude of the inter-species and inter-individual isotopic variations can be simplified to its correlation with the mass of the individual calcite shell.

3.2 Application of inter-individual isotopic variations to estimate more reliable isotopic values of bottom water in the past.

Although we found that the species with smaller inter-individual isotopic deviations are more suitable as environmental proxies, some of those species (*U. akitaensis* and the *Cassidulina* group) had carbon isotopic values that were slightly negative relative to ambient $\delta^{13}\text{C}_{\text{DIC}}$ ($\Delta\delta^{13}\text{C} \approx -2$ ‰). Therefore, we could not correctly estimate the original $\delta^{13}\text{C}_{\text{DIC}}$ values of water by using isotopic evidence from these species. However, Fig. 4 displays a trend of proportionally more negative average $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values of species with increasing inter-individual variations. We determined the relationship between average $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$

Table 2. $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of individual foraminiferal shells collected from the surface sediment at study sites, as well as the averages of multiple specimens analyzed together. $\Delta\delta^{13}\text{C}$, isotopic difference between the $\delta^{13}\text{C}$ of analyzed foraminiferal shell and the $\delta^{13}\text{C}$ of DIC ($\delta^{13}\text{C}_{\text{DIC}}$) in bottom water; $\Delta\delta^{18}\text{O}$, isotopic difference between the $\delta^{18}\text{O}$ of analyzed foraminifera and the $\delta^{18}\text{O}$ of calcite in equilibrium with bottom water ($\delta^{18}\text{O}_{\text{e.c.}}$). Results of $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ analysis for *Cassidulina norvangi* reported in a previous study (Ishimura et al., 2004) are denoted by an asterisk. Living: Rose Bengal-stained. Cmbfsf: cm below the sea floor.

Species	Depth (cmbfsf)	Number of specimens analyzed	Individual weight (μg)	$10^3 \times \delta^{13}\text{C}$ Relative to VPDB	$10^3 \times \delta^{18}\text{O}$	$10^3 \times \Delta\delta^{13}\text{C}$	$10^3 \times \Delta\delta^{18}\text{O}$
<i>Bulimina aculeata</i> (Nankai Trough, water depth 2040 m)							
<i>Bulimina aculeata</i>	0–7	1	23.12	–1.56	+3.12	–0.56	–0.24
<i>Bulimina aculeata</i>	0–7	1	21.30	–1.10	+3.13	–0.10	–0.23
<i>Bulimina aculeata</i>	0–7	1	10.93	–0.84	+2.99	+0.16	–0.37
<i>Bulimina aculeata</i>	0–7	1	12.53	–1.49	+3.14	–0.49	–0.23
<i>Bulimina aculeata</i>	0–7	1	48.33	–1.13	+3.32	–0.13	–0.04
<i>Bulimina aculeata</i>	0–7	1	20.56	–1.32	+3.21	–0.32	–0.15
<i>Bulimina aculeata</i>	0–7	1	17.57	–1.25	+3.27	–0.25	–0.10
<i>Bulimina aculeata</i>	0–7	1	5.50	–1.49	+3.20	–0.49	–0.16
Average, 1σ	0–7	8	19.98	-1.27 ± 0.24	$+3.17 \pm 0.10$	–0.27	–0.19
<i>Cassidulina norvangi</i> (Sea of Okhotsk, water depth 870 m)							
* <i>Cassidulina norvangi</i>	0–1	1	3.33	–1.91	+2.21	–1.51	–0.76
* <i>Cassidulina norvangi</i>	0–1	1	3.66	–1.70	+2.86	–1.30	–0.11
* <i>Cassidulina norvangi</i>	0–1	1	3.96	–1.42	+2.77	–1.02	–0.20
* <i>Cassidulina norvangi</i>	0–1	1	4.46	–1.59	+2.53	–1.19	–0.44
* <i>Cassidulina norvangi</i>	0–1	1	5.71	–1.50	+3.19	–1.10	+0.22
* <i>Cassidulina norvangi</i>	0–1	1	2.80	–1.90	+2.86	–1.50	–0.11
<i>Cassidulina norvangi</i>	0–1	1	4.90	–1.18	+3.08	–0.78	+0.11
* <i>Cassidulina norvangi</i>	0–1	5	4.84 (avg.)	–1.57	+3.15	–1.17	+0.18
* <i>Cassidulina norvangi</i> (living)	0–1	1	3.73	–1.79	+3.17	–1.39	+0.20
* <i>Cassidulina norvangi</i> (living)	0–1	1	4.31	–1.51	+3.07	–1.11	+0.10
* <i>Cassidulina norvangi</i> (living)	0–1	5	4.45 (avg.)	–1.66	+3.03	–1.26	+0.06
<i>Cassidulina norvangi</i> (living)	5–6	1	4.29	–1.85	+3.03	–1.45	+0.06
<i>Cassidulina norvangi</i> (living)	5–6	1	2.58	–1.92	+3.13	–1.52	+0.16
<i>Cassidulina norvangi</i> (living)	5–6	1	1.59	–1.86	+2.85	–1.46	–0.11
<i>Cassidulina norvangi</i> (living)	5–6	1	4.17	–1.12	+1.70	–0.72	–1.26
<i>Cassidulina norvangi</i> (living)	5–6	1	1.72	–1.91	+2.60	–1.51	–0.37
<i>Cassidulina norvangi</i> (living)	7–8	1	5.03	–1.58	+2.84	–1.18	–0.12
<i>Cassidulina norvangi</i> (living)	7–8	1	2.94	–2.18	+2.96	–1.78	–0.01
<i>Cassidulina norvangi</i> (living)	7–8	1	4.66	–1.79	+3.11	–1.39	+0.14
<i>Cassidulina norvangi</i> (living)	7–8	1	2.33	–2.82	+3.19	–2.42	+0.22
Average, 1σ	0–1	19	4.39	-1.61 ± 0.21	$+2.90 \pm 0.30$	–1.21	–0.07
Average, 1σ	5–6	5	2.87	-1.69 ± 0.34	$+2.80 \pm 0.57$	–1.29	–0.17
Average, 1σ	7–8	4	3.74	-2.09 ± 0.55	$+3.02 \pm 0.15$	–1.69	+0.06
Total average, 1σ	0–8	28	4.02	-1.74 ± 0.36	$+2.87 \pm 0.37$	–1.34	–0.10
<i>Brizalina pacifica</i> (Sea of Okhotsk, water depth 870 m)							
<i>Brizalina pacifica</i>	0–1	1	1.06	–5.03	+0.63	–4.63	–2.34
<i>Brizalina pacifica</i>	0–1	1	0.99	–5.25	+0.68	–4.85	–2.29
<i>Brizalina pacifica</i>	0–1	1	0.92	–4.44	+1.37	–4.04	–1.60
<i>Brizalina pacifica</i>	0–1	1	0.74	–5.57	+1.84	–5.17	–1.13
<i>Brizalina pacifica</i>	0–1	1	0.35	–4.33	+1.46	–3.93	–1.51
<i>Brizalina pacifica</i> (living)	0–1	1	0.92	–6.08	+1.67	–5.68	–1.29
<i>Brizalina pacifica</i> (living)	0–1	5	0.68 (avg.)	–4.45	+2.15	–4.05	–0.82
<i>Brizalina pacifica</i> (living)	0–1	1	1.17	–5.65	+2.28	–5.25	–0.69
<i>Brizalina pacifica</i> (living)	0–1	1	1.02	–4.15	+1.00	–3.75	–1.97
<i>Brizalina pacifica</i> (living)	0–1	1	0.82	–4.74	+1.19	–4.34	–1.78
<i>Brizalina pacifica</i> (living)	0–1	1	0.88	–6.87	+1.81	–6.47	–1.16
<i>Brizalina pacifica</i> (living)	0–1	1	0.56	–5.75	+0.70	–5.35	–2.27

Table 2. Continued.

Species	Depth (cmbfsf)	Number of specimens analyzed	Individual weight (μg)	$10^3 \times \delta^{13}\text{C}$ Relative to VPDB	$10^3 \times \delta^{18}\text{O}$	$10^3 \times \Delta\delta^{13}\text{C}$	$10^3 \times \Delta\delta^{18}\text{O}$
<i>Brizalina pacifica</i> (living)	5–6	1	1.13	−1.93	+2.38	−1.53	−0.59
<i>Brizalina pacifica</i> (living)	5–6	1	0.81	−3.96	+2.44	−3.56	−0.52
<i>Brizalina pacifica</i> (living)	5–6	1	0.83	−3.38	+2.42	−2.98	−0.54
<i>Brizalina pacifica</i> (living)	5–6	1	0.59	−3.30	+2.05	−2.90	−0.91
<i>Brizalina pacifica</i> (living)	5–6	1	0.64	−4.16	−1.51	−3.76	−4.48
<i>Brizalina pacifica</i> (living)	7–8	1	0.67	−2.90	+1.17	−2.50	−1.80
Average, 1σ	0–1	16	0.92	-5.19 ± 0.82	$+1.40 \pm 0.57$	−4.79	−1.57
Average, 1σ	5–6	5	0.80	-3.35 ± 0.87	$+1.56 \pm 1.72$	−2.95	−1.41
Total average, 1σ	0–8	22	0.80	-4.55 ± 1.22	$+1.43 \pm 0.96$	−4.15	−1.54
<i>Stainforthia loeblichii</i> (Sea of Okhotsk, water depth 870 m)							
<i>Stainforthia loeblichii</i>	0–1	5	0.23 (avg.)	−4.26	+2.54	−3.86	−0.43
<i>Stainforthia loeblichii</i>	0–1	1	0.42	−6.65	+2.67	−6.25	−0.30
<i>Stainforthia loeblichii</i>	0–1	1	0.46	−4.15	+2.81	−3.75	−0.15
<i>Stainforthia loeblichii</i>	0–1	1	0.26	−5.19	+2.00	−4.79	−0.97
<i>Stainforthia loeblichii</i>	0–1	1	0.36	−4.16	+2.78	−3.76	−0.19
<i>Stainforthia loeblichii</i>	0–1	1	0.20	−8.24	+1.76	−7.84	−1.21
<i>Stainforthia loeblichii</i>	5–6	1	0.49	−3.53	+2.59	−3.13	−0.38
<i>Stainforthia loeblichii</i> (living)	5–6	1	0.25	−4.00	+2.41	−3.60	−0.56
<i>Stainforthia loeblichii</i> (living)	5–6	1	0.26	−5.51	−0.20	−5.11	−3.17
<i>Stainforthia loeblichii</i> (living)	5–6	1	0.41	−3.83	+2.54	−3.43	−0.43
Average, 1σ	0–1	10	0.28	-5.44 ± 1.68	$+2.43 \pm 0.44$	−5.04	−0.54
Average, 1σ	5–6	4	0.35	-4.22 ± 0.88	$+1.83 \pm 1.36$	−3.82	−1.13
Total average, 1σ	0–6	14	0.30	-4.95 ± 1.49	$+2.19 \pm 0.90$	−4.55	−0.78
<i>Brizalina pacifica</i> (Sea of Okhotsk, water depth 1208 m)							
<i>Brizalina pacifica</i>	0–1	1	0.44	−6.80	+0.05	−6.40	−2.92
<i>Brizalina pacifica</i>	1–2	1	0.61	−4.70	+1.49	−4.30	−1.48
<i>Brizalina pacifica</i>	1–2	1	0.39	−11.00	−1.59	−10.60	−4.56
<i>Brizalina pacifica</i>	1–2	1	0.76	−4.45	+1.28	−4.05	−1.69
<i>Brizalina pacifica</i>	7–8	1	0.31	−7.12	−0.77	−6.72	−3.74
Average, 1σ	0–8	5	0.50	-6.81 ± 2.63	$+0.09 \pm 1.32$	−6.41	−2.88
<i>Uvigerina akitaensis</i> (Sea of Okhotsk, water depth 1208 m)							
<i>Uvigerina akitaensis</i>	0–1	1	9.32	−1.59	+2.87	−1.19	−0.09
<i>Uvigerina akitaensis</i>	0–1	1	4.41	−1.90	+2.94	−1.50	−0.02
<i>Uvigerina akitaensis</i>	0–1	1	3.19	−1.62	+3.35	−1.22	+0.39
<i>Uvigerina akitaensis</i>	1–2	1	22.56	−0.99	+3.33	−0.59	+0.37
<i>Uvigerina akitaensis</i> (living)	5–6	1	16.55	−1.19	+3.61	−0.79	+0.64
Average, 1σ	0–6	5	11.21	-1.46 ± 0.36	$+3.22 \pm 0.30$	−1.06	+0.25
<i>Chilostomella ovoidea</i> (Sea of Okhotsk, water depth 1208 m)							
<i>Chilostomella ovoidea</i> (living)	0–1	1	3.19	−2.71	+2.63	−2.31	−0.34
<i>Chilostomella ovoidea</i>	1–2	1	6.99	−3.19	+3.58	−2.79	+0.61
<i>Chilostomella ovoidea</i> (living)	1–2	1	11.16	−2.43	+3.48	−2.03	+0.51
<i>Chilostomella ovoidea</i> (living)	1–2	1	13.98	−2.10	+3.07	−1.70	+0.10
<i>Chilostomella ovoidea</i> (living)	5–6	1	1.75	−3.56	+2.58	−3.16	−0.38
Average, 1σ	0–6	5	7.41	-2.80 ± 0.59	$+3.07 \pm 0.46$	−2.40	+0.10
<i>Islandiella norcrossi</i> (Sea of Okhotsk, water depth 1208 m)							
<i>Islandiella norcrossi</i> (living)	0–1	1	5.03	−0.93	+3.38	−0.53	+0.41
<i>Islandiella norcrossi</i> (living)	0–1	1	2.82	−1.31	+2.88	−0.91	−0.08
<i>Islandiella norcrossi</i> (living)	0–1	1	1.18	−1.76	+2.82	−1.36	−0.15
<i>Islandiella norcrossi</i> (living)	1–2	1	1.84	−1.52	+2.88	−1.12	−0.08
<i>Islandiella norcrossi</i> (living)	7–8	1	4.17	−0.88	+3.18	−0.48	+0.21
Average, 1σ	0–8	5	3.01	-1.28 ± 0.39	$+3.03 \pm 0.24$	−0.88	+0.06
<i>Globobulimina auriculata</i> (Sea of Okhotsk, water depth 1208 m)							
<i>Globobulimina auriculata</i> (living)	0–1	1	1.43	−2.07	+3.32	−1.67	+0.36
<i>Globobulimina</i> cf. <i>auriculata</i> (living)	0–1	1	168.73	−0.69	+3.73	−0.29	+0.77
<i>Globobulimina auriculata</i> (living)	1–2	1	0.72	−3.51	+2.45	−3.11	−0.52
<i>Globobulimina auriculata</i> (living)	5–6	1	1.37	−2.09	+3.43	−1.69	+0.46

Table 2. Continued.

Species	Depth (cmbsf)	Number of specimens analyzed	Individual weight (μg)	$10^3 \times \delta^{13}\text{C}$ Relative to VPDB	$10^3 \times \delta^{18}\text{O}$	$10^3 \times \Delta\delta^{13}\text{C}$	$10^3 \times \Delta\delta^{18}\text{O}$
<i>Globobulimina auriculata</i> (living)	5–6	1	0.51	–3.80	+1.93	–3.40	–1.04
Average, 1σ	0–6	5	34.55	-2.43 ± 1.26	$+2.97 \pm 0.76$	–2.03	+0.01
Average, 1σ (small individual)	0–6	4	1.01	-2.87 ± 0.92	$+2.78 \pm 0.72$	–2.47	–0.19
<i>Takayanagia delicata</i> (Sea of Okhotsk, water depth 1208 m)							
<i>Takayanagia delicata</i> (living)	0–1	1	3.43	–1.58	+3.24	–1.18	+0.27
<i>Takayanagia delicata</i> (living)	0–1	1	2.09	–2.15	+3.01	–1.75	+0.04
<i>Takayanagia delicata</i> (living)	1–2	1	2.37	–2.38	+2.61	–1.98	–0.36
<i>Takayanagia delicata</i> (living)	5–6	1	1.70	–2.52	+2.90	–2.12	–0.06
<i>Takayanagia delicata</i> (living)	7–8	1	3.07	–2.10	+2.88	–1.70	–0.08
Average, 1σ	0–8	5	2.53	-2.14 ± 0.36	$+2.93 \pm 0.23$	–1.74	–0.04
<i>Globobulimina auriculata</i> (Sagami Bay, water depth 1182 m)							
<i>Globobulimina auriculata</i> (living)	0–2	1	31.9	–3.43	+2.92	–3.41	–0.05
<i>Globobulimina auriculata</i> (living)	0–2	1	43.4	–3.65	+3.32	–3.63	+0.35
<i>Globobulimina auriculata</i> (living)	0–2	1	49.8	–1.58	+2.22	–1.56	–0.75
<i>Globobulimina auriculata</i> (living)	0–2	1	7.4	–3.25	+2.78	–3.23	–0.18
<i>Globobulimina auriculata</i> (living)	0–2	1	44.4	–2.43	+3.29	–2.41	+0.32
<i>Globobulimina auriculata</i> (living)	0–2	1	5.9	–1.81	+3.03	–1.79	+0.06
Average, 1σ		6	30.5	-2.69 ± 0.88	$+2.93 \pm 0.41$	–2.67	–0.04
<i>Chilostomella ovoidea</i> (Sagami Bay, water depth 1182 m)							
<i>Chilostomella ovoidea</i> (living)	0–2	1	4.7	–2.95	+2.89	–2.93	–0.08
<i>Chilostomella ovoidea</i> (living)	0–2	1	2.5	–5.25	+1.61	–5.23	–1.36
<i>Chilostomella ovoidea</i> (living)	0–2	1	6.5	–4.34	+2.94	–4.32	–0.03
<i>Chilostomella ovoidea</i> (living)	0–2	1	5.8	–2.49	+1.57	–2.47	–1.40
<i>Chilostomella ovoidea</i> (living)	0–2	1	5.9	–3.27	+2.73	–3.25	–0.24
<i>Chilostomella ovoidea</i> (living)	0–2	1	4.9	–4.94	+3.02	–4.92	+0.05
Average, 1σ		6	5.0	-3.87 ± 1.13	$+2.46 \pm 0.68$	–3.85	–0.51
<i>Rutherfordoides cornuta</i> (Sagami Bay, water depth 1182 m)							
<i>Rutherfordoides cornuta</i> (living)	0–2	1	4.5	–1.18	+0.55	–1.16	–2.42
<i>Rutherfordoides cornuta</i> (living)	0–2	1	6.5	–0.88	+2.92	–0.86	–0.05
<i>Rutherfordoides cornuta</i> (living)	0–2	1	3.9	–1.56	+2.47	–1.54	–0.50
<i>Rutherfordoides cornuta</i> (living)	0–2	1	7.5	–0.96	+3.07	–0.94	+0.10
<i>Rutherfordoides cornuta</i> (living)	0–2	1	5.3	–1.13	+2.58	–1.11	–0.39
<i>Rutherfordoides cornuta</i> (living)	0–2	1	3.9	–0.95	+2.61	–0.93	–0.36
Average, 1σ		6	5.3	-1.11 ± 0.25	$+2.37 \pm 0.92$	–1.09	–0.60

values and inter-individual distributions (SD within species) at the two sites in the Sea of Okhotsk (Fig. 6a and b). The regression lines all have high correlation coefficients ($r > 0.9$, $p < 0.05$), indicating that the coefficients of variation are almost constant. Note that the p-value was calculated for all $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ isotopic values of analyzed species of two stations in the Sea of Okhotsk. Moreover, the intercepts of the regression lines (at $\text{SD} = 0$) for $\delta^{13}\text{C}$ are almost identical to the $\delta^{13}\text{C}_{\text{DIC}}$ (Table 3). In foraminiferal samples collected from Sagami Bay, we found the same trend in the $\delta^{13}\text{C}$ profile (Fig. 6c), and the intercept of the regression line also corresponded to the $\delta^{13}\text{C}_{\text{DIC}}$ values in the bay. The differences between the intercept values and the actual $\delta^{13}\text{C}_{\text{DIC}}$ at these three sites are within 0.3%, substantially closer to $\delta^{13}\text{C}_{\text{DIC}}$ than the $\Delta\delta^{13}\text{C}$ values of all species.

We conclude that we can more reliably estimate $\delta^{13}\text{C}_{\text{DIC}}$ values of bottom water by accounting for the vital effect in foraminifera. This technique promises to yield $\delta^{13}\text{C}$ data for bottom water where water samples are not available. However, there is still uncertainty in $\delta^{18}\text{O}_{\text{e.c.}}$ values owing to changes of bottom-water temperature and analytical errors in measuring $\delta^{18}\text{O}$ of water. In addition, the choice of equation for calculating $\delta^{18}\text{O}_{\text{e.c.}}$ affects the results; for example there is a 0.7‰ difference in $\delta^{18}\text{O}_{\text{e.c.}}$ values at our study sites as calculated by the equations of Friedman and O'Neil (1977) and Kim and O'Neil (1997). We propose applying the magnitude of isotopic disequilibrium using the inter-individual isotopic distributions of species as an index of the reliability of bottom-water isotope values. Also the $\delta^{13}\text{C}_{\text{DIC}}$ can be reconstructed from foraminiferal isotopes without water samples

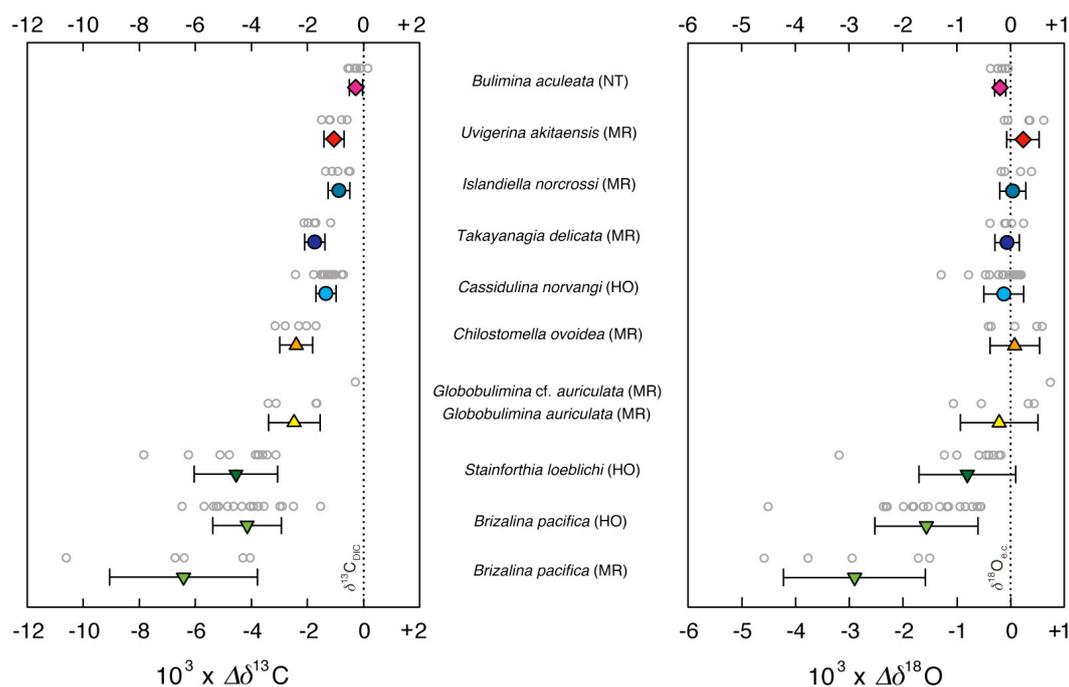


Fig. 4. Variations in $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values of individual benthic foraminifera together with average values. Open circles are individual isotopic values. Colored symbols are average values with standard deviations (error bars). Dotted lines indicate the $\delta^{13}\text{C}$ values of DIC in bottom water ($\delta^{13}\text{C}_{\text{DIC}}$) and the $\delta^{18}\text{O}$ values of calcite in equilibrium with bottom water ($\delta^{18}\text{O}_{\text{e.c.}}$). The $\delta^{18}\text{O}_{\text{e.c.}}$ values were calculated using the equation proposed by Friedman and O'Neil (1977). NT, Nankai Trough (1881 m); MR, Sea of Okhotsk (1208 m); HO, Sea of Okhotsk (870 m).

Table 3. Comparison of predicted isotopic values of calcite in equilibrium with bottom water at the study sites (intercepts of regression lines in Fig. 6) and the actual isotopic values of calcite in equilibrium with bottom water.

Site	Isotopic values of calcite in equilibrium with bottom water			
	Estimated values (intercepts of regression lines in Fig. 6)		Actual values	
	$10^3 \times \delta^{13}\text{C}_{\text{DIC}}$	$10^3 \times \delta^{18}\text{O}_{\text{e.c.}}$	$10^3 \times \delta^{13}\text{C}_{\text{DIC}}$	$10^3 \times \delta^{18}\text{O}_{\text{e.c.}}$
	Relative to VPDB		Relative to VPDB	
Sea of Okhotsk (water depth 1208 m)	-0.6	+4.0	-0.4	+3.0
Sea of Okhotsk (water depth 870 m)	-0.7	+3.6	-0.4	+3.0
Sagami Bay (water depth 1182 m)	-0.3	+3.3	-0.0	+3.0

by correcting for the isotopic disequilibrium. This permits the utilization of stored sediment samples, even those with no associated isotopic data for bottom water. In addition, we can estimate the absolute $\delta^{13}\text{C}$ changes in the ocean bottom by comparing the calibrated isotope values to the relative $\delta^{13}\text{C}$ curves previously reported, and thereby better understand environmental changes.

Although there was a good correlation between SD and $\Delta\delta^{13}\text{C}$ of benthic foraminifera in same sampling site, the empirical relations of SD and $\Delta\delta^{13}\text{C}$ may not always be the same among different environmental conditions (redox condition, flux of organic matter, bottom-water chemistry, etc.). By studying various sites, we were able to clarify the detailed

characteristics of the relationship between the magnitude of SD and $\Delta\delta^{13}\text{C}$ of benthic foraminifera.

3.3 Possible causes of inter-species and inter-individual isotopic variations

Part of the inter-individual variability in $\delta^{13}\text{C}$ may be caused by the isotopic variability of DIC in sediment pore water, which is caused by the decomposition of organic matter, and a wider range of depth habitats may result in large inter-individual variability in isotopic compositions. However, even considering the isotopic variation in sediments owing to the decomposition of organic matter and the presence of a geothermal gradient, we cannot account for the extremely

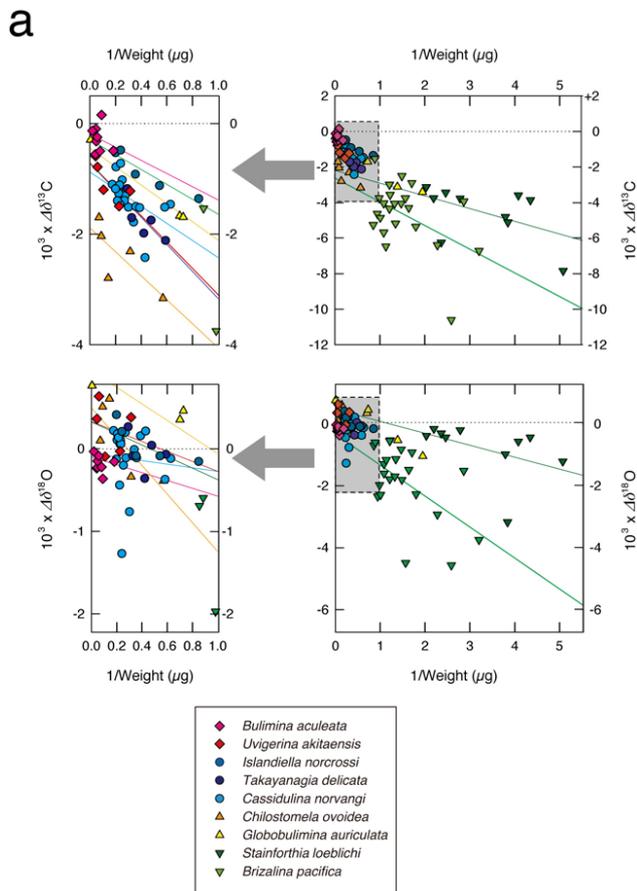


Fig. 5. Relationship between shell weight and $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values of individual benthic foraminifera. (a) $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ values of all individual benthic foraminifera analyzed plotted as a function of inverse shell weight. Lines are linear regressions fitted to each species. (b) Three-dimensional plot of $\Delta\delta^{13}\text{C}$, $\Delta\delta^{18}\text{O}$, and shell weight (animation available as Supplement Fig. 1).

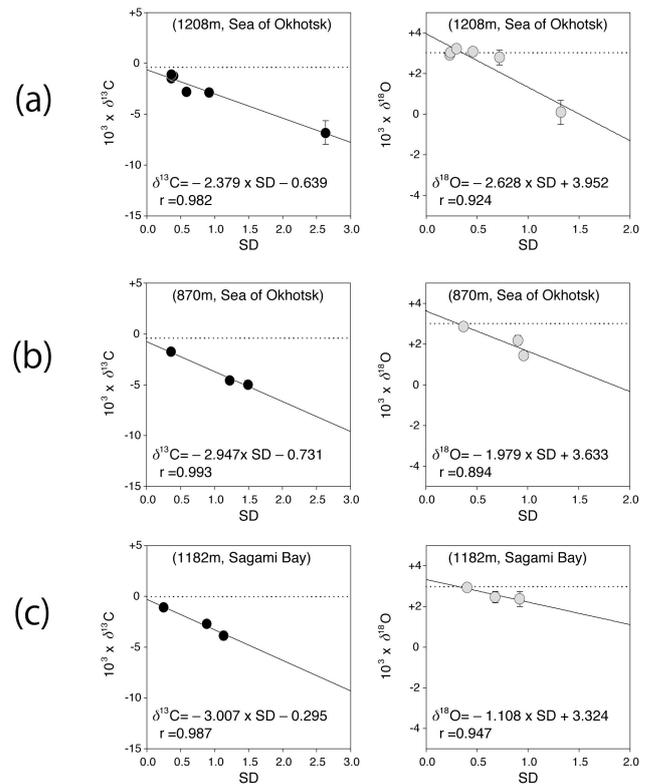


Fig. 6. Relationships between average $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values and inter-individual variations (SD; standard deviation) for each species of benthic foraminifera analyzed. (a) Sea of Okhotsk, 1208 m bottom depth; (b) Sea of Okhotsk, 870 m; (c) Sagami Bay, 1182 m. Solid lines and equations are linear regression results. Dotted lines represent $\delta^{13}\text{C}_{\text{DIC}}$ and $\delta^{18}\text{O}_{\text{e.c.}}$ at each site. Error bars indicate the range of standard error for each species.

negative isotopic values and the large inter-individual variation in $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$, especially observed in *B. pacifica* and *S. loeblichii*. The $\delta^{13}\text{C}$ values of most individuals were much lower than $\delta^{13}\text{C}_{\text{DIC}}$ values of pore water at the sediment depth of which they had been taken (Tables 1 and 2). Furthermore, the $\delta^{18}\text{O}$ variation of 6‰ among species (Fig. 4) corresponds to a temperature difference of $\pm 24^\circ\text{C}$, which is greater than the geothermal gradient. Also, we could not explain the variability of $\delta^{18}\text{O}$ on the basis of differences of microhabitat because the decomposition of organic matter does not change the $\delta^{18}\text{O}$ value of pore water. In fact, the $\delta^{18}\text{O}$ values of pore water at each sediment-depth indicate almost homogeneous isotopic values (the $\delta^{18}\text{O}$ variation among different sediment depths is almost the same as the analytical error). The isotopic variation may instead be related to the carbonate ion concentration ($[\text{CO}_3^{2-}]$) effect that is hypothesized as controlling $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ in calcifying organisms (Spero et al., 1997). $\delta^{18}\text{O}$ in planktonic foraminifera decreases with increasing $[\text{CO}_3^{2-}]$ and pH (Spero et al., 1997). We observed a similar trend in the isotopic shift among benthic species (Fig. 7). This

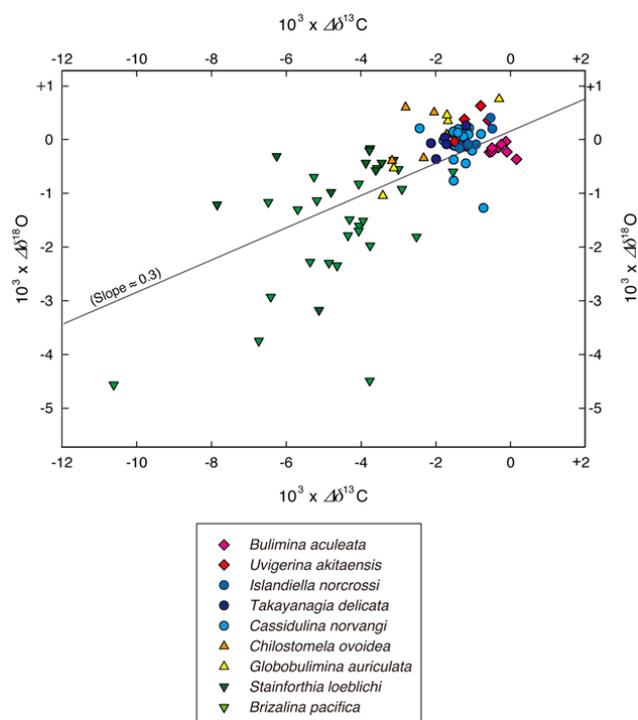


Fig. 7. Relationship between $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$ of individual benthic foraminifera. Solid line indicates the reported trend of the isotopic shift caused by the carbonate ion concentration effect (Spero et al., 1997).

similarity between laboratory experiments with planktonic foraminifera and our natural samples of benthic foraminifera suggests that the inter-individual and inter-species isotopic variations are strongly affected by the $[\text{CO}_3^{2-}]$ effect. However, infaunal species, such as *B. pacifica* and *S. loeblichii* in particular, should have positive isotopic values because of the low pH and decreased $[\text{CO}_3^{2-}]$ of ambient water in sediment resulting from decomposition of organic matter (Bemis et al., 1998). Our results show the opposite trend.

Recent findings suggest a novel explanation for the extremely negative and heterogeneous inter-individual isotopic compositions of some benthic foraminifera. One study of intracellular calcification found that a widespread strategy among benthic foraminifera is to elevate the pH at the site of calcification to promote calcite precipitation, and described a mechanism that produces a ^{13}C -depleted foraminiferal shell (de Nooijer et al., 2009). This mechanism would explain the negative shift of $\Delta\delta^{13}\text{C}$ and $\Delta\delta^{18}\text{O}$. Heterogeneity of pH in foraminiferal cells may thus be one cause of the inter-individual differences in isotopic compositions observed in this study. In addition, the inter-species variability in isotopic compositions may be caused by the magnitude of pH control by different species (or by the inter-species variability in the pool of respired CO_2 that participates in calcification).

Another study reported that some benthic foraminifera accumulate intracellular nitrate stores that are respired through

denitrification (Risgaard-Petersen et al., 2006), and nitrate pooling is reported as common in foraminifera, including species closely related to *B. pacifica* and *S. loeblichii*, from very diverse benthic marine environments (Pina-Ochoa et al., 2010). High concentrations of nitrate in such species suggest that they make semi-closed system of their cell to store the materials to use. Evidence for isotopic fractionation of $\delta^{18}\text{O}$ through denitrification has been reported previously (Sigman et al., 2005), indicating that ^{18}O -depleted nitrate is respired during denitrification. Although details of the isotopic fractionation through respiration and calcification in foraminiferal cells are not well known, its respiration through intracellular denitrification and the contributions of respired CO_2 in semi-closed system are expected to alter the isotopic composition and the magnitude of isotopic variation. Moreover, both intracellular calcification and denitrification occur at the individual level and any associated isotopic fractionation would affect individuals and not the entire population, thus potentially leading to larger inter-individual isotopic differences.

However, vital effects were originally thought to result from the incorporation of isotopically depleted carbon and oxygen compounds derived from the metabolic CO_2 pool within an organism into its shell (Grossman, 1987; Erez, 2003). It appears that these metabolic vital effects, some other vital effects (Rohling and Cooke, 1999; Mackensen, 2008), and microenvironmental heterogeneity (Mackensen et al., 1993) play roles in the observed inter-individual differences in isotopic signatures, and their intensity is reflected in the magnitude of inter-individual isotopic variations. In addition, previous studies reported that *Stainforthia fusiformis* has kleptoplasts (Bernhard and Alve, 1996), and that *B. pacifica* has ectobionts (Bernhard et al., 2010b; reported as *Bolivina pacifica*). As discussed in these studies, the activities of symbionts might more or less affect the isotopic compositions of those species. Further in situ biological observations and culture experiments (e.g. Spero et al., 1997; Hintz et al., 2004; Nomaki et al., 2005; Risgaard-Petersen et al., 2006; McCorkle et al., 2008; de Nooijer et al., 2009; Pina-Ochoa et al., 2010) should help to clarify the mechanisms responsible for large inter-individual isotopic variations in foraminifera.

4 Conclusions

In this study, we show a simplified analysis of the magnitude of isotopic disequilibrium in benthic foraminifera based on inter-individual isotopic variations, and its application to estimating bottom-water conditions precisely.

We analyzed isotopes in individual foraminifera of multiple benthic species from four sites in continental margin of the northwestern Pacific Ocean to characterize the magnitude of inter-individual isotopic variations. We found that the species with low inter-individual isotopic dispersion

are more suitable as direct proxies of the bottom-water environment. Moreover, the magnitude of the inter-species and inter-individual isotopic variations can be simplified to its correlation with the mass of the individual calcite shell.

Comparing the isotopic values of benthic foraminifera collected from the same sampling site, we found that the magnitude of the isotopic disequilibrium in each species (inter-species isotopic variations) was correlated with the inter-individual isotopic dispersion. Furthermore, by using the simplified interpretation of the isotopic disequilibrium established in this study, we can reconstruct $\delta^{13}\text{C}$ values of dissolved inorganic carbon in bottom water from foraminiferal isotopic compositions, by correcting for the isotopic disequilibrium. By studying various sites, we will be able to clarify the more detailed characteristics of the relationship between the magnitude of inter-individual isotopic dispersions and $\Delta\delta^{13}\text{C}$ of benthic foraminifera.

Appendix A

Faunal reference list

***Bulimina aculeata* d'Orbigny, 1826** [Fig. 3b]. *Bulimina aculeata* d'Orbigny, 1826, p. 269; Parker, Jones and Brady, 1871, p. 172, pl. 11, Fig. 128.

***Brizalina pacifica* (Cushman and McCulloch, 1942)** [Fig. 3j] *Bolivina acerosa* var. *pacifica* Cushman and McCulloch, 1942, p. 185, pl. 21, Figs. 2–3.

***Cassidulina norvangi* Thalmann, 1952** [Fig. 3d] *Cassidulina islandica* Nørvang forma *minuta* Nørvang, 1945, p. 43, text-Figs. 8a–c (fide Ellis and Messina, 1940 et seq.). *Cassidulina islandica* Nørvang var. *norvangi* Thalmann in Phleger, 1952, p. 83, footnote 1; Phleger, 1952, pl. 14, Fig. 30. *Cassidulina norvangi* Thalmann, Nomura, 1983b, p. 53, pl. 4, Figs. 12, 13; pl. 23, Figs. 10–12.; pl. 24, Figs. 1–3.

***Chilostomella ovoidea* Reuss, 1850** [Fig. 3c] *Chilostomella ovoidea* Reuss, 1850, p. 380., pl. 48, Figs. 12a–e.

***Globobulimina auriculata* (Bailey)** [Fig. 3g] *Bulimina auriculata* Bailey, 1851, p. 12, Figs. 25–27. *Globobulimina auriculata* (Bailey); Asano, 1958, p. 9, pl. 2, Figs. 1–3.

***Islandiella norcrossi* (Cushman)** [Fig. 3f] *Cassidulina norcrossi* Cushman, 1933, p. 7, pl. 2, Fig. 7a–c. *Islandiella norcrossi* (Cushman, 1933), Nørvang, 1958 (part), p. 32, pl. 7, Figs. 8, 9, 11, pl. 8, Fig. 14 (not pl. 7, Figs. 10, 12, 13).

***Rutherfordoides cornuta* (Cushman)** [Fig. 3h] *Virgulina cornuta* Cushman, 1913, p. 637, pl. 80, Figs. 1a–c; *Rutherfordoides cornuta* (Cushman, 1913), McCulloch, 1977, p. 249, pl. 105, Fig. 5. (Generic name was originally *Rutherfordia*, but it has been replaced by the same author (McCulloch, 1981, p. 6)).

***Stainforthia loeblichii* (Feyling-Hanssen)** [Fig. 3i] *Virgulina loeblichii* Feyling-Hanssen, 1954, p. 191, pl. 1, Figs. 14–18; text-Fig. 8; *Stainforthia concave* var., *loe-*

blichii (Feyling-Hanssen), Feyling-Hanssen, 1973, p.123, pl. 5, Figs. 10.

***Takayanagia delicata* (Cushman, 1927)** [Fig. 3e] *Cassidulina delicata* Cushman, 1927, p. 168, pl. 6, Fig. 5.; *Takayanagia delicata* (Cushman); Nomura, 1983a, p. 53, pl. 1, Figs. 3a–c; pl. 7, Figs. 1–5.

***Uvigerina akitaensis* Asano, 1950** [Fig. 3a] *Uvigerina akitaensis* Asano, 1950, pt. 2, p. 14, Figs. 60–62.

Supplementary material related to this article is available online at: <http://www.biogeosciences.net/9/4353/2012/bg-9-4353-2012-supplement.zip>.

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