Temperature characteristics of bacterial sulfate reduction in continental shelf and slope sediments

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Abstract. The temperature responses of sulfate-reducing microbial communities were used as community temperature characteristics for their in situ temperature adaptation, their origin, and dispersal in the deep sea. Sediments were collected from a suite of coastal, continental shelf, and slope sediments from the southwest and southeast Atlantic and permanently cold Arctic fjords from water depths ranging from the intertidal zone to 4327 m. In situ temperatures ranged from 8 °C on the shelf to −1 °C in the Arctic. Temperature characteristics of the active sulfate-reducing community were determined in short-term incubations with 35S-sulfate in a temperature gradient block spanning a temperature range from 0 to 40 °C. An optimum temperature (T\text{opt}) between 27 °C and 30 °C for the South Atlantic shelf sediments and for the intertidal flat sediment from Svalbard was indicative of a psychrotolerant/mesophilic sulfate-reducing community, whereas T\text{opt} ≤ 20 °C in South Atlantic slope and Arctic shelf sediments suggested a predominantly psychrophilic community. High sulfate reduction rates (20–50 %) at in situ temperatures compared to those at T\text{opt} further support this interpretation and point to the importance of the ambient temperature regime for regulating the short-term temperature response of sulfate-reducing communities. A number of cold (< 4 °C) continental slope sediments showed broad temperature optima reaching as high as 30 °C, suggesting the additional presence of apparently mesophilic sulfate-reducing bacteria. Since the temperature characteristics of these mesophiles do not fit with the permanently cold deep-sea environment, we suggest that these mesophilic microorganisms are of allochthonous origin and transported to this site. It is likely that they were deposited along with the mass-flow movement of warmer shelf-derived sediment. These data therefore suggest that temperature response profiles of bacterial carbon mineralization processes can be used as community temperature characteristics, and that mixing of bacterial communities originating from diverse locations carrying different temperature characteristics needs to be taken into account to explain temperature response profiles of bacterial carbon mineralization processes in sediments.

1 Introduction

About 95 % of the seafloor is permanently cold with in situ temperatures below 4 °C (Levitus and Boyer, 1994). Bacteria carrying out carbon mineralization in the cold seabed must be adapted to operate effectively under such low temperatures. Hence, psychrophilic bacteria with suitable cell membrane fluidity and cold-adapted enzymes are particularly abundant in the deep sea (Margesin and Miteva, 2011). As a result of such microbial adaptation to low temperature, the rate and efficiency of organic carbon mineralization in the cold may be as high as in temperate and warm habitats (Kostka et al., 1999). However, bacteria that grow at temperatures extending into the mesophilic range and that do not show cold adaptations were also isolated from the cold deep seafloor (Rüger, 1989; Rüger and Tan, 1992; Finster and Bak, 1993; Chen et al., 2003; Aono et al., 2010).

Thermal adaptation and flexibility of microorganisms to function effectively in different temperature regimes is related to underlying molecular mechanisms, e.g. to the presence of a large number of genes induced at low
temperatures and cold-adapted proteins (Ting et al., 2011; Casanueva et al., 2010). Phenotypically these molecular adaptations are expressed as different cardinal temperatures of growth or respiration, i.e. a temperature minimum \( T_{\text{min}} \), maximum \( T_{\text{max}} \) and optimum \( T_{\text{opt}} \), the \( T_{\text{opt}} \) indicating the temperature of the highest rate. By definition, psychrophilic bacteria have a temperature minimum \(< 0^\circ\text{C}\), optimum \(\leq 15^\circ\text{C}\), and maximum \(\leq 20^\circ\text{C}\). Psychrotolerant bacteria have a minimum \(\leq 0^\circ\text{C}\), optimum \(\leq 25^\circ\text{C}\), and maximum \(\leq 35^\circ\text{C}\). Mesophilic bacteria have a minimum \(> 0^\circ\text{C}\), optimum \(25–40^\circ\text{C}\), and maximum \(35–40^\circ\text{C}\) (Morita, 1975).

Bacterial sulfate reduction is the predominant anaerobic carbon mineralization pathway in most continental shelf and slope sediments (Kasten and Jørgensen, 2006) and is also detected in sediments of the continental rise and the abyssal plains (Ferdeman et al., 1999; D’Hondt et al., 2002; Lee et al., 2008). Sulfate reduction rates can be quantified with high sensitivity and precision by incubating sediment with \(^{35}\text{S}\)-sulfate and measuring the rate of \(^{35}\text{S}\)-sulfide formation (Fossing, 1995; Kallmeyer et al., 2004). This technique is one of the few that link the quantification of organic carbon mineralization rates to a physiologically defined group of microorganisms (Leloup et al., 2007; Lee et al., 2008).

Incubation of sediments or of pure cultures in a temperature gradient has been used to determine the cardinal temperatures for growth or respiration of microbial communities (Batley, 1964). Temperature–activity curves possess multiple properties that define the temperature characteristics of a community. Rates of metabolism at in situ temperatures relative to the rates at \( T_{\text{opt}} \) indicate how well bacteria perform under ambient low temperatures (Knoblauch and Jørgensen, 1999). The Arrhenius plot can be used for a graphical representation of the temperature dependence of bacterial metabolism (Arrhenius, 1908). Thereby, the logarithm of the rate of bacterial respiration or growth versus the inverse absolute temperature yields a linear relationship in the temperature range where the bacteria are well adapted, whereas deviation from a linear slope at the upper or lower extreme expresses the inability of Sulfate Reducing Bacteria (SRB) to maintain a well-controlled metabolic activity and may indicate decoupled electron flow that maintains local high sedimentation rates (Arrhenius, 1908). The Arrhenius plot can be used for a graphical representation of the temperature dependence of bacterial metabolism (Arrhenius, 1908).

### 2 Material and methods

Sediments from the South Atlantic were collected in 2008 and 2009 at six stations located on the shelf and slope off central Namibia and off Uruguay and Argentina, respectively. Sediments from four stations in the Arctic were collected in 1998, 1999 and 2007 from fjords and an intertidal flat on the west coast of Svalbard. Samples were taken from the sediment zone of highest sulfate reduction rates, which was typically in the depth range of 3–10 cm. Sediments were stored in gas-tight plastic bags at 4°C until further processing in the laboratory within a few days. For measurements of rates in intact sediment cores, cores of 26 mm diameter and ca. 15 cm length were taken, sealed at both ends with rubber stoppers leaving air in the headspace, and stored at 4°C. Coordinates for the study sites, in situ temperatures, and water depths at which sediments were collected are given in Table 1.

#### 2.1 Oceanography and sedimentary setting

##### 2.1.1 Namibian shelf and slope

The Benguela upwelling system has extremely high primary productivity of 767–1051 g C m\(^{-2}\) yr\(^{-1}\) and is the most productive coastal upwelling area on Earth (Carr, 2002). The bottom water temperatures of the Namibian shelf vary annually between 7 and 10°C on the shelf and between 1 and 4°C on the slope (Lass and Mohrholz, 2005). The shelf and slope system is characterized by seaward and downslope particle transport that maintains local high sedimentation rates in a depocenter at 1000–1500 m water depth (Inthorn et al., 2005). Sediments were collected from two stations during the RV Meteor cruise M76/1 (Zabel and Ferdelman, 2008), one on the shelf in 130 m water depth and one on the continental slope in 2000 m water depth. Sediments from the shelf and from the slope have high concentrations of organic carbon...
of up to 20% dry weight that decrease to about 7–8% in the depocenter on the continental slope (Inthorn et al., 2005). This decrease corresponds well with the decrease in sulfate reduction rates (SRR) with increasing water depth (Table 2, Brüchert et al., 2003).

### 2.1.2 SW Atlantic margin and basin

The region off Uruguay and Argentina has dynamic oceanographic conditions due to the confluence of two different water masses that causes high primary productivity and high deposition of organic matter (e.g. Behrenfeld and Falkowski, 1997). Sediments from the SW Atlantic were collected during RV Meteor cruise M78/3 on the continental shelf off Uruguay and on the slope off Argentina (Krastel and Wefer, 2009). Shelf bottom water temperatures vary between 7 and 10°C, and continental slope bottom water temperatures are between 1 and 4°C (Hansen et al., 2003; Ortega and Martínez 2007). This region is characterized by high sedimentation rates, gravity mass flows due to major turbidities and slides, and strong surface currents (Riedinger et al., 2005). The sediments in the study area are characterized by low carbonate concentrations and high concentrations of organic carbon and iron oxides (Hansen et al., 2003). SRR and total organic carbon (TOC) contents are of the same magnitude as shelf sediments off Namibia (Table 2). The TOC contents of Argentine slope sediments decrease with water depth from about 5% to 1% dry weight.

### 2.1.3 West coast of Svalbard

Along the west coast of Svalbard, primary productivity is controlled by light availability, ice cover, and the influence of the West Spitsbergen current, which transports relatively warm water (0.5 to 2°C) from the North Atlantic to this latitude (Carmack et al., 2006). These Arctic sediments have very constant temperatures throughout the year, ranging from −1°C to +2°C. Intertidal mud flats on the west coast of Svalbard have summer temperatures as high as 6°C, but drop to −20°C during the winter (Nordli, 2005). The mean annual primary production along the west coast is around 120 g C m⁻² yr⁻¹ (Sakshaug, 2004). SRR are also relatively high and comparable to rates of many temperate shelf areas (Sagemann et al., 1998). Sediment was collected from four stations along the west coast of the main island of the Svalbard archipelago, Spitsbergen. Three stations were located centrally in fjords while the fourth was on an intertidal mud flat. The fjord sediments were taken in July 1998 and July 1999 with a Haps corer while the intertidal flat was sampled in August 2008 from the shore.

### 2.2 Sulfate reduction rate measurements

Sulfate reduction rates (SRR) were measured in two parallel sediment cores using the whole core incubation method by Jørgensen (1978). These data are henceforth termed in situ SRR. 5 µl of carrier-free $^{35}$SO$_4^{2-}$ tracer solution in 4% NaCl (~100 kBq per injection) was injected at 1 cm intervals to a depth of 16 cm. Incubation time was 8 h at the in situ temperature. Samples that were collected in 1998 and 1999 were distilled by the hot chromium reduction method described by Fossing and Jørgensen (1989). Samples collected in 2008 and 2009 were analyzed using the low-blank cold chromium distillation method described by Kallmeyer et al. (2004). Briefly, centrifuged sediment was diluted with 10 ml dimethylformamide and placed in a distillation flask. Total reduced inorganic sulfide (TRIS) was acid-distilled under nitrogen at room temperature after adding 12 ml 6N HCL and 12 ml 1M chromium chloride. The TRIS was recovered as zinc sulfide in traps containing 7 ml of 5% w/v zinc acetate solution and $^{35}$S was counted in a liquid scintillation counter (Packard, Tricarb 2500 TR). The scintillation cocktail was Lumasafe Plus (Lumac BV, Groningen, The Netherlands) mixed in a ratio of 2:1 (v/v) with the ZnS suspension.

### 2.3 Temperature dependence of SRR

The temperature dependence of SRR was determined in temperature gradient incubation experiments using a thermostated aluminum block (Isaksen and Jørgensen, 1996).
Table 2. Bulk geochemical analysis. Carbon and nitrogen concentrations determined for Namibian, Uruguayan, Argentine and Arctic sediments.

<table>
<thead>
<tr>
<th>Station</th>
<th>Organic carbon (wt %)</th>
<th>Nitrogen (wt %)</th>
<th>C/N</th>
<th>SRR nmol cm⁻² d⁻¹ (SR zone mean)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Namibia 130 m</td>
<td>4.4</td>
<td>0.5</td>
<td>10.5</td>
<td>65.8</td>
</tr>
<tr>
<td>Namibia 2000 m</td>
<td>6.5</td>
<td>0.9</td>
<td>10.4</td>
<td>3.59</td>
</tr>
<tr>
<td>Uruguay 244 m</td>
<td>5.0</td>
<td>0.6</td>
<td>10.0</td>
<td>43.6</td>
</tr>
<tr>
<td>Argentina 627 m</td>
<td>2.3</td>
<td>0.3</td>
<td>10.3</td>
<td>nd</td>
</tr>
<tr>
<td>Argentina 3400 m</td>
<td>1.3</td>
<td>0.2</td>
<td>9</td>
<td>nd</td>
</tr>
<tr>
<td>Argentina 4327 m</td>
<td>nd</td>
<td>nd</td>
<td>nd</td>
<td>6</td>
</tr>
<tr>
<td>Arctic 0 m</td>
<td>1.4</td>
<td>0.1</td>
<td>16</td>
<td>11</td>
</tr>
<tr>
<td>Krossfjord 80 m</td>
<td>nd</td>
<td>nd</td>
<td>nd</td>
<td>4.29</td>
</tr>
<tr>
<td>Kongsfjord 110 m</td>
<td>nd</td>
<td>nd</td>
<td>nd</td>
<td>12.6</td>
</tr>
<tr>
<td>Smeerenburgfjord 215 m</td>
<td>1.2</td>
<td>0.2</td>
<td>8.8</td>
<td>19.4</td>
</tr>
</tbody>
</table>

The temperature in the gradient block ranged from −5 °C to +40 °C with a temperature increment between each sample of 1.5 °C. On an absolute temperature scale, this temperature difference is small with regard to the temperature dependence of enzymatic responses. We do not provide a standard error for each measurement, but the degree of scatter of neighboring data points indicates the precision of the method and serves as an estimation of the analytical uncertainty and sample variability. The interpolated trend for each temperature response experiment is thus statistically substantiated as the temperature trend integrates the analytical uncertainty. Hence, the sediments were not incubated in replicates, with the exception of the Argentine sediment from the 3400 m depth. Sediment slurries were prepared from Na₂SO₄ and sealed with butyl rubber stoppers. The Hungate tubes were immediately placed in the temperature gradient block and pre-incubated to allow the sediments to reach thermal equilibrium. Then ³⁵S-labeled carrier-free sulfate (100 kBq final activity) was injected and the slurries were incubated with radiotracer. The South Atlantic sediments were pre-incubated for 12 h, and subsequently incubated with ³⁵S-labeled sulfate for 24 hours. The sediments from the three Arctic fjords (Kongsfjorden, Krossfjorden, and Smeerenburgfjorden) were pre-incubated for one hour, and then incubated for 12 h with ³⁵S-labeled sulfate for Kongsfjorden and Krossfjorden sediment, and 8 hours for Smeerenburgfjorden sediment. The Argentine sediment from the 3400 m depth. Sediment slurries were prepared from Na₂SO₄ and sealed with butyl rubber stoppers. The Hungate tubes were immediately placed in the temperature gradient block and pre-incubated to allow the sediments to reach thermal equilibrium. Then ³⁵S-labeled carrier-free sulfate (100 kBq final activity) was injected and the slurries were incubated with radiotracer. The South Atlantic sediments were pre-incubated for 12 h, and subsequently incubated with ³⁵S-labeled sulfate for 24 hours. The sediments from the three Arctic fjords (Kongsfjorden, Krossfjorden, and Smeerenburgfjorden) were pre-incubated for one hour, and then incubated for 12 h with ³⁵S-labeled sulfate for Kongsfjorden and Krossfjorden sediment, and 8 hours for Smeerenburgfjorden sediment. Incubations were stopped by transferring the sediment to 50 ml polyethylene centrifuge tubes with 20 ml 20 % zinc acetate to stop bacterial activity and to fix sulfides. Samples were kept frozen until further analysis. Subsequent distillation followed procedures described in Kallmeyer et al. (2004) for samples collected in 2008 and 2009 and in Fossing and Jørgensen (1989) for samples collected in 1998 and 1999.

2.4 Arrhenius plot and Q₁₀

Activation energies were calculated from the linear range of data in an Arrhenius-type diagram where the ³⁵S-sulfate reduction rate, k, is plotted as a function of the inverse temperature (1/T):

$$\ln(k) = \ln(A) + \left(\frac{-E_a}{R} \cdot \frac{1}{T}\right)$$  (1)

where $E_a$ is the activation energy (J mol⁻¹), $k$ is the rate of sulfate reduction (nmol cm⁻³ day⁻¹), $A$ is a constant, $R$ is the molecular gas constant (8.314 J K⁻¹ mol⁻¹), and $T$ is the absolute temperature (K).

$Q_{10}$ values between 0 °C and 10 °C were calculated according to:

$$Q_{10} = \exp\left[\frac{E_a \cdot 10}{RT(T + 10)}\right].$$  (2)

2.5 Solid phase analyses

Freeze-dried and homogenized sediment was analyzed for total carbon (TC) and total nitrogen (TN) with a Fisons NA 1500 Series 2 Elemental Analyzer. Total inorganic carbon (TIC) was measured with a CM 5240 Orbis BV coulometer. Total organic carbon (TOC) was calculated by subtracting TIC from TC.

2.6 Sulfate measurements

Porewater sulfate concentrations were determined after centrifugation of sediment at 3500 rpm in capped centrifuge tubes with nitrogen headspace at 4 °C for 15 min. Supernatant pore water (1 ml) was preserved with 200 µl 1 % (w/v) Zn-acetate solution and stored frozen at −20 °C. Sulfate concentrations were measured by suppressed ion chromatography after 1:100 dilution with deionized water on a Metrohm 761 compact ion chromatograph. Sulfate standards were prepared from Na₂SO₄, with concentrations ranging from 5 to
400 µM using an eight-point calibration curve. Quality control samples, treated as unknowns, were prepared from calibrated seawater (IAPSO) and analyzed at the start and end of every sample run sequence.

3 Results

3.1 Characterization of study sites

The organic carbon and nitrogen content, expressed as % dry weight, and the molar C:N ratio are listed in Table 2. The highest organic carbon content was measured in Namibian (4.4 % TOC) and Uruguayan (5.0 % TOC) shelf sediments, and Namibian (6.5 % TOC) slope sediment. In the other sediments, TOC ranged from 1.2 % to 2.3 % TOC. The total nitrogen (TN) concentration ranged from 0.1 % to 0.9 %. Except for the Arctic intertidal flat sediments, where the molar C:N ratio was 16, C:N ratios calculated for other sediments ranged between 8 and 10. C:N ratios near 10 determined for South Atlantic sediments are typical for sediments with high TOC content deposited under highly productive marine systems with associated high organic matter fluxes (Meyers, 1994). C:N ratios of <10 generally characterize labile organic matter easily accessible to microorganisms.

Average rates of in situ SRR in the zone of highest sulfate reduction (top 3 to 9 cm) are presented in Table 2. Highest rates were found in the Namibian shelf sediment, 65.8 nmol cm⁻³ d⁻¹, and Uruguayan shelf sediment, 43.6 nmol cm⁻³ d⁻¹. The lowest SRR was found in Argentinean sediment from 3400 m water depth, 6 nmol cm⁻³ d⁻¹, in the Arctic fjord sediment from Krossfjorden, 4.29 nmol cm⁻³ d⁻¹, and in Namibian slope sediment from 2000 m water depth, 3.59 nmol cm⁻³ d⁻¹.

3.2 Temperature dependence of SRR, South Atlantic

The SRR measured in slurried sediments in the temperature gradient block are not representative of in situ rates, yet they are clearly related to ambient temperatures, water depth, and the availability of organic matter. In situ temperatures of Namibian and Uruguayan shelf sediments define a temperate environment conducive for mesophilic to psychrotolerant microorganisms. This is supported by the temperature curves of sulfate reduction in the shelf sediments (Fig. 1a, b), and the three cardinal temperatures, $T_{\text{min}}$, $T_{\text{opt}}$, and $T_{\text{max}}$ (Table 3). In both shelf sediments $T_{\text{opt}}$ was between 25 and 30 °C. SRR in Namibian slope sediments were 7 nmol cm⁻³ d⁻¹ at the in situ temperature and increased to 72 nmol cm⁻³ d⁻¹ at $T_{\text{opt}}$. In the Uruguayan sediment SRR were 3 nmol cm⁻³ d⁻¹ at the in situ temperature and increased to 16 nmol cm⁻³ d⁻¹ at the $T_{\text{opt}}$.

In situ temperatures in the Namibian and Argentinean slope sediments range annually between 1 °C and 4 °C (Siedler, 1996). In these sediments the $T_{\text{opt}}$ of sulfate reduction was distinctly lower than in the corresponding shelf sediments. The highest SRR were found at 22 °C and 20 °C (Fig. 1c, d, f), and the $T_{\text{max}}$ was near 30 °C. A $T_{\text{opt}}$ of only 12 °C was found in one of the sediments from the Argentinean slope in 3400 m water depth (Fig. 1e). This sediment showed a broad temperature profile between 0 °C and 15 °C, but above 20 °C SRR dropped below the detection limit.
3.3 Temperature dependence of SRR, Arctic

In the Arctic fjord sediments, the highest rates were measured in Smeerenburgfjorden (Fig. 1j). In this fjord, the SRR were 53 nmol cm⁻³ d⁻¹ at the in situ temperature and increased to 200 nmol cm⁻³ d⁻¹ at the \( T_{opt} \). In the Arctic intertidal mud flat (Ymerbukta) and in Kongsfjorden sediment, SRR were about 40% lower than at the \( T_{opt} \). In Ymerbukta sediment, rates increased from 4 at \(-4^\circ C\) to 120 nmol cm⁻³ d⁻¹ at \( T_{opt} \). In the Arctic intertidal mud flat (Ymerbukta) and in Kongsfjorden sediment, they increased from 27 at \(-4^\circ C\) to 141 nmol cm⁻³ d⁻¹ at the \( T_{opt} \). In the other Arctic fjord sediments, rates increased from 4 to 44 nmol cm⁻³ d⁻¹ at \( T_{opt} \) (Fig. 1h).

The temperature profile of SRR in Ymerbukta increased in the temperature range from \(-4^\circ C\) to \( 30^\circ C\) and dropped to near the detection limit at a \( T_{max} \) of 35°C (Fig. 1g). In Smeerenburgfjorden and Krossfjorden sediment, the temperature response curve was broad and increased from \(-4^\circ C\) to \( 27^\circ C\) and \( 23^\circ C\) with a \( T_{max} \) at 40 and 34°C, respectively (Fig. 1g, h). Likewise, in Kongsfjorden sediment, SRR showed a psychrotolerant response as activity increased from \(-4^\circ C\) to a maximum at \( 18^\circ C\), but was barely detectable above \( 34^\circ C\).

3.4 Arrhenius plots and \( Q_{10} \)

For all stations the relative metabolic rates at in situ temperatures (compared to those at \( T_{opt} \)) ranged between 9% and 50% (Table 3, Fig. 2). Activation energies calculated from the Arrhenius plots ranged from 25 to 55 kJ mol⁻¹ while the \( Q_{10} \) factors were in the range of 2 to 3 (Table 3). In Namibian sediments activation energies ranged from 33 to 55 kJ mol⁻¹ and in Argentine sediments from 26 to 55 kJ mol⁻¹.

4 Discussion

4.1 Temperature response of SRR depends on in situ \( T^\circ C\) and on water depth

The most important property of temperature–SRR curves is the relative rate at in situ temperature in relation to the rate at \( T_{opt} \) (Knoblauch and Jørgensen, 1999). High rates at in situ temperature reflect a better adaptation of a process to the ambient temperature. The \( T_{opt} \) is an additional characteristic of a metabolic process that is dependent on the ambient temperature. In permanently cold sediments the \( T_{opt} \) of respiration and growth is always higher than the environmental temperatures the microorganisms experience (Feller, 2007).

Therefore, the \( T_{opt} \) itself is not a good indicator of a community’s adaptation to temperature, because the \( T_{opt} \) more reflects the threshold temperature above which enzymatic functions rapidly degrade.

Our results support the basic notion that the ambient temperature regime of marine environments selects for microbial populations with the best physiology for the respective environment. The permanently cold fjord sediments of the Arctic thus appear to host predominantly psychrophilic to psychrotolerant populations while the SRR–temperature profile in temperate sediments reflects more mesophilic populations (cf. Isaksen et al., 1994; Sagemann et al., 1998; Isaksen and Jørgensen, 1996).

In the three Arctic deep-fjord sediments with bottom water temperatures permanently near \( 0^\circ C\), temperature selected for well-adapted SRB communities, because the relative respiration rates at in situ temperatures were high compared to the rates at \( T_{opt} \). The Arctic intertidal flat sediment of Ymerbukta on Svalbard has seasonal temperature fluctuations between \(-20^\circ C\) and \(+9^\circ C\), which should favor a community with mixed temperature characteristics. This interpretation is in accordance with the lower relative respiration rates at in situ temperatures and the higher \( T_{opt} \) than in the three Arctic permanently cold deep-fjord sediments. Likewise, the
Fig. 2. Arrhenius plots of data in Fig. 1.

relatively warm bottom water temperatures of the Uruguayan and Namibian shelf of 7–10 °C are consistent with relatively high $T_{\text{opt}}$ of 25 to 30 °C and therefore suggest a greater abundance of mesophilic populations in these sediments (Fig. 1a, b, j).

In the South Atlantic sediments, the $T_{\text{opt}}$ and $T_{\text{max}}$ of sulfate reduction also decreased along with the decrease in in situ temperature at greater water depths (Fig. 1b, c, d, f), but the temperature characteristics differed between the slope sediments. The most extreme demonstration of the temperature selection for an adapted microbial community that is reflected by all properties of the temperature–SRR curve is seen in the Argentine sediment from 3400 m depth, where we measured a $T_{\text{opt}}$ of only 12 °C after 36 h of incubation. Such a low $T_{\text{opt}}$ in combination with a high relative respiration rate of 50 % at in situ temperature suggests the presence of a well-adapted psychrophilic community. A similar low $T_{\text{opt}}$ of 12.5 °C for sulfate reduction was observed in sediment from Antarctica in the Weddell Sea (Isaksen and Jørgensen, 1996). The authors incubated sediment for 8 days and reasoned that this allowed the growth of sulfate-reducing bacteria. The observed $T_{\text{opt}}$ therefore likely reflected the growth rate optimum of a predominantly psychrophilic community (Isaksen and Jørgensen, 1996).

The incubations in our studies were short (max. 36 h) and the SRR therefore reflect the short-term enzymatic capabilities of the in situ sulfate-reducing community. Prolonged incubations in similar experiments have shown that the temperature optimum decreases over time (Brüchert et al., 2001; Finke and Jørgensen, 2008). It is apparent that sulfate-reducing microorganisms from these cold sediments maintain high activity at the highest temperatures only for a limited time. Instead, relative respiration rates of an adapted community remain constant for a longer time even when sediment is exposed to higher temperature (Robador et al., 2009). The temperature–sulfate reduction rate diagrams shown in our study should therefore not be misunderstood as indicators of the true $T_{\text{opt}}$ of the sulfate-reducing community, but in combination with $T_{\text{min}}$, $T_{\text{max}}$ and relative respiration rates should be regarded as temperature characteristics of the existing populations. These community temperature characteristics allow comparisons with populations in other sediments and, along with additional verifying methods, temperature response profiles may be indicative of the origins of these populations and their temperature characteristics.

4.2 SRR in shelf and slope sediments fall into three temperature groups

In the South Atlantic and Arctic sediments, the multiple parameters of the temperature–SRR diagrams imply the presence of mesophilic, psychrotolerant, and psychophilic SRBs. The high percentage of SRR at the in situ temperature and the low $T_{\text{opt}}$ measured for the Argentinian slope sediment from 3400 m depth indicate the predominance of a psychrotolerant and psychrophilic community (Fig. 1e). In other sediments, these distinct characteristics are more obscured and suggest the existence of mixed populations of mesophilic and/or psychrotolerant SRBs.

In the Namibian shelf and slope sediments, the rates at in situ temperature were low and less than 10 % of the rates at $T_{\text{opt}}$ (Table 3), suggesting a poorly adapted sulfate-reducing community. By contrast, the $E_a$ values of 38 and 55 kJ mol$^{-1}$ suggested that SRB were well adapted to the ambient temperature. The corresponding $Q_{10}$ values for SRR were around 2, which is typical for experimental studies from many marine environments (Kirchman et al., 2009). These low $Q_{10}$ values were found repeatedly for metabolic processes in Arctic and Antarctic sediments and imply a microbial community well adapted to ambient temperature. $E_a$ values reported for
permanently cold sediments vary in the range of 25–31 kJ mol⁻¹, depending on whether the community is more psychrotolerant or psychrophilic. Our $E_a$ values were similar to those measured for sulfate reduction, denitrification and anammox in Arctic sediments from eastern Greenland and western Svalbard fjords (Fig. 2, Table 3) (Rysgaard et al., 1998; Gihring et al., 2010).

In the three Arctic deep-fjord sediments, the sulfate-reducing communities also had some inconsistent temperature characteristics: Although the relative respiration rates were all close to 20%, the $T_{\text{opt}}$ varied between 18 and 27°C. The relative respiration rates classify the temperature response as psychrophilic and psychrotolerant, but the combination of a relative respiration rate of 20% and a high $T_{\text{opt}}$ of 27°C suggests a temperature response of SRBs from the intertidal sediment that is psychrotolerant to mesophilic.

The temperature response characteristics of some South Atlantic slope sediments also differed in terms of $T_{\text{opt}}$ and broadness of the curves. $T_{\text{opt}}$ and low relative SRR of Argentinean (4327 m) and Namibian slope (2000 m) sediment suggest the presence of psychrotolerant bacteria, but also the additional presence of mesophilic populations. We hypothesize that these differences can be attributed to the strength of the influx of a seed population from warmer environments – in this case warmer shelf locations. This could explain the mesophilic signature of microorganisms derived from the shelf by sediment transport (Fig. 1c, d, f).

4.3 Sediment transport effects on experimentally-determined temperature–activity relationships

Passive movement of microorganisms with currents and water masses is an important mechanism of marine microbial dispersal (Martiny et al., 2006). Benthic bacteria in surface sediments can be suspended to the water column due to physical disturbance by near-bed currents and macroorganisms (Queric and Soltwedel, 2007; Schauer et al., 2010). Sediment transport and hydrographic conditions could explain the 9°C difference in $T_{\text{opt}}$ in three nearby deep-fjord Arctic sediments. Smeerenburgfjord, at the northwestern tip of Spitzbergen, is a narrow protected coastal strait open to the Fram Strait in the west and towards the Arctic Ocean in the north. It is strongly affected by warm, Atlantic-derived water that is constantly supplied from the south and flushed to the north (Caroll et al., 2011). This process might affect benthic conditions and contribute to the import of less-adapted microorganisms with higher $T_{\text{opt}}$ or even displace well-adapted microorganisms. The other two sites were located deep within the Kongsfjorden and Krossfjorden. Kongsfjorden is an open fjord and creates, along with Krossfjorden, a system that merges before opening into Fram Strait. Both coastal and Atlantic water flows into it and there is significant freshwater and sediment influx in the summer from melting snow, precipitation, glacial calving and runoff (Cottier et al., 2005). This material may also transport an allochthonous microbial community with higher temperature characteristics that leaves an imprint in the form of different $T_{\text{opt}}$.

In the South Atlantic slope sediments, resuspension and dispersion of suboptimally adapted microorganisms may inherit a mesophilic signature of a shelf SRB community to these deep-sea sediments. Organic matter from the coastal Benguela upwelling system on the Namibian shelf undergoes continuous suspension and redeposition leading to a net downslope transport. The shelf material accumulates in depocenters at 1000–1500 m water depth, where the sediment is rich in organic matter (Inthorn et al., 2005, 2006). The slope sediments off Uruguay and Argentina are also characterized by dynamic depositional conditions (Riedinger et al., 2005). Attempts have been made to resolve the contribution of shelf-derived material through benthic nepheloid layer transport by means of $^{230}$Th flux measurements and $^{14}$C radiocarbon analyses for the Namibian and Argentinian slope (Mollenhauer et al., 2006, 2008). The interpretations presented in these studies emphasize the temporal and spatial variability in lateral export intensity. Such variability likely affects the proportional mixtures of microorganisms with different temperature adaptations.

We suggest that the downslope transport of sediment material from the warmer continental shelf to the permanently cold continental slope may seed deep-water sediments with mesophilic organisms. Anaerobic sulfate-reducing bacteria face the problem of toxic exposure to oxygenated water during transport. However, a number of SRB have been shown to develop molecular strategies to remove oxygen or to use oxygen temporarily for respiration to reduce its toxicity (Cypionka, 2000; Dolla et al., 2006). In addition, mass transport often occurs in the form of aggregates, whose insides can be anoxic (Ploug et al., 1997).

Mesophilic and psychrotolerant SRB deposited on cold slope sediments would appear less competitive compared to their autochthonous psychrophilic counterparts (Knoblauch and Jørgensen, 1999; Robador et al., 2009). For long-term survival and success, rate of growth and growth yield are true measures of microbial adaptation to environmental temperatures. Environmental conditions in organic matter-limited deep-sea sediment impose nutritional constraints that can impede microbial growth. It is conceivable that psychrotolerant bacteria transported downslope have adaptive mechanisms to maximize and maintain a high growth yield at low temperature (Bakermans and Nealson, 2004), but it is not clear how allochthonous, mesophilic SRB derived from the shelf may proliferate when relocated to greater depths. It is noteworthy, however, that in laboratory studies both psychrophilic and mesophilic sulfate-reducing bacteria have shown comparable growth rates when grown on lactate (Knoblauch and Jørgensen, 1999, Sass et al., 1998), suggesting that proliferation of mesophiles in permanently cold environments is conceivable. The hypothesis that allochthonous mesophilic SRB are transported from the shelf and remain viable in permanently cold sediment is also supported by sulfate reduction.
rates measured in intertidal temperate sediment that showed a characteristic mesophilic temperature response even when the sediment was stored for two years at 0 °C, not changing the temperature response very much over this period of time (Robador et al., 2009). Clearly, despite the low ambient temperatures, the mesophilic sulfate-reducing community proliferated (Robador et al., 2009).

5 Conclusions

$T_{\text{min}}$, $T_{\text{max}}$, and $T_{\text{opt}}$ for SRR in various continental shelf and slope sediments from the Atlantic and Arctic Oceans indicate the coexistence of different thermal groups of SRB. This observation has implications for predicting the likely outcome of long-term temperature changes of marine sediments in Arctic shelf and deep-sea habitats. The seeding of permanently cold deep-sea habitats with mesophilic microorganisms must affect growth rates of the bacterial community at changed bottom water temperatures and can thus potentially affect the efficiency of organic carbon mineralization rates. The existing evidence from this study and from previous comparisons of carbon mineralization in permanently cold and temperate habitats points to the availability of reactive organic matter as the overriding limiting factor for long-term carbon mineralization in marine sediment (Kostka et al., 1998). Incubation experiments of SRR over two years (Robador et al., 2009) and recent experimental studies from lake sediment (Gudasz et al., 2010) suggest, however, a more prominent role of temperature limitation. These different observations can be reconciled if one considers that an apparent temperature limitation effect may be temporary and is alleviated by community replacement with better temperature adaptations. The regional comparison indicates that over longer time scales, bacteria with the best physiological adaptation to exploit the available conserved energy in the respective environment will likely prevail. In the future it will be important to explore at what time scale such a community change to the best-adapted community occurs.

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