N and P as ultimate and proximate limiting nutrients in the northern Gulf of Mexico: implications for hypoxia reduction strategies

Katja Fennel and Arnaud Laurent
Department of Oceanography, Dalhousie University, 1355 Oxford Street, Halifax B3H 4R2, Nova Scotia, Canada

Correspondence: Katja Fennel (katja.fennel@dal.ca)

Received: 2 November 2017 – Discussion started: 6 November 2017
Revised: 10 April 2018 – Accepted: 4 May 2018 – Published: 24 May 2018

Abstract. The occurrence of hypoxia in coastal oceans is a long-standing and growing problem worldwide and is clearly linked to anthropogenic nutrient inputs. While the need for reducing anthropogenic nutrient loads is generally accepted, it is costly and thus requires scientifically sound nutrient-reduction strategies. Issues under debate include the relative importance of nitrogen (N) and phosphorus (P) as well as the magnitude of the reduction requirements. The largest anthropogenically induced hypoxic area in North American coastal waters (of 15 000 ± 5000 km²) forms every summer in the northern Gulf of Mexico where the Mississippi and Atchafalaya rivers deliver large amounts of freshwater and nutrients to the shelf. A 2001 plan for reducing this hypoxic area by nutrient management in the watershed called for a reduction of N loads. Since then evidence of P limitation during the time of hypoxia formation has arisen, and a dual nutrient-reduction strategy for this system has been endorsed.

Here we report the first systematic analysis of the effects of single and dual nutrient load reductions from a spatially explicit physical–biogeochemical model for the northern Gulf of Mexico. The model has been shown previously to skillfully represent the processes important for hypoxic formation. Our analysis of an ensemble of simulations with stepwise reductions in N, P, and N and P loads provides insight into the effects of both nutrients on primary production and hypoxia, and it allows us to estimate what nutrient reductions would be required for single and dual nutrient-reduction strategies to reach the hypoxia target. Our results show that, despite temporary P limitation, N is the ultimate limiting nutrient for primary production in this system. Nevertheless, a reduction in P load would reduce hypoxia because primary production is P limited in the region where density stratification is conducive to hypoxia development, but reductions in N load have a bigger effect. Our simulations show that, at present loads, the system is almost saturated with N, in the sense that the sensitivity of primary production and hypoxia to N load is much lower than it would be at lower N loads. We estimate that reductions of 63 ± 18 % in total N load or 48 ± 21 % in total N and P load are necessary to reach a hypoxic area of 5000 km², which is consistent with previous estimates from statistical regression models and highly simplified mechanistic models.

1 Introduction

Coastal eutrophication as a result of anthropogenic nutrient inputs is a longstanding and growing problem worldwide with negative effects that include hypoxia (Diaz and Rosenberg, 2008), degradation of habitat and harmful algal blooms (Huisman et al., 2005). The most important limiting nutrients in aquatic systems are nitrogen (N) and phosphorus (P), and both have major anthropogenic sources (Seitzinger et al., 2010). Fixed N (i.e., N in its bioavailable forms nitrate, nitrite and ammonium) enters aquatic systems mainly through leakage of industrial N-based fertilizer applied in agriculture and through deposition of N resulting from combustion of fossil fuels. P is added to aquatic systems primarily from urban and industrial wastewater as well as fertilizers. Because these nutrients have different sources, the management actions required to reduce one or the other are different. While P can be controlled by wastewater treatment, control of N requires a decrease in N-based fertilizer use. The need for reducing anthropogenic nutrient inputs to aquatic systems is well recognized, but doing so comes at a significant cost and can be met with substantial political obstacles, in particular...
with regard to N. Sound nutrient management strategies are thus needed.

There is a long-standing debate about whether controlling only N or P or both of these nutrients is most appropriate for reversing the detrimental effects of eutrophication. As summarized by Conley et al. (2009), in the 1970s it was established that P is the primary limiting nutrient in several Canadian lakes (Schindler et al., 2008; Schindler, 1974). Widespread reductions in P loads to North American and European aquatic systems led to improvements in water quality in many lakes, but not in estuarine and coastal systems (see Conley et al., 2009, and references therein). It was concluded that N input needs to be controlled for coastal waters (Howarth and Marino, 2006), and N has been the main target of nutrient load reduction strategies for many estuarine and coastal systems.

In this context the concept of ultimate versus proximate limiting nutrient is useful. Tyrrell (1999) defines the proximate limiting nutrient as the one that is locally or temporarily limiting primary production; its addition would lead to an immediate enhancement of primary production. In contrast, supply of the ultimate limiting nutrient determines the productivity of a system over long timescales. Clarity about the relevant timescale is important when using these terms. On geological timescales (millennia and longer), P is thought to be the ultimate limiting nutrient of ocean productivity, while N is thought to be limiting only in the proximate sense (Tyrrell, 1999). On timescales of years to centuries, productivity in the present ocean is clearly limited by the supply of fixed N. In pristine lakes, P is the ultimate limiting nutrient (Schindler et al., 2008). Estuarine and coastal systems that receive heavy nutrient loads can be limited by P or N or both (Conley, 1999; Sylvan et al., 2007). Which nutrient is limiting can vary significantly in time and space (Malone et al., 1996; Sylvan et al., 2007). Establishing for a given estuarine or coastal system which of the two nutrients is the ultimate limiting one (on timescales of years to decades) should inform the design of sound nutrient-reduction strategies.

The largest hypoxic area in US coastal waters is located in the northern Gulf of Mexico, where hypoxic conditions affect a region of 15 000 km² every summer (Rabalais et al., 2002; Bianchi et al., 2010). Hypoxia in this system is driven by nutrient and freshwater inputs from the Mississippi–Atchafalaya River system, which stimulate high levels of primary production, subsequent decay of organic matter, and vertical density stratification that prevents ventilation (Rabalais et al., 2002; Bianchi et al., 2010; Yu et al., 2015b). Interannual variability in the size of the hypoxic region is large and hypoxic conditions are restricted to a relatively thin layer above the bottom (Wiseman et al., 1997; Fennel et al., 2013).

N is generally limiting primary production in the Gulf of Mexico; however, observations (Sylvan et al., 2006, 2007) and models (Laurent et al., 2012; Laurent and Fennel, 2014) have shown that in spring and early summer, during the time when hypoxic conditions are established, P is limiting in the Mississippi River plume. The effect of P limitation on hypoxia in this system has been debated. Scavia and Donnelly (2007) have speculated, based on evidence from other systems (Conley, 1999; Paerl et al., 2004), that P limitation exacerbates hypoxia by spreading the detrimental effects of elevated N over a larger area. In contrast, the model analysis of Laurent and Fennel (2014) indicates that P limitation mitigates hypoxia in the northern Gulf of Mexico. The model shows that, although P limitation does shift excess N further downstream as suggested by Scavia and Donnelly (2007), the downstream N concentrations are sufficiently diluted so that less hypoxia is generated overall.

Nutrient-reduction strategies for the northern Gulf have initially focused on N, implicitly assuming that N is the ultimate limiting nutrient. In 2001, the action plan put forth by the Mississippi River/Gulf of Mexico Watershed Nutrient Task Force, an alliance of multiple state and federal agencies and tribes, called for a reduction of the size of the hypoxic zone to a running 5-year mean of 5000 km² by the year 2015 through nutrient management in the watershed (Task Force, 2001). The 2001 Action Plan stated that “the best current science indicates [...] a 30 % reduction [...] in nitrogen discharges [...] is consistent with this goal” (p. 21). In 2007, a special hypoxia advisory panel was convened by the task force and the US Environmental Protection Agency and recommended adoption of a dual-nutrient strategy with the goal of reducing N and P loads by “at least 45 %” (EPA, 2007, p. ii). In the updated action plan from 2008, a dual nutrient strategy is called for but without a specific target (Task Force, 2008, p. 29). In 2013, the hypoxia task force released a reassessment that called for a “decrease [in the] scientific uncertainty of nitrogen and phosphorus effects on hypoxia” (Task Force, 2013, p. 49). No noticeable nutrient load reductions have occurred since 2001, and the region affected by seasonal hypoxia has not decreased.

Here we use a biogeochemical model for the hypoxic region in the northern Gulf of Mexico (Fennel et al., 2011; Laurent et al., 2012; Laurent and Fennel, 2014; Laurent et al., 2017) to analyze how reduction in N and P loads affect the system. The motivation for this study is twofold. On the one hand, we aim to determine whether N or P is the ultimate limiting nutrient in this system and to elucidate how their interplay affects hypoxia development. On the other hand, we address the more practical question of how far N or P loads would have to be reduced to reach the desired reductions in hypoxia. Thus far, the primary modeling tool for defining nutrient-reduction targets has been statistical modeling that relates spring nutrient loads to summer hypoxic extent (Scavia et al., 2003; Greene et al., 2009; Forrest et al., 2011; Turner et al., 2012; Obenour et al., 2015). These models are not spatially explicit, and they ignore or highly simplify the mechanisms underlying hypoxia generation. Variations in spring nutrient load, although significantly correlated with summer hypoxic area, explain only 24 % of vari-
ability in hypoxic area in the study of Forrest et al. (2011).
When other factors like directional wind strength and fresh-
water discharge are incorporated as independent variables the
correlation improves markedly (Forrest et al., 2011). This il-
lustrates the importance of variations in atmospheric forcing
and circulation patterns in determining hypoxic conditions
on the shelf and suggests that a spatially explicit, mechanis-
tic approach is valuable. While our study is specific to the
northern Gulf of Mexico, the findings should also be relevant
to other coastal shelf systems that receive high nutrient loads,
e.g., the North Sea and the East China Sea.

2 Methods

Our biogeochemical model (Fig. 1) uses a high-resolution
implementation of the Regional Ocean Modeling System
(ROMS; Haidvogel et al., 2008) for the northern Gulf of
Mexico coupled with the relatively simple N cycle model of
Fennel et al. (2006). The original N cycle model has been ex-
panded to include phosphate as an additional nutrient (Lau-
rent et al., 2012), dissolved oxygen (Fennel et al., 2013) and
river-derived dissolved organic matter (Yu et al., 2015b). An
up-to-date description of the model equations is provided in
the supplemental information of Laurent et al. (2017). The
model is configured for the shelf region of the northern Gulf
of Mexico that frequently experiences hypoxia (Fig. 1). It has
been extensively validated by comparing standing stocks and
rates against available measurements and has been shown to
represent the biogeochemical dynamics of the system well
(see Fennel et al., 2011, 2013; Laurent et al., 2012; Laurent
and Fennel, 2014; Yu et al., 2015b).

The model is forced with 3-hourly winds from the
NCEP North American Regional Reanalysis (Mesinger et al.,
2006), climatological surface heat and freshwater fluxes from
da Silva et al. (1994), and daily freshwater discharge from the
Mississippi and Atchafalaya rivers recorded by the US Army
Corps of Engineers at Tarbert Landing and Simmesport, re-
spectively. Inputs of nutrients and particulate and dissolved
organic matter are based on monthly flux estimates from the
US Geological Survey using their composite method (Aulen-
bach et al., 2007). Model simulations start on 1 January 2000
and end on 31 December 2016.

In our previous studies, we have used several, qualitatively
different parameterizations for the interaction between the
sediment and overlying water column: an instant remineral-
ization (IR) parameterization, which assumes that all or-
ganic matter is remineralized immediately upon reaching the
sediment (Fennel et al., 2013); an empirical parameteriza-
tion where sediment oxygen consumption and nutrient efflux
from the sediment depend on bottom-water temperature and
oxygen (e.g., Yu et al., 2015a); and sediment flux parameteri-
zations based on a metamodel analysis of a diagenetic model
(Laurent et al., 2016). The simple empirical parameterization,
in particular, has proven useful as a computationally
efficient and accurate bottom boundary layer (Fennel et al.,
2013, 2016). However, neither the metamodel nor the empir-
cal parameterization are appropriate for this study because
they do not explicitly consider the depositional flux of or-
ganic matter to the sediment.

Since the objective here is to quantify the response of hy-
poxia to nutrient reductions, the IR parameterization has to
be used. One known issue with IR is that the simulated hy-
hypoxic area is biased low compared to the empirical parameterization and observations. In order to address this problem, we debiased the simulated hypoxic area by normalization against the observation-based estimates of Obenour et al. (2013). We calculated the average ratio between the observed and simulated hypoxic area during the hypoxic monitoring cruises (Rabalais et al., 2002) for the 12 years for which the simulation period and data set overlap and applied this ratio to all simulated hypoxic area estimates.

We conducted one 17-year simulation using the current nutrient loads as described above. Then we repeated the same simulation with reduced loads of total nitrogen (TN), dissolved inorganic phosphorus (DIP) and reduced loads of both (TN and P). We chose to reduce TN because we assume the task force goals of reducing N load are referring to the sum on inorganic and organic N. It should be noted that a reduction in the organic matter load implies not only a reduction in N but also a slight reduction in organic P load. Conversely a reduction of organic P would imply a much larger reduction in N (by a factor of 16 if Redfield stoichiometry is assumed for the composition of organic matter). Hence, we reduced only the inorganic P fraction in the DIP-reduction experiments. In all three cases (TN, DIP, and TN and DIP reductions) the loads were decreased by 20, 40, 60 and 80 %. This resulted in a total of 13 simulations. The freshwater discharge was not changed.

3 Results

3.1 Seasonal evolution of nutrients, primary production and hypoxia area

To illustrate the effects of nutrient reductions, we first compare time series of shelf-averaged surface nutrients, primary production and bottom-water hypoxia in 2009. These are shown in Fig. 2 for four simulations: the one with current nutrient loads, the one with a 60 % reduction in TN load, the one with a 60 % reduction in DIP load and the one with a 60 % reduction in both.

In the simulation with current nutrient loads (dark orange lines in Fig. 2), the shelf-averaged surface nitrate concentration is high in winter with \( \sim 10 \text{ mmol N m}^{-3} \). Nitrate increases further in spring and early summer to 18 mmol N m\(^{-3}\) due to nutrient input in spring, then decreases rapidly in mid-summer to about 5 mmol N m\(^{-3}\) as a result of algal uptake, and continues to decrease more slowly in late summer and early fall until reaching its minimum in October and beginning to increase again in late fall.

Phosphate behaves differently, reaching its maximum concentration between 0.6 and 0.8 mmol P m\(^{-3}\) in winter and its minimum concentration of 0.3 mmol P m\(^{-3}\) in summer. The amplitude of seasonal changes (i.e., the ratio between maximum and minimum concentration) is smaller for phosphate than nitrate. The difference in the seasonal cycles of nitrate and phosphate is a result of P limitation, which, according to observations (Sylvan et al., 2006) and model simulations (Laurent et al., 2012), occurs in plume waters during spring and early summer after the annual maximum in riverine nutrient input. P limitation of primary production results in an accumulation of inorganic nitrogen in early summer. In late summer and early fall, as river-derived high-nitrate waters...
mix with marine waters that hold an excess of phosphate relative to nitrate, P limitation is relieved and most of the accumulated nitrate is eventually taken up until minimum concentrations are reached in fall.

In the simulation with 60% reduction in DIP load (light orange lines in Fig. 2), average phosphate is, as expected, much smaller than for current loads. This makes P limitation in the plume in spring and early summer more severe and further amplifies the accumulation of surface nitrate in summer (i.e., average nitrate concentrations in summer are larger in the simulation with P reductions than in the simulation without nutrient reductions).

A 60% reduction in TN load (light purple lines in Fig. 2) results in the smallest surface nitrate concentrations of all four simulations. The June nitrate maximum of 18 mmol N m\(^{-3}\) in the current-load simulation is reduced to 2.5 mmol N m\(^{-3}\), a decrease of 86% that is much larger than the 60% reduction in river TN load; i.e., nitrate accumulation in summer is less pronounced. In the simulation where both TN and P loads are reduced by 60% (dark purple lines), the summer nitrate concentrations are slightly larger than in the case where only TN load is reduced, indicating again that more severe P limitation in the summer results in nitrate accumulation.

Figure 3. Bottom-water oxygen on 28 July 2009 (the day of maximum simulated hypoxic extent in 2009) for the simulations with current nutrient load, 60% reduction in DIP load, 60% reduction in TN load, and 60% in both (same simulations as shown in Fig. 2), and primary production averaged over May, June and July in 2009. The bold black lines in the bottom panels outline the hypoxic zones (< 2 mg O\(_2\) L\(^{-1}\)) shown in the top panels, and the thin lines outline the area with < 3 mg O\(_2\) L\(^{-1}\).
Shelf-averaged primary production is, as expected, largest in the current-load simulation. Compared to this, primary production slightly decreases in early summer when the DIP load is reduced, but decreases significantly from March to November for a reduced TN load. In the simulation with TN and P load reduction, primary production is similar to the result for the TN load reduction, except for a brief period in early summer when primary production is even lower. The effect of DIP load reductions on primary production is thus minor as illustrated by the two pairs of simulations (current load versus DIP load reduction, and TN load reduction versus TN and P load reduction), while the TN load reduction has a big effect. Annually integrated values of primary production for 2009, i.e., the year shown in Fig. 2, and averaged over the years 2000 to 2016 are listed in Table 1.

Hypoxic conditions occur in all four simulations shown in Fig. 2 from early June to the end of September, but the simulated spatial extent of hypoxia is different in all of them. Hypoxia is most expansive in the current-load simulation, decreases significantly in the simulation with P load reduction, is reduced further in the simulation with N load reduction, and is smallest in the simulation with N and P load reduction (Fig. 3). The corresponding spatial distributions of average primary production are shown in Fig. 3. The values of annually integrated hypoxic area, primary production, and H are given in Table 1.

Given the relatively minor effect of DIP load reductions on shelf-averaged primary production, the large sensitivity of the simulated hypoxic area to DIP load is perhaps surprising. However, as discussed in previous publications, hypoxic conditions are spatially and temporally constrained by the “stratification envelope,” i.e., the existence of a stratification regime that is conducive to hypoxia by preventing ventilation of bottom waters (Hetland and DiMarco, 2008). Hypoxic extent is thus sensitive to the spatiotemporal alignment between peak primary production and the occurrence of the stratification envelope (Laurent and Fennel, 2014). Temporal and spatial shifts, and changes in magnitude of peak primary production in summer, which result from variations in P load, can thus have a notable effect on hypoxia without altering shelf-averaged primary production significantly.

Next, we systematically compare annually integrated values of primary production and hypoxic area for the different nutrient load reductions and determine their sensitivity to nutrient load decreases.

### 3.2 Sensitivity of primary production and hypoxia to nutrient load reductions

We define the dimensionless sensitivity $S$ of a system property (e.g., shelf-averaged PP or H) to nutrient load reduction as the ratio between the change in this property (in %) to the imposed change in nutrient load (in %). In other words, a sensitivity of PP equal to 1 implies that for a 10 % decrease in nutrient load, a 10 % decrease in PP can be expected. If the sensitivity is less than 1, a 10 % decrease in nutrient load would bring about a decrease in PP of $S \times 10 \%$, i.e., smaller than 10 %. When $S$ is larger than 1, the change in PP would be larger than 10 %.

| Table 1. May total nitrogen (TN) load, simulated annual primary production (PP) and annually integrated hypoxic area H for 2009 (normal font) and averaged over the years 2000 to 2016 (in bold) for current loads and selected nutrient load reduction simulations. |
|---|---|---|---|
| May TN load (10$^9$ mol N) | 12.8 (year 2009) | **11.7** (avg 2000–2016) |  |
| Annual PP (mol O$_2$ m$^{-2}$ yr$^{-1}$) | 25.4 | **26.7** |  |
| Hypoxic area H (10$^{13}$ km$^2$ yr) | 4.61 | **3.83** |  |

<table>
<thead>
<tr>
<th>DIP load reduction</th>
<th>20 %</th>
<th>40 %</th>
<th>60 %</th>
<th>80 %</th>
</tr>
</thead>
<tbody>
<tr>
<td>May TN load (10$^9$ mol N)</td>
<td>12.8</td>
<td>same</td>
<td>same</td>
<td>same</td>
</tr>
<tr>
<td>Annual PP (mol O$_2$ m$^{-2}$ yr$^{-1}$)</td>
<td>25.1</td>
<td>24.7</td>
<td>24.2</td>
<td>23.4</td>
</tr>
<tr>
<td>Hypoxic area H (10$^{13}$ km$^2$ yr)</td>
<td>3.89</td>
<td>3.01</td>
<td>1.99</td>
<td>1.02</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>TN load reduction</th>
<th>20 %</th>
<th>40 %</th>
<th>60 %</th>
<th>80 %</th>
</tr>
</thead>
<tbody>
<tr>
<td>May TN load (10$^9$ mol N)</td>
<td>10.3</td>
<td>7.69</td>
<td>5.12</td>
<td>2.56</td>
</tr>
<tr>
<td>Annual PP (mol O$_2$ m$^{-2}$ yr$^{-1}$)</td>
<td>23.5</td>
<td>21.2</td>
<td>17.8</td>
<td>12.9</td>
</tr>
<tr>
<td>Hypoxic area H (10$^{13}$ km$^2$ yr)</td>
<td>3.56</td>
<td>2.46</td>
<td>1.00</td>
<td>0.10</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>TN and P load reduction</th>
<th>20 %</th>
<th>40 %</th>
<th>60 %</th>
<th>80 %</th>
</tr>
</thead>
<tbody>
<tr>
<td>May TN load (10$^9$ mol N)</td>
<td>10.3</td>
<td>7.69</td>
<td>5.12</td>
<td>2.56</td>
</tr>
<tr>
<td>Annual PP (mol O$_2$ m$^{-2}$ yr$^{-1}$)</td>
<td>23.2</td>
<td>20.7</td>
<td>17.4</td>
<td>12.7</td>
</tr>
<tr>
<td>Hypoxic area H (10$^{13}$ km$^2$ yr)</td>
<td>2.92</td>
<td>1.36</td>
<td>0.29</td>
<td>0.004</td>
</tr>
</tbody>
</table>
At current DIP loads, the sensitivity of H to TN load reductions is larger than the sensitivity of PP, and also increases for smaller nutrient loads from 1.1 at loads around $10 \times 10^9$ mol N to 1.6 for loads less than $5 \times 10^9$ mol N. At current loads, the sensitivity to DIP load reduction is 0.78, smaller than the sensitivity to TN load reduction. In other words, a 10 % decrease in TN load would shrink the hypoxic area more than a 10 % decrease in DIP load. At lower DIP loads, the sensitivity of H to N load reduction decreases.

The results in Fig. 4 illustrate that reductions in P load would have a much smaller effect on system-wide primary production than reductions in N load. In other words, P might be limiting temporarily, but that has little bearing on the overall system productivity. N is the ultimate limiting nutrient in this system. The system is currently on a trajectory toward N saturation, i.e., larger nutrient loads would not increase PP significantly, because the system is almost saturated in N. These results also imply that there would be little or no P limitation without the excessive N loads. As shown in Fig. 4, at high N loads reductions in P load have a small effect on overall primary production, but this effect is much reduced for decreasing N loads and practically disappears for the 80 % N load reduction. Another important implication is that initial nutrient reductions from present loads will have a more modest effect than similar reductions would accomplish at lower N loads. However, these results also indicate that nutrient load reductions would likely bring about larger decreases in hypoxic area than in PP.

For the year 2009, which we have considered thus far, the response to nutrient load reductions is well behaved and suggests that predictive relationships can be derived. It is important to recognize that interannual variability in the phenology of freshwater and nutrient discharges and in shelf circulation results in very different hypoxia expressions from year to year (e.g., Feng et al., 2014). Next we account for interannual variability and derive predictive relationships that consider this source of uncertainty.

3.3 Defining nutrient-reduction targets

In Fig. 5 we show the simulated hypoxic area in summer in comparison with the corresponding observed estimates of Obenour et al. (2013) for the 12 years that both records overlap. This comparison illustrates that the model roughly agrees with the observations and has a similar response to variations in May N load.

May N load varies considerably from year to year (Fig. 5), but, even when comparing years with similar load, hypoxic area estimates are highly variable in the observations and the model. A large degree of interannual variability, even when nutrient loads are similar, is not surprising given the oceanographic characteristics of the system. The region that is prone to hypoxia is an open shelf system influenced by a highly dynamic river plume. The evolution of the plume distribution, vertical stratification and hypoxia is strongly affected by shelf circulation, which is determined by variable surface forcing (e.g., Feng et al., 2014), the passage of atmospheric disturbances with high wind, and meso- and submesoscale dynamics (e.g., Marta-Almeida et al., 2013; Mattern et al., 2013).

We account for the uncertainty resulting from this interannual variability by using all 17 years of our simulations. Summer hypoxic areas for all years and all simulations with TN load reductions (but without reductions in DIP load) are shown in Fig. 6a. The orange squares with error bars are means $\pm$ 1 SD of the binned data. At high nutrient loads of $> 10 \times 10^9$ mol N, SDs are large and hypoxic area is relatively insensitive to N load. In other words, the system is almost saturated in N. Below loads of $10 \times 10^9$ mol N, hypoxic area decreases with N load reductions.

Using piecewise linear regression to estimate at which TN load the hypoxic area would be reduced to 5000 km$^2$ yields an estimate of $4.3 \pm 2.1 \times 10^9$ mol N (a reduction of 63.2 $\pm$ 17.9 % of the current TN load). The same analysis is

---

**Figure 4.** Shelf-averaged annual primary production (PP) and annually integrated hypoxic area H for 2009 plotted over May N loads for 20, 40, 60 and 80 % reductions in TN and DIP loads. Gray lines and numbers indicate sensitivity to load reductions (3) as defined in the text.
repeated for the simulations with proportional reductions in TN and P (Fig. 6b). In this case, the targeted hypoxic area of 5000 km$^2$ would be reached at a load of 6.1 ± 2.5 × 10$^9$ mol N (a reduction of 48.4 ± 21.1 % of the current TN and P load).

Our model does not account for the possibility of a “legacy effect” as proposed by Turner et al. (2006, 2008). Turner and co-authors suggested that organic matter is accumulating in the sediments resulting in an increase in sediment oxygen consumption from year to year even as nutrient loads and system-wide productivity are stable. Our model does not include organic matter storage in sediments and, thus, in its present form, cannot address the question of legacy.

### 3.4 Nutrient targets in comparison to previous studies

We now compare our estimates of the reductions in nutrient load that are necessary to reach a summer hypoxic area of 5000 km$^2$ with previous estimates from the literature (see Table 2). The first estimate, based on box modeling work described in Brezonik et al. (1999) and published by the Hypoxia Task Force (2001), was a 30 % reduction. Since then a number of further estimates have been published using a range of modeling approaches.

Scavia et al. (2003, 2013) developed a one-dimensional model that simulates oxygen downstream of organic matter sources and accounts for oxygen consumption due to organic matter decomposition and resupply by ventilation. The model of Obenour et al. (2015) is a mass-balance model that simulates nutrient-stimulated primary production, organic matter sedimentation, decomposition of organic matter in water column and sediments, and ventilation. Both models are based on mechanistic assumptions but highly simplify the physical and biogeochemical processes affecting oxygen. Inspired by the linear regression model of Turner et al. (2006), the models of Greene et al. (2009); Forrest et al. (2011) and Turner et al. (2012) are purely empirical, with the first two based on multivariate linear regressions and different combinations of predictive variables and the latter based on a curvilinear fit between nutrient load and hypoxic area. In addition to these simple mechanistic models and purely empirical regressions, coupled physical–biogeochemical mod-

---

**Figure 5.** (a) Simulated midsummer hypoxic area for current nutrient loads plotted over May N load in comparison to observed values by Obenour et al. (2013) for the years 2000 to 2011. (b) Observed over simulated midsummer hypoxic area for the same years.

**Figure 6.** Simulated midsummer hypoxic area plotted over May N load for nutrient load reduction scenarios. (a) Simulated hypoxic area for current loads and all simulations with reduced N loads (black dots), binned means ± 1 SD (orange squares with error bars) and value at which target of 5000 km$^2$ hypoxic area would be reached ± 1 SD (dashed line and gray shade). (b) Same as middle panel but for proportional reductions of TN and P load.
Table 2. Previously estimated nutrient load reductions necessary to reach the hypoxic area target of 5000 km$^2$ and estimates from this study. In the right column, numbers in square brackets are 95% confidence intervals. Numbers that follow ± are SDs. TN refers to total nitrogen. P refers to phosphate. NO$_x$ refers to nitrate + nitrite. BN refers to bioavailable nitrogen as defined in the referenced study.

<table>
<thead>
<tr>
<th>Reference</th>
<th>Estimated load reduction</th>
</tr>
</thead>
<tbody>
<tr>
<td>Task force (2001)</td>
<td>30% N load</td>
</tr>
<tr>
<td>Scavia et al. (2003)$^a$</td>
<td>40–45% TN load</td>
</tr>
<tr>
<td>EPA (2007)</td>
<td>45% N and P load</td>
</tr>
<tr>
<td>Scavia and Donnelly (2007)$^a$</td>
<td>37–45% TN load</td>
</tr>
<tr>
<td></td>
<td>40–50% P load</td>
</tr>
<tr>
<td>Greene et al. (2009)$^b$ model 11</td>
<td>50% NO$_x$ load</td>
</tr>
<tr>
<td></td>
<td>42% NO$_x$ and P load</td>
</tr>
<tr>
<td>Forrest et al. (2011)$^b$ UEDC</td>
<td>68% NO$_x$ load</td>
</tr>
<tr>
<td>Turner et al. (2012)$^c$</td>
<td>57% TN load$^e$</td>
</tr>
<tr>
<td>Scavia et al. (2013)$^a$</td>
<td>62% TN load</td>
</tr>
<tr>
<td>Laurent and Fennel (2014)$^d$</td>
<td>&gt; 50% NO$_x$ and P load</td>
</tr>
<tr>
<td>Feist et al. (2016)$^d$</td>
<td>69% NO$_x$ and P load</td>
</tr>
<tr>
<td>Scavia et al. (2017) UM$^a,1$</td>
<td>58% [49–70%] TN load</td>
</tr>
<tr>
<td>NCSU$^a,2$</td>
<td>56% [50–62%] BN load</td>
</tr>
<tr>
<td>LSU$^c,3$</td>
<td>56% [50–64%] NO$_x$ load</td>
</tr>
<tr>
<td>VIMS$^b,4$</td>
<td>80 ± 70% NO$_x$ load</td>
</tr>
<tr>
<td>This study</td>
<td>63 ± 18% TN load</td>
</tr>
<tr>
<td></td>
<td>48 ± 21% TN and P load</td>
</tr>
</tbody>
</table>

$^a$ Highly simplified mechanistic model
$^b$ Multi-linear regression model
$^c$ Curvilinear regression model
$^d$ Physical–biogeochemical model
$^e$ Converted from 70/000 metric tons
$^1$ Same model as in Scavia et al. (2013)
$^2$ Same model as in Obenour et al. (2015)
$^3$ Same model as in Turner et al. (2012)
$^4$ Same model as in Forrest et al. (2011)

4 Conclusions

We presented the first comprehensive analysis of the effects of differential reductions in N and P load on primary production and hypoxia in the northern Gulf of Mexico using a spatially explicit physical–biogeochemical model. Our first objective was to determine which nutrient is the ultimate limiting one and how the interplay between N and P limitation affects hypoxia development. An ensemble of scenario simulations, where riverine loads of TN, DIP or both were reduced in a stepwise manner, shows that, while there is temporary P limitation, N supply determines shelf-averaged, time-integrated primary production. System-wide primary production is much more sensitive to variations in N load than P load. Temporary P limitation is the result of excessive N loading. This is fully consistent with the notion that N is the ultimate limiting nutrient in this system, and P is limiting only in a proximate sense. Furthermore, the sensitivity of primary production to TN load varies. At the high end of the range of current loads, the sensitivity is relatively low (∼0.4) but increases to almost 1 when TN load is reduced by at least 60%. This indicates that the system is approaching N saturation.

Our model further suggests that, although P load reductions have little effect on overall primary production, they would lead to a significant decrease in hypoxia. This is because hypoxic extent is sensitive to the spatiotemporal alignment between elevated primary production and the presence of vertical density stratification. Intensified P limitation in summer decreases the peak in production of organic matter, thus reducing the supply of organic matter in the shelf region where density stratification is conducive to hypoxia. However, hypoxia is more sensitive to N load reductions than reductions in P. As with primary production, the sensitivity of hypoxia to N load reduction changes for different N loads. Consequently, statistical extrapolation outside the historically observed range of conditions should be treated with caution.

The second objective of our study was to estimate the load reduction targets to reach the hypoxia reduction goal set by the Task Force (2001). Previously published, simple predictive models relate summer hypoxic area to May N load (see Table 2), but interannual variability in hypoxic area is large, even among years with similar May N load, because of year-to-year variations in ocean circulation and in the phenomenology of river inputs. By considering an ensemble of 17-year simulations we account for interannual variability when estimating the load reductions that would be required to reach the goal of 5000 km$^2$. Piecewise linear regression of simulated
summer hypoxic area against May N load suggests that reductions of 63±18 % of TN load or 48±21 % TN and P load are necessary. In other words, a dual nutrient strategy would be most effective in reducing hypoxia. These estimates are consistent with the previously published estimates.

Data availability. All model output is available at www.ioos.us/comt/projects/gom_hypoxia.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. This work was supported by the NOAA Coastal and Ocean Modeling Testbed (COMT) and NOAA CSCOR grants NA09NOS4780208 and NA06NOS4780198. NGOMEX publication no. 228.

Edited by: Gerhard Herndl
Reviewed by: three anonymous referees

References


