



# Trends in element incorporation in hyaline and porcelaneous foraminifera as a function of $p\text{CO}_2$

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**Abstract.** In this study we analyzed the impact of seawater carbonate chemistry on the incorporation of elements in both hyaline and porcelaneous larger benthic foraminifera. We observed a higher incorporation of Zn and Ba when  $p\text{CO}_2$  increases from 350 to 1200 ppm. Modeling the activity of free ions as a function of  $p\text{CO}_2$  shows that speciation of some elements (like Zn and Ba) is mainly influenced by the formation of carbonate complexes in seawater. Hence, differences in foraminiferal uptake of these might be related primarily by the speciation of these elements in seawater. We investigated differences in trends in element incorporation between hyaline (perforate) and porcelaneous (imperforate) foraminifera in order to unravel processes involved in element uptake and subsequent foraminiferal calcification. In hyaline foraminifera we observed a correlation of element incorporation of different elements between species, reflected by a general higher incorporation of elements in species with higher Mg content. Between porcelaneous species, inter-element differences are much smaller. Besides these contrasting trends in element incorporation, however, similar trends are observed in element incorporation as a function of seawater carbonate chemistry in both hyaline and porcelaneous species. This suggests similar mechanisms responsible for the transportation of ions to the site of calcification for these groups of foraminifera, although the contribution of these processes might differ across species.

## 1 Introduction

Calcareous foraminifera, cosmopolitan unicellular protists, are widely used to reconstruct past environmental conditions, since the chemical composition of the carbonate shells reflect a wide variety of environmental parameters. For instance, the Mg / Ca of foraminiferal shells is primarily determined by seawater temperature (Nürnberg et al., 1996; Allen and Sanders, 1994) and seawater Mg / Ca (Segev and Erez, 2006; Evans et al., 2015) and has been widely applied as a paleothermometer (Elderfield and Ganssen, 2000; Lear et al., 2000). The use of foraminifera as proxies for the inorganic carbon system in the past (seawater pH, alkalinity, saturation state, etc.) has more recently been added to the foraminiferal proxy toolbox. For example, the concentrations of trace elements in foraminiferal shells, including U (Keul et al., 2013; Russell et al., 2004), Zn (Marchitto et al., 2000; van Dijk et al., 2017) and B (Yu and Elderfield, 2007), correlate with seawater carbonate ion concentration ( $[\text{CO}_3^{2-}]$ ), while the boron isotopic composition of foraminiferal calcite is used as a proxy for pH (Sanyal et al., 1996). However, insight into vital effects (Erez, 2003) and inter-specific differences in trace element incorporation (Bentov and Erez, 2006; Toyofuku et al., 2011; Wit et al., 2012) is needed to increase robustness of these proxies.

On the broadest taxonomic scale, foraminifera produce tests using either one of two fundamentally different mechanisms. These calcification strategies reflect the evolutionary separation of foraminiferal groups dating back to the Cambrian diversification, from where the imperforate porcelaneous species and perforate hyaline foraminifera developed

independently (Pawlowski et al., 2003). Previously observed species-specific differences in partitioning and fractionation of elements most likely primarily reflect these differences in calcification strategy, since these offsets are largest between hyaline and porcelaneous species (for a summary, see Toyofuku et al., 2011). The calcification process of the latter group has been studied more extensively than that of the porcelaneous species (De Nooijer et al., 2014). Although many aspects of perforate calcification remain unsolved, there is consensus that chamber formation takes place extracellularly, but within a (semi-)enclosed space, generally termed the site of calcification (SOC). The first layers of calcite precipitate on an organic matrix (the POS or primary organic sheet) that serves as a template for the calcite layer that forms the chamber wall (Erez, 2003; Hemleben et al., 1977). To promote calcification, foraminifera furthermore need to remove Mg ions and/or protons (Zeebe and Sanyal, 2002) from the seawater entering the SOC. A number of larger benthic foraminifera form hyaline shells, although the amount of Mg in their shells is often more than 10 times higher than that of planktonic and small benthic hyaline species, hence covering a larger range in Mg / Ca<sub>CALCITE</sub> values. The calcification strategy of porcelaneous foraminifera is less well studied, which may be partly explained by their limited application in paleoceanography. Porcelaneous foraminifera use a different mode of calcification (Debenay et al., 1998; De Nooijer et al., 2009; Berthold, 1976; Hemleben et al., 1986) and produce shells without pores (hence the term imperforate) consisting of tablets or needles (Erez, 2003; Bentov and Erez, 2006; Debenay et al., 1998). These calcitic needles (2–3 µm) are precipitated intracellularly (Berthold, 1976), after which they are transported out of the foraminifer to form a new chamber (Angell, 1980). At the outer and inner layers of these chambers, the needles are arranged along the same orientation so that they form an optically homogenous surface, giving it a shiny (hence the term “porcelaneous”) appearance. In general, the Mg / Ca values of the shells of porcelaneous foraminifera are high.

Remarkably, despite this large biological control, incorporation of minor and trace elements still reflects environmental conditions, in both hyaline and porcelaneous foraminiferal shells. Systematic offsets between different species, interdependence of trace elements incorporated (Langer et al., 2016) and the different response of element incorporation on element speciation (Wit et al., 2013; Keul et al., 2013; van Dijk et al., 2017) potentially provide useful clues for determining which processes play an important role in the biomineralization pathways. Here we present the results from a controlled growth experiment for which we used several (intermediate- and high-Mg) hyaline and porcelaneous species combined with an inter-species comparison of trace element incorporation. We assessed the impact of bio-calcification on element incorporation as a function of  $p\text{CO}_2$  in order to explore the proposed inorganic carbonate proxies (e.g., Zn / Ca; van Dijk et al., 2017) and the impact of different calcification strate-

gies (hyaline versus porcelaneous) on element partitioning. We cultured eight benthic foraminiferal species (four hyaline and four porcelaneous) under four different  $p\text{CO}_2$  conditions, analyzing incorporation of Mg, Sr, Na, Zn and Ba.

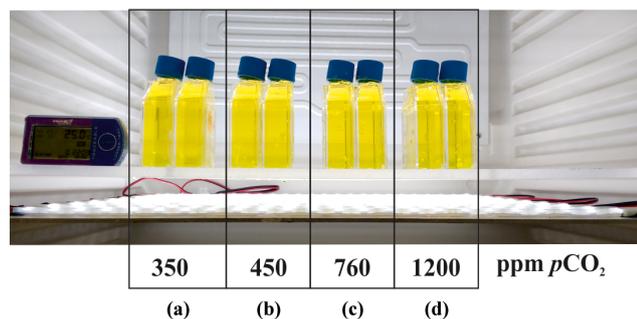
## 2 Methods

### 2.1 Foraminiferal collection

Large samples of macroalgae (*Dictyota* sp.) were collected in November 2015 at a depth of 2–3 m in Gallows Bay, St. Eustatius (17°28'31.6" N, 62°59'9.4" W). Salinity was ~34 and temperature was ~29 °C at the site of collection. The collected macroalgae were transported to the laboratory at the Caribbean Netherlands Science Institute (CNSI), where they were placed in a 5 L aquarium with aerated and unfiltered seawater. From this stock, small amounts of algae and debris were gently sieved over a 90 and 600 µm mesh to carefully dislodge foraminifera. Several species of foraminifera were picked from the resulting 90–600 µm fraction and directly from the macroalgae. Living specimens of *Sorites marginalis* Lamarck, 1816, *Amphistegina gibbosa* d'Orbigny, 1839, *Laevipeneroplis bradyi* Cushman, 1930 and *Archaias angulatus* Fichtel and Moll, 1798 and limited amounts (< 20) of *Peneroplis pertusus* Forskål, 1775, *Asterigerina carinata* d'Orbigny, 1839, *Heterostegina antillarum* d'Orbigny, 1839, and *Planorbulina acervalis* Brady, 1884, characterized by yellow cytoplasm and pseudopodial activity, were selected for the culturing experiments.

### 2.2 Culture setup

Four barrels, each filled with 100 L of seawater (5 µm filtered), were connected to a Li-Cor CO<sub>2</sub>/H<sub>2</sub>O analyzer (LI-7000) to regulate the CO<sub>2</sub> level in the barrels' headspace. The set points were maintained by addition of CO<sub>2</sub> and/or CO<sub>2</sub>-scrubbed air according to the monitored  $p\text{CO}_2$ . The set points for  $p\text{CO}_2$  were 350, 450, 760 and 1400 ppm, resulting in four batches of seawater (treatments A–D) differing only in their inorganic carbon chemistry. Salinity (34.0 ± 0.2) was monitored with a salinometer (VWR CO310). The fluorescent compound calcein (bis[*N,N*-bis(carboxymethyl)aminomethyl]-fluorescein) was added to the culture media (5 mg L<sup>-1</sup> seawater) to enable determination of newly formed chambers during the culture experiment (Bernhard et al., 2004). Short-term exposure (< 3 weeks) to calcein has no detectable impact on the physiology of benthic foraminifera (Kurtarkar et al., 2015), and the presence of calcein has no effect on the incorporation of Mg and Sr in foraminiferal calcite (Dissard et al., 2009). Culture media were stored air-free in portions of 250 mL in Nalgene bottles with Teflon-lined caps at 4 °C until further use.



**Figure 1.** Photograph of the culture setup. Four duplicate bottles with culture media (with calcein added) and adult specimens of foraminifera. Treatments with corresponding set points are A, 350 ppm; B, 450 ppm; C, 760 ppm; D, 1200 ppm  $p\text{CO}_2$ .

Foraminifera were divided over the different treatments in duplicate and placed in 70 mL Falcon<sup>®</sup> tissue bottles with gas-tight caps in a thermostat set at 25 °C (Fig. 1). The thermostat was monitored by a temperature logger (Traceable Logger Trac, Maxi Thermal), recording the temperature every minute. The average temperature over the whole experiment was  $25 \pm 0.2$  °C. To create uniform light conditions, the thermostat was equipped with two LED shelves, which resulted in high light conditions 12 h/12 h. Culture media were replaced every four days to avoid build-up of organic waste and to obtain stable seawater element concentrations and carbon chemistry. Foraminifera were fed after every water change with 0.5 mL of concentrated freeze-dried *Dunaliella salina* cells, pre-diluted with the corresponding treatment seawater. After 21 days, the experiment was terminated. Foraminifera were rinsed three times with deionized water, dried at 40 °C and stored in micropaleontology slides until further analysis at the Royal Netherlands Institute for Sea Research (NIOZ).

## 2.3 Analytical methods

### 2.3.1 Seawater carbon parameters

At the start and termination of the experiment, 125 mL samples of the seawater at each of the different experimental conditions were collected to analyze dissolved inorganic carbon (DIC) and total alkalinity (TA) on a Versatile Instrument for the Determination of Titration Alkalinity (VINDTA) at the CNSI. Using the measured DIC and TA values and the software CO2SYS v2.1, adapted to Excel by Pierrot et al. (2006), the other carbon parameters (including  $[\text{CO}_3^{2-}]$  and  $\Omega_{\text{CALCITE}}$ ) were calculated. For this we used the equilibrium constants for K1 and K2 from Lueker et al. (2000) and  $\text{KHSO}_4$  from (Dickson, 1990) (Table 1).

### 2.3.2 Seawater element concentrations

At the start and end of the experiment and during replacement of the culture media, subsamples were collected in duplicate using 50 mL LDPE Nalgene bottles and immediately frozen at  $-80$  °C. After transportation to the NIOZ, defrosted samples were acidified with triple quartz-distilled HCl to  $\text{pH} \sim 1.8$ , and the seawater composition of the samples was analyzed on an Element 2 sector field double-focusing mass spectrometer (SF-ICP-MS) run in medium-resolution mode. IAPSO (International Association for the Physical Sciences of the Ocean) standard seawater was used as a drift monitor. Analytical precision (relative standard deviation) was 3 % for Ca, 4 % for Mg, 1 % for Na, 1 % for Sr and 5 % for Ba. We obtained average values of  $5.25 \pm 0.06$  mol mol<sup>-1</sup> for Mg / Ca,  $44.6 \pm 0.6$  mol mol<sup>-1</sup> for Na / Ca,  $8.63 \pm 0.05$  mol mol<sup>-1</sup> for Sr / Ca, and  $9.04 \pm 0.47$   $\mu\text{mol mol}^{-1}$  for Ba / Ca.

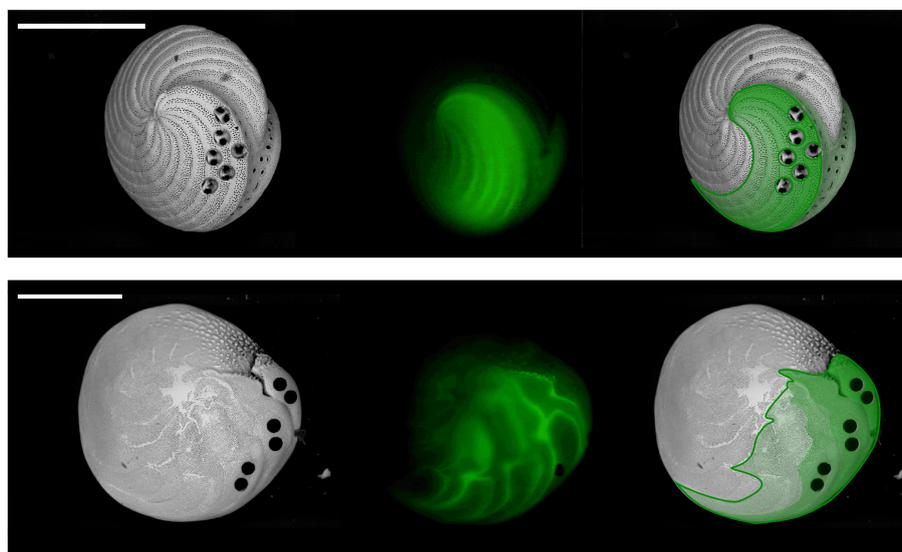
A subsample was analyzed using a commercially available pre-concentration system, SeaFAST S2. With the SeaFAST system, elements with low concentrations are pre-concentrated to values above detection limit of the SF-ICP-MS. Accordingly, we measured Cd, Pb, U, B, Ti, Mn, Fe, Co, Ni, Cu, and Zn. In short, 10 mL of sample was mixed with an ammonium acetate buffer to  $\text{pH} 6.2$  and loaded on a column containing NOBIAS chelating agent. After rinsing the column with a diluted ammonium acetate buffer the metals were eluted in 750  $\mu\text{L}$  of quartz-distilled 1.5 M  $\text{HNO}_3$  before being quantified on the SF-ICP-MS. Here we use the Zn data only, as this was analyzed in the foraminifera as well. Analytical precision (relative standard deviation) was 5 % for Zn. We obtained an average value of  $15.3 \pm 0.5$   $\mu\text{mol mol}^{-1}$  for seawater Zn / Ca for all treatments. Although these values are clearly above natural open ocean values, the concentrations are very uniform between treatments and when comparing start and end of the experiments. The contamination with Zn hence might have already occurred when filling the culture setup with the waters from the bay adjacent to the culture facility. Concentrations are well below values considered harmful for foraminifera (Nardelli et al., 2016).

### 2.3.3 Cleaning methods

After termination of the experiment, foraminiferal shells were cleaned following an adapted version of Barker et al. (2003). Per treatment duplicate, all foraminifera were transferred to 10 mL polyethylene vials. To each vial, 10 mL of 1 %  $\text{H}_2\text{O}_2$  solution (buffered with 0.5 M  $\text{NH}_4\text{OH}$ ) was added to remove organic matter. The vials were heated for 10 min in a water bath at 95 °C, and placed in an ultrasonic bath for 30 s (degas mode, 80 kHz, 50 % power), after which the oxidizing reagent was removed. These steps (organic removal procedure) were repeated five times. Foraminiferal samples were rinsed five times with ultrapure water, after which the vials were stored overnight in a laminar flow cabinet at room temperature to dry. Dried foraminifera were

**Table 1.** Carbon parameters (TA, total alkalinity,  $n = 2$ ; DIC, dissolved inorganic carbon,  $n = 2$ ) with (relative) standard deviation of the culture water per treatment of the  $p\text{CO}_2$  experiment. CO2SYS was used to calculate seawater carbonate ion concentration, calcite saturation state and pH from measured TA and DIC.

Treatment	Set point	Measured		Calculated CO2SYS		
	$p\text{CO}_2$ ppm	TA $\mu\text{mol kg}^{-1}$	DIC $\mu\text{mol kg}^{-1}$	$[\text{CO}_3^{2-}]$ $\mu\text{mol kg}^{-1}$	pH (total scale)	$\Omega_{\text{CALCITE}}$
A	350	$2302.8 \pm 8.2$	$2007.5 \pm 10.7$	220.7	8.06	5.4
B	450	$2305.2 \pm 5.8$	$2021.3 \pm 12.5$	200.0	8.01	4.9
C	760	$2304.4 \pm 0.9$	$2100.8 \pm 13.4$	153.7	7.87	3.7
D	1200	$2300.3 \pm 0.7$	$2201.4 \pm 4.1$	92.2	7.61	2.2



**Figure 2.** SEM (left) and fluorescence microscope (middle) photographs of *A. angulatus* (top series) and *A. gibbosa* (bottom series) to assess newly formed chambers for laser ablation (right). Scale bar = 500  $\mu\text{m}$ .

placed on double-sided tape on LA-ICP-MS stubs. Pictures were taken of individual foraminifera with a ZEISS Axio-plan 2 fluorescence microscope equipped with appropriate excitation and emission optics and a ZEISS AxioCam MRC 5 camera to assess the number of chambers added during the experiment based on the incorporation of calcein.

### 2.3.4 LA-ICP-MS

Element concentrations of individual fluorescent chambers were analyzed by laser ablation ICP-MS (Reichert et al., 2003; van Dijk et al., 2017). To determine foraminiferal element concentrations, the laser system (NWR193UC, New Wave Research) at the Royal NIOZ was equipped with a two-volume sample cell (New Wave Research), characterized by a wash-out time of 1.8 s (1% level) and hence allowing detection of variability in obtained element to Ca ratios within chamber walls. Single chambers were ablated in a helium environment using a circular laser spot with a diameter of 80  $\mu\text{m}$  (*S. marginalis*) or 60  $\mu\text{m}$  (other species). We ablated

all calcein-stained chambers twice, except for the first 1–2 chambers that formed during the experiment to avoid contamination of calcite of chambers formed prior to the experiments and overlapped by the labeled chambers (Fig. 2).

All foraminiferal samples were ablated with an energy density of  $1 \pm 0.1 \text{ J cm}^{-2}$  and a repetition rate of 6 Hz. The resulting aerosol was transported on a helium flow through an in-house-built smoothing device, being mixed with a nitrogen flow ( $5 \text{ mL min}^{-1}$ ) before entering the quadrupole ICP-MS (iCAP-Q, Thermo Scientific). Monitored masses included  $^7\text{Li}$ ,  $^{11}\text{B}$ ,  $^{23}\text{Na}$ ,  $^{24}\text{Mg}$ ,  $^{25}\text{Mg}$ ,  $^{27}\text{Al}$ ,  $^{43}\text{Ca}$ ,  $^{44}\text{Ca}$ ,  $^{66}\text{Zn}$ ,  $^{88}\text{Sr}$  and  $^{137}\text{Ba}$ . Contrary to  $^{67}\text{Zn}$  and  $^{68}\text{Zn}$ ,  $^{66}\text{Zn}$  is free of interferences when measuring calcium carbonate and SRM NIST (National Institute of Standard and Technology, USA, standard reference material) glass standards (Jochum et al., 2012). Potential contamination or diagenesis of the outer or inner layer of calcite was excluded by monitoring the Al signal. At the start of each series, we analyzed SRM NIST612 and NIST610 glass standard in triplicate (using an energy

**Table 2.** Total number of LA-ICP-MS measurements per species, per treatment (A–D).

Species	<i>n</i> measurements ( <i>n</i> specimens)			
	A: 350 ppm	B: 450 ppm	C: 760 ppm	D: 1200 ppm
<i>A. angulatus</i>	62 (19)	72 (21)	76 (21)	51 (14)
<i>S. marginalis</i>	48 (14)	49 (15)	57 (18)	33 (11)
<i>A. gibbosa</i>	106 (28)	126 (32)	75 (18)	59 (15)
<i>L. bradyi</i>	21 (5)	38 (13)	27 (5)	16 (4)
<i>A. carinata</i>	12 (2)	14 (1)	19 (4)	5 (1)
<i>P. pertusus</i>		12 (2)		11 (2)
<i>H. antillarum</i>		12 (1)		14 (2)
<i>P. acervalis</i>		8 (2)		
Total	187 (49)	331 (87)	254 (66)	189 (49)

density of  $5 \pm 0.1 \text{ J cm}^{-2}$ ), JCT-1 (coral carbonate) and two in-house standards, namely NFHS (NIOZ foraminifera house standard; Mezger et al., 2016) and the Iceland spar NCHS (NIOZ calcite house standard). We further analyzed JCp-1 (coral, *Porites* sp.; Okai et al., 2002) and MACS-3 (synthetic calcium carbonate) at the start of each series and monitored drift after every 10 samples. All element to calcium ratios were calculated with an adapted version of the MATLAB-based program SILLS (Signal Integration for Laboratory Laser Systems; Guillong et al., 2008). SILLS was modified by NIOZ to evaluate LA-ICP-MS measurements on foraminifera, allowing import of Thermo Qtegra software sample list, laser data reduction and laser LOG files. Major adaptations include improved automated integration and evaluation of (calibration and monitor) standards, quality control report of the monitor standards and export in element to calcium ratios ( $\text{mmol mol}^{-1}$ ). Calibration was performed against the MACS-3 carbonate standard, with  $^{43}\text{Ca}$  as an internal standard, and we used the multiple measurements of MACS-3 for a linear drift correction. Relative analytical precision (relative standard deviation (RSD) of all MACS-3 analyses) is 3 % for  $^{23}\text{Na}$ , 3 % for  $^{24}\text{Mg}$ , 3 % for  $^{25}\text{Mg}$ , 4 % for  $^{66}\text{Zn}$ , 3 % for  $^{88}\text{Sr}$  and 3 % for  $^{137}\text{Ba}$ . In total, 961 analyses were performed on 251 specimens covering eight species cultured in four experimental conditions (see Table 2 for details).

We calculated the standard deviation (SD), RSD and standard error ( $\text{SD}/\sqrt{n}$ ; SE) per treatment. The partitioning coefficient (*D*) of an element (*E*) between seawater and foraminiferal calcite is expressed as  $D_E = (E/\text{Ca}_{\text{CALCITE}}) / (E/\text{Ca}_{\text{SW}})$ . Partition coefficients and element versus calcium ratio parameters were statistically compared with different experimental parameters (such as  $p\text{CO}_2$  or  $[\text{CO}_3^{2-}]$ ) using a two-sided *t* test with 95 % confidence levels. This also allows for the calculation of 95 % confidence intervals over the average per treatment. Pairwise comparisons were made for per *E* / Ca per species and culture conditions using ANOVA. Groups that showed significant difference were assigned different letters. When com-

paring partition coefficients to other studies, *E* /  $\text{Ca}_{\text{SW}}$  data were, in some studies, not measured. For these studies we used average seawater *E* /  $\text{Ca}_{\text{SW}}$  to calculate  $D_E$  (see also Table S1 in the Supplement), allowing comparison of partitioning coefficients.

### 3 Results

#### 3.1 Element / Ca as a function of ocean acidification

In hyaline species  $\text{Mg}/\text{Ca}_{\text{CALCITE}}$  varies between 25.9 and  $141.3 \text{ mmol mol}^{-1} \text{ Mg/Ca}$ . In contrast,  $\text{Mg}/\text{Ca}_{\text{CALCITE}}$  of porcelaneous species ranges from 121.3 to  $149.3 \text{ mmol mol}^{-1}$  (Table 3). This large spread in foraminifera *E* / Ca of hyaline species is also observed for Sr ( $1.7\text{--}3.1 \text{ mmol mol}^{-1}$ ), Na ( $3.4\text{--}19.5 \text{ mmol mol}^{-1}$ ), Zn ( $9.0\text{--}97.0 \text{ } \mu\text{mol mol}^{-1}$ ) and Ba ( $2.7\text{--}20.1 \text{ } \mu\text{mol mol}^{-1}$ ), while porcelaneous species only vary over a narrow range ( $\text{Sr} = 2.0\text{--}2.2 \text{ mmol mol}^{-1}$ ;  $\text{Na} = 3.8\text{--}5.8 \text{ mmol mol}^{-1}$ ;  $\text{Zn} = 53.0\text{--}140.8 \text{ } \mu\text{mol mol}^{-1}$ ;  $\text{Ba} = 18.0\text{--}29.0 \text{ } \mu\text{mol mol}^{-1}$ ).

In both porcelaneous and hyaline species we find an increase in  $\text{Zn}/\text{Ca}_{\text{CALCITE}}$  and  $\text{Ba}/\text{Ca}_{\text{CALCITE}}$  with  $p\text{CO}_2$ , while foraminiferal Sr / Ca, Mg / Ca and Na / Ca remain similar across the experimental conditions (Fig. 3 and Table 4). Sensitivity of both foraminiferal Zn / Ca and Ba / Ca to changes in seawater  $p\text{CO}_2$  differs between the studied porcelaneous and hyaline species. When  $p\text{CO}_2$  changes from 350 to 1200 ppm, Zn / Ca of hyaline foraminifera increases by a factor of 3.7 (*A. carinata*) or 4.5 (*A. gibbosa*), while porcelaneous foraminiferal Zn / Ca increases only by 1.3 (*S. marginalis*), 1.8 (*A. angulatus*) and 2.1 (*L. bradyi*). Sensitivity of foraminiferal Ba / Ca to the same change in  $p\text{CO}_2$  also shows a similar pattern, with Ba / Ca of hyaline species increasing by a factor of 3.6 (*A. carinata*) or 3.7 (*A. gibbosa*), while porcelaneous species increase Ba / Ca only by a factor of 1.8 (*S. marginalis*), 1.6 (*A. angulatus*) or 2.1 (*L. bradyi*).

**Table 3.** Overview of element to Ca ratios in foraminiferal calcite (Avg, average; SE, standard error) and partition coefficients  $D_E$ , with  $D_E$  of ambient conditions (treatment B) in bold.

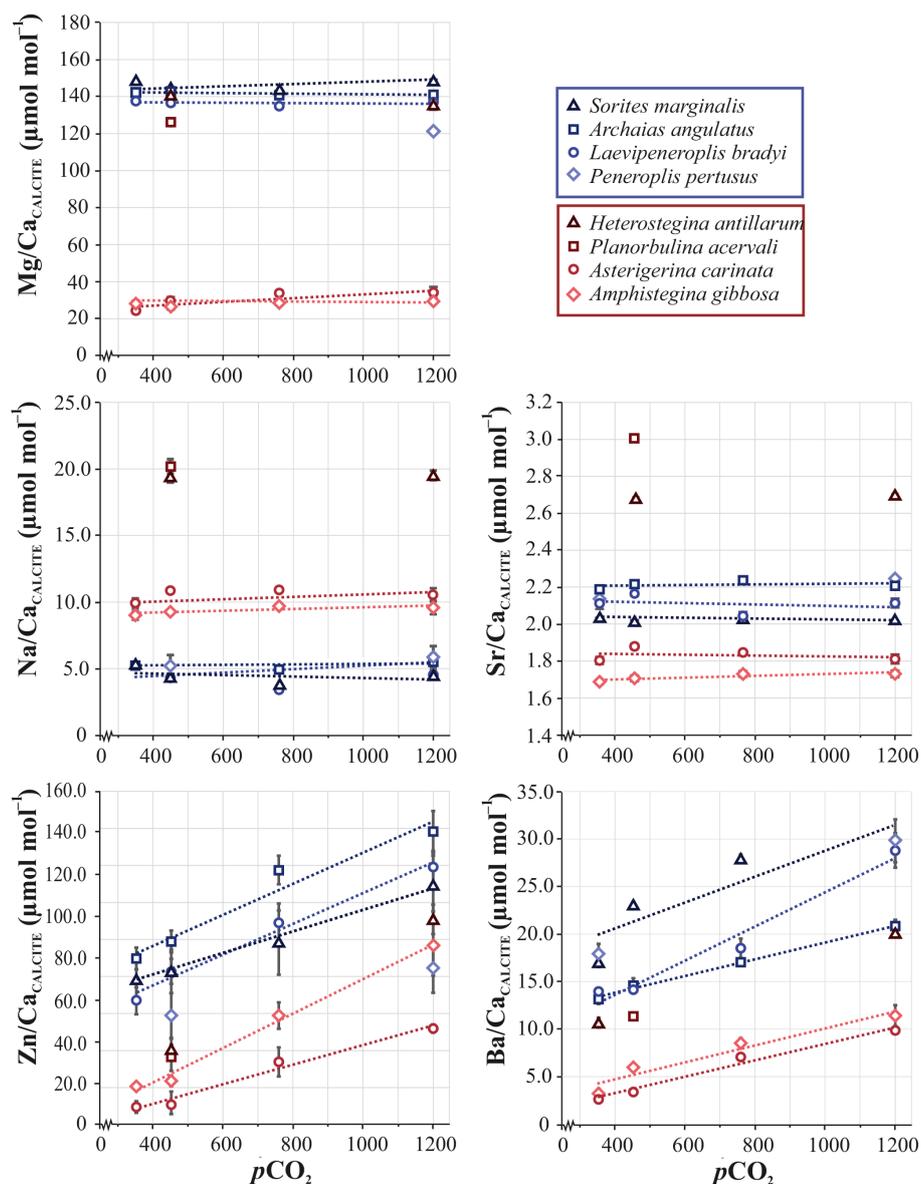
Species	$p\text{CO}_2$	Mg / Ca mmol mol <sup>-1</sup>		Na / Ca mmol mol <sup>-1</sup>		Sr / Ca mmol mol <sup>-1</sup>		Zn / Ca μmol mol <sup>-1</sup>		Ba / Ca μmol mol <sup>-1</sup>	
		Avg ± SE	$D_{\text{Mg}}$ × 10 <sup>-3</sup>	Avg ± SE	$D_{\text{Na}}$ × 10 <sup>-3</sup>	Avg ± SE	$D_{\text{Sr}}$	Avg ± SE	$D_{\text{Zn}}$	Avg ± SE	$D_{\text{Ba}}$
Porcelaneous species											
<i>A. angulatus</i>	350	139.4 ± 0.6 <sup>a</sup>	26.6	5.2 ± 0.1 <sup>a</sup>	0.12	2.2 ± 0.02 <sup>a</sup>	0.25	80.0 ± 5.1 <sup>a</sup>	5.3	13.2 ± 0.5 <sup>a</sup>	1.5
	<b>450</b>	<b>137.7 ± 0.5<sup>b</sup></b>	<b>26.3</b>	<b>4.3 ± 0.1<sup>b</sup></b>	<b>0.10</b>	<b>2.2 ± 0.01<sup>a</sup></b>	<b>0.26</b>	<b>88.1 ± 5.2<sup>b</sup></b>	<b>5.8</b>	<b>14.6 ± 0.5<sup>b</sup></b>	<b>1.6</b>
	760	137.4 ± 0.7 <sup>b</sup>	26.2	4.9 ± 0.1 <sup>c</sup>	0.11	2.2 ± 0.01 <sup>a</sup>	0.26	122.6 ± 7.0 <sup>c</sup>	8.1	17.0 ± 0.6 <sup>b</sup>	1.9
	1200	138.6 ± 1.1 <sup>a</sup>	26.4	5.4 ± 0.2 <sup>a</sup>	0.12	2.2 ± 0.02 <sup>a</sup>	0.26	140.8 ± 9.9 <sup>d</sup>	9.3	20.9 ± 0.2 <sup>c</sup>	2.3
<i>S. marginalis</i>	350	147.7 ± 0.6 <sup>a</sup>	28.2	4.8 ± 0.1 <sup>a</sup>	0.11	2.0 ± 0.01 <sup>a</sup>	0.24	70.0 ± 10.1 <sup>a</sup>	4.6	17.0 ± 0.5 <sup>a</sup>	1.9
	<b>450</b>	<b>144.2 ± 0.8<sup>b</sup></b>	<b>27.5</b>	<b>4.1 ± 0.1<sup>b</sup></b>	<b>0.09</b>	<b>2.0 ± 0.01<sup>a</sup></b>	<b>0.23</b>	<b>74.0 ± 10.6<sup>b</sup></b>	<b>4.9</b>	<b>23.1 ± 0.5<sup>b</sup></b>	<b>2.6</b>
	760	143.0 ± 0.6 <sup>a</sup>	27.3	3.8 ± 0.1 <sup>a</sup>	0.09	2.0 ± 0.01 <sup>a</sup>	0.23	87.7 ± 15.5 <sup>c</sup>	5.8	27.9 ± 0.6 <sup>c</sup>	3.1
	1200	148.3 ± 0.5 <sup>b</sup>	28.3	4.5 ± 0.2 <sup>c</sup>	0.10	2.0 ± 0.01 <sup>a</sup>	0.23	115.6 ± 15.3 <sup>d</sup>	7.6	30.1 ± 0.2 <sup>d</sup>	3.3
<i>L. bradyi</i>	350	137.8 ± 1.3 <sup>a</sup>	26.3	5.2 ± 0.2 <sup>c</sup>	0.12	2.1 ± 0.03 <sup>a</sup>	0.24	60.0 ± 6.5 <sup>a</sup>	4.0	14.0 ± 0.5 <sup>a</sup>	1.5
	<b>450</b>	<b>136.2 ± 0.7<sup>a</sup></b>	<b>26.0</b>	<b>4.3 ± 0.1<sup>b</sup></b>	<b>0.10</b>	<b>2.2 ± 0.01<sup>b</sup></b>	<b>0.25</b>	<b>73.8 ± 6.0<sup>b</sup></b>	<b>4.9</b>	<b>14.2 ± 0.5<sup>a</sup></b>	<b>1.6</b>
	760	134.4 ± 1.2 <sup>b</sup>	25.6	3.4 ± 0.1 <sup>a</sup>	0.08	2.0 ± 0.02 <sup>c</sup>	0.24	97.5 ± 9.4 <sup>c</sup>	6.4	18.5 ± 0.6 <sup>b</sup>	2.1
	1200	136.9 ± 1.1 <sup>a</sup>	26.1	6.2 ± 0.2 <sup>d</sup>	0.14	2.1 ± 0.02 <sup>a</sup>	0.24	124.2 ± 7.8 <sup>d</sup>	8.2	28.8 ± 0.2 <sup>c</sup>	3.2
<i>P. pertusus</i>	350										
	<b>450</b>	<b>126.1 ± 1.8<sup>a</sup></b>	<b>24.0</b>	<b>5.2 ± 0.3<sup>a</sup></b>	<b>0.12</b>	<b>2.1 ± 0.07<sup>a</sup></b>	<b>0.25</b>	<b>53.0 ± 10.8<sup>a</sup></b>	<b>3.5</b>	<b>18.0 ± 0.5<sup>a</sup></b>	<b>2.0</b>
	760										
	1200	121.3 ± 1.0 <sup>a</sup>	23.1	5.8 ± 0.2 <sup>a</sup>	0.13	2.2 ± 0.02 <sup>a</sup>	0.26	75.5 ± 11.9 <sup>b</sup>	5.0	29.8 ± 0.2 <sup>b</sup>	3.3
Hyaline species											
<i>H. anillarium</i>	350										
	<b>450</b>	<b>141.3 ± 0.3<sup>a</sup></b>	<b>26.9</b>	<b>19.4 ± 0.5<sup>a</sup></b>	<b>0.44</b>	<b>2.7 ± 0.02<sup>a</sup></b>	<b>0.31</b>	<b>36.0 ± 14.7<sup>a</sup></b>	<b>2.4</b>	<b>10.7 ± 0.5<sup>a</sup></b>	<b>1.2</b>
	760										
	1200	136.9 ± 1.6 <sup>a</sup>	26.1	19.5 ± 0.4 <sup>a</sup>	0.44	2.7 ± 0.02 <sup>a</sup>	0.31	97.0 ± 18.3 <sup>b</sup>	6.4	20.1 ± 0.2 <sup>b</sup>	2.2
<i>P. acervalis</i>	350										
	<b>450</b>	<b>139.1 ± 1.2</b>	<b>26.5</b>	<b>19.5 ± 0.7</b>	<b>0.46</b>	<b>3.1 ± 0.02</b>	<b>0.36</b>	<b>31.6 ± 6.6</b>	<b>2.1</b>	<b>11.3 ± 0.5</b>	<b>1.3</b>
	760										
	1200										
<i>A. carinata</i>	350	23.6 ± 1.5 <sup>a</sup>	4.5	9.9 ± 0.4 <sup>a</sup>	0.22	1.8 ± 0.02 <sup>a</sup>	0.21	9.0 ± 2.6 <sup>a</sup>	0.6	3.2 ± 0.5 <sup>a</sup>	0.4
	<b>450</b>	<b>28.5 ± 2.4<sup>b</sup></b>	<b>5.4</b>	<b>10.8 ± 0.1<sup>a</sup></b>	<b>0.24</b>	<b>1.9 ± 0.01<sup>a</sup></b>	<b>0.22</b>	<b>10.9 ± 5.5<sup>a</sup></b>	<b>0.7</b>	<b>6.0 ± 0.5<sup>b</sup></b>	<b>0.7</b>
	760	33.1 ± 1.2 <sup>b</sup>	6.3	10.9 ± 0.2 <sup>a</sup>	0.24	1.8 ± 0.01 <sup>a</sup>	0.21	30.7 ± 7.0 <sup>b</sup>	2.0	8.5 ± 0.6 <sup>c</sup>	0.9
	1200	33.5 ± 3.1 <sup>b</sup>	6.4	10.6 ± 0.5 <sup>a</sup>	0.24	1.8 ± 0.03 <sup>a</sup>	0.21	46.4 ± 2.1 <sup>b</sup>	3.1	11.4 ± 0.2 <sup>d</sup>	1.3
<i>A. gibbosa</i>	350	27.8 ± 0.5 <sup>a</sup>	5.3	9.0 ± 0.1 <sup>a</sup>	0.20	1.7 ± 0.01 <sup>a</sup>	0.20	19.0 ± 1.8 <sup>a</sup>	1.3	2.7 ± 0.5 <sup>a</sup>	0.3
	<b>450</b>	<b>25.9 ± 0.6<sup>b</sup></b>	<b>4.9</b>	<b>9.2 ± 0.1<sup>a</sup></b>	<b>0.21</b>	<b>1.7 ± 0.02<sup>a</sup></b>	<b>0.20</b>	<b>21.5 ± 2.5<sup>b</sup></b>	<b>1.4</b>	<b>3.4 ± 0.5<sup>a</sup></b>	<b>0.4</b>
	760	28.2 ± 0.7 <sup>a</sup>	5.4	9.7 ± 0.1 <sup>b</sup>	0.22	1.7 ± 0.02 <sup>a</sup>	0.20	52.8 ± 6.1 <sup>c</sup>	3.5	7.1 ± 0.6 <sup>b</sup>	0.8
	1200	28.7 ± 0.6 <sup>a</sup>	5.5	9.6 ± 0.1 <sup>b</sup>	0.21	1.7 ± 0.02 <sup>a</sup>	0.20	85.8 ± 11.3 <sup>d</sup>	5.7	9.9 ± 0.2 <sup>c</sup>	1.1

Superscript letters (a–d) indicate (per species per  $E / \text{Ca}$ ) groups that are statistically different (one-way ANOVA).

### 3.2 Inter-species differences in element incorporation

When comparing Mg incorporation to that of the other elements studied here (Ba, Zn, Sr and Na) between hyaline species (treatment B; Table 3), we observe a positive relation between  $D_{\text{Mg}}$  with  $D_{\text{Sr}}$  ( $p < 0.0025$ ),  $D_{\text{Na}}$  ( $p < 0.0005$ ),  $D_{\text{Ba}}$  ( $p < 0.05$ ) and  $D_{\text{Zn}}$  ( $p < 0.005$ ). In general, hyaline species are enriched similarly in all elements (Fig. 4). Com-

pared to porcelaneous species, the hyaline shell-building species which incorporate most Mg ( $> 100 \text{ mmol mol}^{-1}$  Mg / Ca) incorporate more Na and Sr, while incorporating less Zn and Ba. Element incorporation across porcelaneous species is less variable than observed for hyaline species. Including data from the literature (both culture and field calibrations; see Table S1 in the Supplement), prefer-



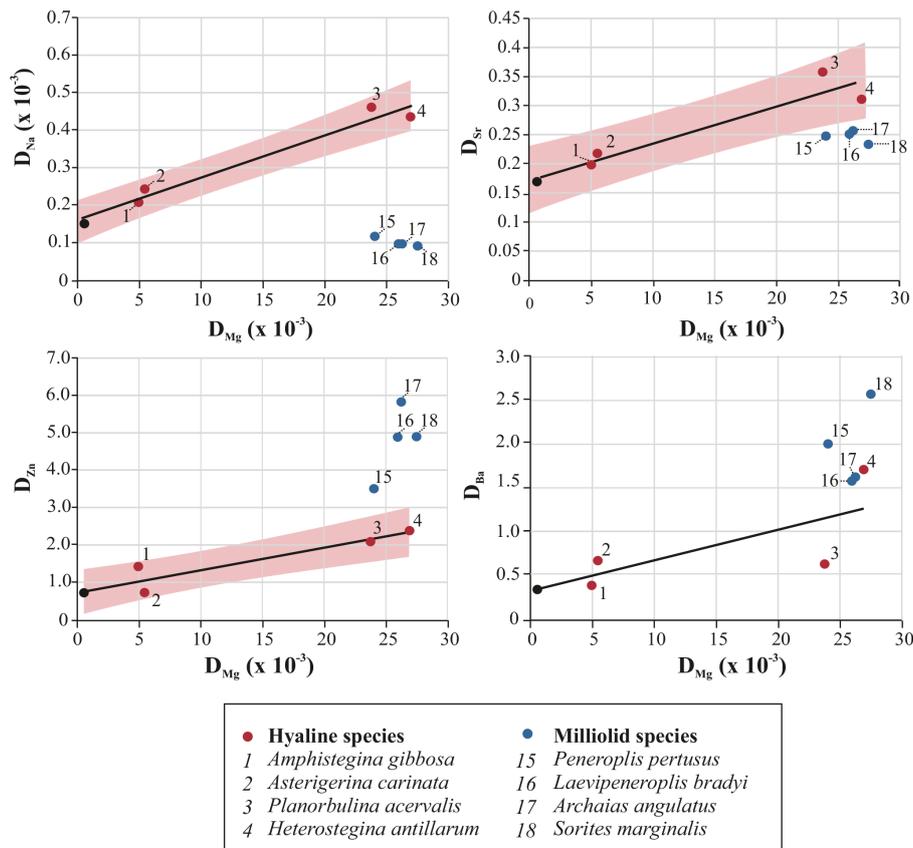
**Figure 3.** Element to Ca ratios ( $\pm$ SE) of different species of foraminifera over a range of  $p\text{CO}_2$  values. In some cases, the error bar is smaller than the symbol. Porcelaneous species in blue (triangles, *S. marginalis*; squares, *A. angulatus*; circles, *L. bradyi*; squares, *P. pertusus*) and hyaline species in red (triangles, *H. antillarum*; squares, *P. acervali*; circles, *A. carinata*; diamonds, *A. gibbosa*).

ably those in which both Mg/Ca and at least one other element (Na, Sr, Ba or Zn) are measured, shows that the relation based on the Caribbean species studied here is also more generally applicable when including more species ( $D_{\text{Sr}} = p < 0.0025$ ;  $D_{\text{Na}} = p < 0.0005$ );  $D_{\text{Ba}} = p < 0.005$ ;  $D_{\text{Zn}} = p < 0.01$ ), even though these compiled data (labeled “All studies” in Table S2) cover a somewhat wider range in environmental and experiment conditions.

## 4 Discussion

### 4.1 Effect of ocean acidification on Element / Ca

Foraminiferal Mg/Ca, Na/Ca and Sr/Ca do not systematically change with  $p\text{CO}_2$  for either porcelaneous or hyaline species. The impact of pH (and/or  $[\text{CO}_3^{2-}]$ ) on Mg/Ca<sub>CALCITE</sub> and Sr/Ca<sub>CALCITE</sub> in foraminifera has been the subject of discussion (e.g., Dissard et al., 2010; Elderfield et al., 1996). In deep-sea benthic species, incorporation of certain elements is governed by carbonate system as observed for Zn by Marchitto et al. (2005) and Cd



**Figure 4.** Partition coefficient of Na, Sr, Zn and Ba versus  $D_{Mg}$  of hyaline (red symbols) and porcelaneous (blue symbols) species. Black lines represent linear trend lines. The 95 % confidence intervals of significance trends ( $p < 0.025$ ) are indicated in red. Black dots represent the NFHS, in-house carbonate standard, consisting of planktonic foraminifera. Numbers correspond to foraminiferal species analyzed; numbers 5–14 and 19–21 represent values from previously published species and are included in Fig. S1 (in the Supplement).

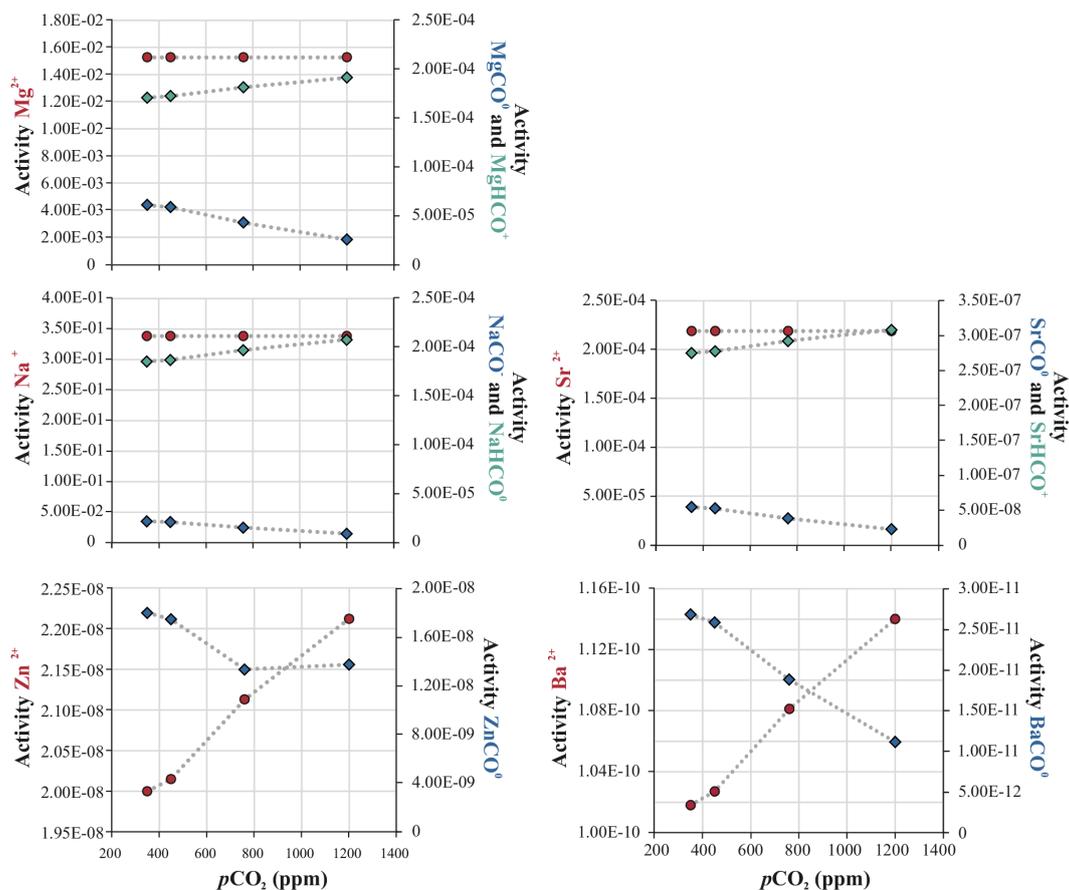
**Table 4.** Regression and  $p$  values of foraminiferal Zn / Ca and Ba / Ca versus  $pCO_2$  values of different species (Fig. 4).

Species	Zn / Ca		Ba / Ca	
	$R^2$	$p$ value	$R^2$	$p$ value
<i>S. marginalis</i>	0.99	< 0.0005	0.81	< 0.025
<i>A. angulatus</i>	0.95	< 0.0025	0.99	< 0.0005
<i>L. bradyi</i>	0.98	< 0.0005	0.97	< 0.0025
<i>A. carinata</i>	0.98	< 0.001	0.94	< 0.005
<i>A. gibbosa</i>	0.99	< 0.0005	0.98	< 0.001

and Ba by McCorkle et al. (1995). Observed responses to changes in carbonation ion concentration are in these studies mainly due to calcification in undersaturated seawater, as described by “the carbonate ion saturation hypothesis” (Elderfield et al., 2006). In some low-Mg benthic species, both  $Mg / Ca_{CALCITE}$  and  $Sr / Ca_{CALCITE}$  do not seem to depend on inorganic carbon system parameters, e.g., pH or  $[CO_3^{2-}]$  (Allison et al., 2011; Dueñas-Bohórquez et al., 2011). However, for several planktonic species pH does in-

fluence  $Mg / Ca_{CALCITE}$  and  $Sr / Ca_{CALCITE}$  (Russell et al., 2004; Lea et al., 1999; Evans et al., 2016). The effect of pH on  $Sr / Ca_{CALCITE}$  might be explained via increased growth rates due to pH-associated changes in  $[CO_3^{2-}]$  (Dissard et al., 2010). However, due to the limited experimental setup, we are not able to disentangle the effects of the different carbon parameters in this study. Still, here we show that incorporation of Mg, Sr and Na of the selected larger benthic hyaline and porcelaneous foraminifera is not significantly impacted when cultured over a limited range of  $pCO_2$  and thus  $[CO_3^{2-}]$  and pH values.

In contrast, foraminiferal Zn / Ca and Ba / Ca are significantly impacted by  $pCO_2$  for all species studied here (Table 4; Fig. 3). Although Hönisch et al. (2011) suggested that the impact of carbonate chemistry on Ba incorporation is negligible, their data do suggest a trend over the same interval in pH as studied here. In hyaline foraminifera, Zn / Ca and Ba / Ca increases more as a function of  $pCO_2$  (factor of 3.7–4.5 and 3.6–3.7, respectively, when  $pCO_2$  increases from 350 to 1200 ppm) compared to the porcelaneous species (1.3–2.1 and 1.6–2.1 times, respectively). This observation



**Figure 5.** Speciation of Mg, Na, Sr, Zn and Ba in the different seawater treatments modeled in PHREEQC (Parkhurst and Appelo, 1999). Activities of free ions (red) and element ( $E$ )–carbonate complexes ( $\text{ECO}_3$ , blue diamonds;  $\text{EHC}_3$ , green diamonds).

suggests that the mechanisms involved in uptake of Ba and Zn are similar for porcelaneous and hyaline species, although the contribution of this process is different, which leads to differences in the sensitivity of Zn and Ba incorporation with  $p\text{CO}_2$

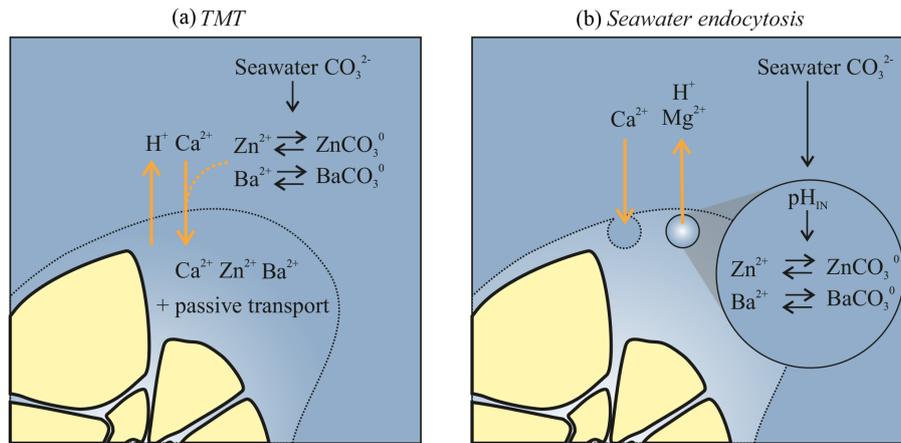
## 4.2 Speciation in the microenvironment

In the culture setup used, increasing  $p\text{CO}_2$  increases DIC, reduces pH and thereby decreases seawater  $[\text{CO}_3^{2-}]$ . Speciation of Zn, Ba and also other elements, like U (Keul et al., 2013; Russell et al., 2004; van Dijk et al., 2017), is primarily controlled by seawater  $[\text{CO}_3^{2-}]$ . The speciation of all elements studied here (Mg, Na, Sr, Zn and Ba) for our different seawater treatments was modeled using PHREEQC, a computer program for speciation, batch reaction, one-dimensional transport, and inverse geochemical calculations (Parkhurst and Appelo, 1999). For this we used the standard in-software LLNL database, a PHREEQC database that implements most of the inorganic aqueous species and minerals in the thermodynamic data and includes many elements not available in any other PHREEQC database. We observed a

decrease in free ions ( $\text{Zn}^{2+}$  and  $\text{Ba}^{2+}$ ) and an increase in Ba and Zn carbonate complexes ( $\text{BaCO}_3^0$  and  $\text{ZnCO}_3^0$ ), with increasing  $p\text{CO}_2$  (Fig. 5), while the activity of  $\text{Mg}^{2+}$ ,  $\text{Na}^+$  and  $\text{Sr}^{2+}$  remained similar. This suggests that element incorporation in foraminiferal calcite might be depending on the availability of free ions, which in the case of Ba and Zn changes with  $p\text{CO}_2$ . This is contrary to inorganic precipitation, where carbonate complexes (e.g.,  $\text{MgCO}_3^0$ ) are easily incorporated into the calcite crystal lattice.

## 4.3 Trends in element incorporation

Element incorporation in hyaline foraminifera is highly interdependent, i.e., species with increased Mg content also incorporate more Sr, Na, Ba and Zn (Fig. 3). The linear increase in Sr incorporation with increasing Mg/Ca has also been observed for other hyaline species, like *Operculina ammonoides* (Evans et al., 2015), *Amphistegina lessonii* and *Ammonia aomoriensis* (Mewes et al., 2015), which both fall into the same trend as that for inorganic calcite (Mucci and Morse, 1983). Evans et al. (2015) hypothesize that the incorporation of other alkali elements such as Na is also re-



**Figure 6.** Schematic overview of different mechanisms for ion transport during foraminiferal calcification. Orange arrows indicate transport of ions. **(a)** The transmembrane transport (TMT) mixing model as proposed by Nehrke et al. (2013). The amount of free ions, e.g.,  $\text{Zn}^{2+}$  and  $\text{Ba}^{2+}$ , available for transport by  $\text{Ca}^{2+}$  in exchange for protons (Toyofuku et al., 2017) is influenced by speciation due to changes in seawater  $[\text{CO}_3^{2-}]$ . The composition of the seawater at the site of calcification is determined by both TMT and passive transport. **(b)** Simplified overview of seawater endocytosis (Bentov et al., 2009; Erez, 2003), where speciation of  $\text{Zn}^{2+}$  and  $\text{Ba}^{2+}$  is determined by changes in internal pH ( $\text{pH}_{\text{IN}}$ ), which changes with ambient seawater carbonate chemistry. Based on our observations, panel **(a)** likely applies to hyaline species, whereas panel **(b)** represents porcelaneous calcification.

lated, due to lattice distortion by the incorporation of Mg, which has been found in inorganic calcite (Okumura and Kitano, 1986). However, although this mechanism might explain some of the observed species-specific element incorporation in hyaline foraminifera, this does not explain the observed difference between hyaline and porcelaneous foraminifera (Fig. 4). Porcelaneous foraminifera have generally high Mg/Ca, but we actually observe lower incorporation of Na and Sr compared to hyaline species with similar  $D_{\text{Mg}}$  (Fig. 4, upper right and left panel:  $D_{\text{Mg}}$  versus  $D_{\text{Na}}$  and  $D_{\text{Sr}}$ ). When including porcelaneous species from other studies we also observe no increase in  $D_{\text{Sr}}$  over a larger range in  $D_{\text{Mg}}$  (Fig. S1, upper left panel:  $D_{\text{Mg}}$  versus  $D_{\text{Sr}}$ ). Therefore, even though this interdependence might partly stem from mechanisms associated with crystallography, differences between hyaline and porcelaneous foraminifera suggest it could also be caused by mechanisms involved in taking up the ions ( $\text{Ca}^{2+}$  and  $\text{CO}_3^{2-}$ ) necessary for chamber formation, which are different for hyaline and porcelaneous species, reflected in the different trends observed here.

#### 4.4 Ion transport models

Both porcelaneous and hyaline foraminifera promote calcification by increasing their internal pH (De Nooijer et al., 2009). Still, they might use different mechanisms to take up the ions ( $\text{Ca}^{2+}$  and  $\text{CO}_3^{2-}$ ) necessary for chamber formation, which is reflected in the different trends observed here. In both porcelaneous and hyaline foraminifera,  $E/\text{Ca}$  responds similarly to changes in  $p\text{CO}_2$  (Fig. 3), suggesting uptake of all these elements is controlled by the same process. However, we observed different inter-element relations

between hyaline and porcelaneous foraminifera (Fig. 4), indicating the mechanisms for ion transport might be different for these two groups. Two of the main concepts of ion transport in foraminifera are transmembrane transport (Nehrke et al., 2013) and the inclusion of seawater by seawater endocytosis (Bentov et al., 2009). Here we try to validate these two concept by comparing them with our observations.

##### 4.4.1 Transmembrane transport

During calcification,  $\text{Ca}^{2+}$  is proposed to be transported from seawater to the SOC via ion channels (Nehrke et al., 2013), likely in exchange for protons (Toyofuku et al., 2017). This so-called transmembrane transport (TMT) through  $\text{Ca}^{2+}$  channels has also been found for other marine organisms, including coccolithophores (Gussone et al., 2006). These  $\text{Ca}^{2+}$  channels may not discriminate perfectly between Ca ions and elements like Mg, Sr, Ba, Na (Sather, 2005; Allen and Sanders, 1994; Hess and Tsien, 1984), causing accidental transport of these elements into the SOC. How much of a certain element will enter the SOC in this way, depends on (1) the selectiveness of the channels and the characteristics of the transported ions (like atomic radius), (2) the element to calcium ratio in the foraminiferal microenvironment and (3) the concentration gradient between seawater and the SOC. The availability of some free ions, like Ba and Zn, changes as a function of  $p\text{CO}_2$  due to the formation of carbonate complexes (Fig. 5). When Zn and Ba form stable complexes with carbonate ions, they are no longer available for (sporadic) transport through the  $\text{Ca}^{2+}$  channels, decreasing the availability at the site of calcification and, subsequently, incorporation into the foraminiferal calcite. Thus,

transmembrane transport of ions by  $\text{Ca}^{2+}$  to the SOC is in agreement with our results (Fig. 3). The amount of Zn and Ba available at the site of calcification is proportional to the concentration of the ratio between  $\text{Ca}^{2+}$  and free  $\text{Zn}^{2+}$  and  $\text{Ba}^{2+}$  in the foraminiferal microenvironment. In turn, the amount of free Zn and Ba ions in seawater is controlled by their respective concentration in seawater concentration, as well as carbonate chemistry (Fig. 5). Foraminiferal Mg / Ca, Na / Ca and Sr / Ca is not detectably affected, since the availability of  $\text{Mg}^{2+}$ ,  $\text{Na}^+$  and  $\text{Sr}^{2+}$  does not change over the range of  $[\text{CO}_3^{2-}]$  studied here. However, the large range in Mg / Ca values in hyaline species suggests that TMT might play a variable role in the calcification process of these species. This may result in an interdependence between all these elements studied such as observed here for the hyaline species if the selectivity for  $\text{Ca}^{2+}$  of these channels varies between species.

#### 4.4.2 Seawater endocytosis

Another proposed ion transport mechanism is seawater endocytosis (Bentov et al., 2009; Erez, 2003), in which seawater is vacuolized, altered and then used as a calcifying fluid. The chemistry or elemental composition of the vacuolized seawater or the calcification fluid is in this case dependent on the seawater concentration. Inter-species differences would therefore be minimized, as is observed for porcelaneous foraminifera in our study (Fig. 4). However, this concept cannot explain the observed changes in Zn / Ca and Ba / Ca as a function of  $p\text{CO}_2$ , which are caused by the speciation of elements in seawater due to changes in the carbonate chemistry (Fig. 5). The concentration of total Zn, for example, in the seawater vacuoles does not change for the different treatments, only the species of Zn present. Only if the internal pH in the vacuole depends on ambient seawater pH, which is currently unknown, is there a potential for changes in speciation. In theory, if pH changes in concert with ambient seawater, such a change in internal pH from  $> 9$  (De Nooijer et al., 2009) to  $> 8.6$  ( $\Delta\text{pH}=0.4$  in our treatment) will change  $[\text{CO}_3^{2-}]$  and thus the speciation of, for example, Zn (and Ba) at the site of calcification. However, over this range the change in  $[\text{CO}_3^{2-}]$  will be rather limited and hence such an effect of differential speciation within the calcifying fluid does not suffice to explain the observed sensitivity of Zn and Ba to  $p\text{CO}_2$  in our study. This is in line with recent evidence regarding Zn / Ca in foraminifera, which suggests Zn incorporation is not primarily governed by changes in seawater pH but by carbonate ion concentration, which does not change much at these high pH's (van Dijk et al., 2017).

#### 4.5 Consequences for calcification in hyaline and porcelaneous species

Both ion transport mechanisms, and their consequences for Zn and Ba incorporation, are summarized in Fig. 6. The

observed species-specific element incorporation in hyaline foraminifera (Fig. 4) is compatible with the transmembrane transport mixing model proposed by Nehrke et al. (2013), where species-specific differences in  $E/\text{Ca}$  are explained by the relative contribution of transmembrane transport and so-called passive transport. In contrast to hyaline species, the porcelaneous species show much less inter-species variation in element composition (Fig. 3), suggesting that this group of foraminifera calcifies from a fluid comparable to ambient seawater (Ter Kuile and Erez, 1987) by, for example, seawater endocytosis (Fig. 6, panel b), with only minor alteration of the elemental composition of the calcifying fluid by ion channels. However, the observed correlation between  $p\text{CO}_2$  and Ba and Zn (Fig. 3) suggests that Ca channels still play a (modest) role in supplying  $\text{Ca}^{2+}$  to the porcelaneous SOC, since possible speciation of minor and trace elements in the SOC caused by a change in the internal pH is probably not sufficient to explain observed patterns (Sect. 4.4.2). However, since porcelaneous species already obtain calcium by including seawater in their calcification vesicle prior to calcite precipitation, contribution of  $\text{Ca}^{2+}$  through TMT is likely smaller than in hyaline species, which may explain the observed lower sensitivity of, for example, foraminiferal Zn / Ca and Ba / Ca to changes in seawater  $[\text{CO}_3^{2-}]$  in porcelaneous species (Fig. 3). This approximately 2 times lower sensitivity of porcelaneous foraminifera compared to hyaline species suggests that porcelaneous foraminifera acquire half of the necessary  $\text{Ca}^{2+}$  through Ca channels compared to hyaline species. Element incorporation in porcelaneous foraminifera will therefore be mainly governed by their respective concentrations in seawater, and to a lesser extent by the selectivity for  $\text{Ca}^{2+}$ /permeability for other ions during TMT.

## 5 Conclusions

Trends in element incorporation in larger benthic foraminifera can be explained by a combination of differences in calcification strategy and seawater chemistry. Carbonate chemistry of seawater determines speciation and therefore availability of some ions (e.g.,  $\text{Zn}^{2+}$  and  $\text{Ba}^{2+}$ ), which are available for ion transport to the site of calcification. For hyaline foraminifera, we observed species-specific interdependence of element incorporation, which can be explained by a previously proposed transmembrane transport model and the bioavailability of ions in seawater during calcification. For porcelaneous foraminifera, species-specific differences are small, suggesting a higher contribution of another ion source, such as seawater endocytosis.

## 6 Data availability

Data are available and can be requested from the corresponding author (inge.van.dijk@nioz.nl).

The Supplement related to this article is available online at doi:10.5194/bg-14-497-2017-supplement.

*Competing interests.* The authors declare that they have no conflict of interest.

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