



# Fast-freezing with liquid nitrogen preserves bulk dissolved organic matter concentrations, but not its composition

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**Abstract.** Freezing can affect concentrations and spectroscopic properties of dissolved organic matter (DOM) in water samples. Nevertheless, water samples are regularly frozen for sample preservation. In this study we tested the effect of different freezing methods (standard freezing at  $-18^{\circ}\text{C}$  and fast-freezing with liquid nitrogen) on DOM concentrations measured as organic carbon (DOC) concentrations and on spectroscopic properties of DOM from different terrestrial ecosystems (forest and grassland). Fresh and differently frozen throughfall, stemflow, litter leachate and soil solution samples were analyzed for DOC concentrations, UV-vis absorption and fluorescence excitation–emission matrices combined with parallel factor analysis (PARAFAC). Fast-freezing with liquid nitrogen prevented a significant decrease of DOC concentrations observed after freezing at  $-18^{\circ}\text{C}$ . Nonetheless, the share of PARAFAC components 1 ( $\text{EX}_{\text{max}} < 250 \text{ nm}$  ( $340 \text{ nm}$ ),  $\text{EX}_{\text{max}}$ :  $480 \text{ nm}$ ) and 2 ( $\text{EX}_{\text{max}}$ :  $335 \text{ nm}$ ,  $\text{EX}_{\text{max}}$ :  $408 \text{ nm}$ ) to total fluorescence and the humification index (HIX) decreased after both freezing treatments, while the shares of component 3 ( $\text{EX}_{\text{max}} < 250 \text{ nm}$  ( $305 \text{ nm}$ ),  $\text{EX}_{\text{max}}$ :  $438 \text{ nm}$ ) as well as  $\text{SUVA}_{254}$  increased. The contribution of PARAFAC component 4 ( $\text{EX}_{\text{max}}$ :  $280 \text{ nm}$ ,  $\text{EX}_{\text{max}}$ :  $328 \text{ nm}$ ) to total fluorescence was not affected by freezing. We recommend fast-freezing with liquid nitrogen for preservation of bulk DOC concentrations of samples from terrestrial sources, whereas immediate measuring is preferable to preserve spectroscopic properties of DOM.

## 1 Introduction

In addition to dissolved organic carbon (DOC) concentrations, properties of dissolved organic matter (DOM) are crucial for its role in biogeochemical cycles of carbon and nutrients as well as for its effect on pollutant dynamics (Bolan et al., 2011). Spectroscopic methods like UV-vis absorption and fluorescence spectroscopy used as single excitation–emission scans, synchronous scans and excitation–emission matrices (EEMs) in combination with different indices and/or parallel factor analysis (PARAFAC) are increasingly applied to characterize chromophoric dissolved organic matter (cDOM) in various environments (e.g., Murphy et al., 2008; Yamashita et al., 2010; Stedmon and Markager, 2005; Graeber et al., 2012; Otero et al., 2007; Traversa et al., 2014; Kalbitz et al., 1999).

The applicability of optical methods for characterizing DOM and the comparability of results in multidisciplinary studies relies on the preservation of samples prior to their analysis. DOM properties depend on many physicochemical and biological boundary conditions, so that artifacts caused by sample storage or sample pre-treatment may be produced easily. For these reasons it is recommended to directly filter samples after collection and store them in the cold and dark prior to measurement as soon as possible (Santos et al., 2010; Spencer and Coble, 2014). However, immediate measurement is often not possible for practical reasons such as a large number of samples or remote or separated sampling sites, so that freezing of filtered DOM samples is often the selected storage method (Murphy et al., 2008; Yamashita et al., 2010;

Graeber et al., 2012). Freezing can affect the physicochemical composition of samples (Edward and Cresser, 1992), so that improved conservation techniques, which avoid or minimize potential artifacts of freezing, are required. During the freezing process, DOM is preferentially excluded from the ice phase and enriched in the remaining liquid phase (Belzile et al., 2002; Xue et al., 2015). The increasing solute concentrations and changing physical conditions in the remaining liquid phase during the freezing process could promote conformational and configurational changes of DOM molecules as well as particle and complex formation depending on DOM composition and sample type (Zaritzky, 2006; Edward and Cresser, 1992). One potential technique for minimizing these effects could be fast freezing with liquid N<sub>2</sub>, by radically reducing the freezing time.

Whereas studies on sample preservation of marine waters (Del Castillo and Coble, 2000; Yamashita et al., 2010a; Conmy et al., 2009) showed only small freezing effects on DOM fluorescence characteristics, research with a variety of freshwater samples produced inconsistent results. Fellman et al. (2008) measured DOC concentrations and UV absorption in fresh and frozen and/or thawed Alaska stream water samples and reported a significant decrease of DOC concentration and specific ultraviolet absorption at 254 nm (SUVA<sub>254</sub>). They recommended freezing as an acceptable storage method for freshwater samples with low DOC concentration and/or low SUVA<sub>254</sub> values. In contrast, Yamashita et al. (2010) observed only minor changes in absorption-based indices after freezing and thawing of Venezuela river water but significant alterations (decrease and increase) for PARAFAC component intensities. A freeze–thaw experiment with water samples from a large number of UK locations conducted by Spencer et al. (2007) showed large and variable changes (decreasing and increasing) in DOM fluorescence intensity and absorbance after freezing and thawing. Likewise Peakock et al. (2015) found strong and inconsistent effects of freezing and thawing on absorbance properties of cDOM in water from bog pools, fen ditches and lakes. In a study of sample preservation on rainwater cDOM fluorescence, Santos et al. (2010) found a decrease of protein-like fluorescence intensity due to freezing.

While many studies investigated the influence of different soil sample pre-treatments on DOC concentrations and DOM composition (e.g., Christ and David, 1994; Sun et al., 2015) only few studies focused on the influence on these properties when using different preservation methods for the extracted soil solutions. Otero et al. (2007) conducted freeze–thaw experiments on salt marsh pore water and found no changes in characteristics of synchronous fluorescence scans.

The impact of sample preservation like freezing seems highly variable depending on sample and DOM characteristics. While most studies focused on samples from marine or freshwater ecosystems, there is a lack of information on sample pre-treatment effects on cDOM properties of water samples from terrestrial ecosystems, especially soil solution. Due

to different sources of DOM in land and water environments (Bolan et al., 2011) and therefore different chemical characteristics, it is unlikely that insights regarding the alterations of samples during storage can be transferred from one sample type to another. To help close this gap, we investigate in this study the influence of freezing and thawing on DOC concentration, spectral absorption and fluorescence properties for a wide range of water samples (throughfall, litter leachate and soil solution) from different terrestrial ecosystems (grasslands and forests). We tested in how far fast-freezing with liquid nitrogen might prevent concentration and partitioning effects and minimize structural changes of DOM. We hypothesized (i) that sample type affects freeze–thaw effects on DOC concentrations and DOM properties, because of different physical and chemical DOM characteristics and therefore different response to changing conditions during freezing and (ii) that fast-freezing with liquid nitrogen reduces these freeze–thaw effects, because it minimizes the freezing time and thus prevents partitioning effects and their physical consequences.

## 2 Material and methods

### 2.1 Study sites

The study was conducted on experimental plots in the Schorfheide Chorin Exploratory of the German “Biodiversity Exploratories”, which were established as platform for large-scale and long-term functional biodiversity research (Fischer et al., 2010). The experimental plots are located in a young glacial landscape in NE Germany with an annual mean temperature of 8 to 8.5 °C and an annual mean precipitation of 500 to 600 mm. The forest plots are dominated either by pine (*Pinus sylvestris* L.) or beech (*Fagus sylvatica* L.) on Cambisols (IUSS working group WRB, 2014). The grassland plots are meadows, pastures and mown pastures on Histosols, Gleysols and Cambisols.

### 2.2 Sampling and sample preparation

For the experiments, we collected solution samples from five forest and three grassland plots on 17 and 18 June 2014 within a bi-weekly 2-day sampling routine of above and below-ground water samples in the DFG priority program “Biodiversity Exploratories”. Together we collected 27 samples for the freezing experiment including six throughfall (TF), five stemflow (SF), five forest litter leachate (LL) as well as six top- and five subsoil solution samples. Volume-weighted composite samples for the experiment were produced from replicated samplers of the same type (e.g., throughfall collectors, shallow suction cups) of one plot in “aged” 500 mL PE bottles. The bottles were bi-weekly used in the field for the same samples, after washing in the dishwasher and with deionized water. TF was sampled with funnel-type collectors (diameter 0.12 m, polyethylene) 0.3 m

above soil surface. We pooled five replicates at grassland and 20 replicates arranged in two lines of 10 samplers in a cross-shaped form at forest sites. To minimize alterations of the sample and contamination such as evaporation, photochemical reactions and algae growth, the sampling bottles were wrapped with aluminium foil and closed with a 1.6 mm polyester mesh and a table-tennis ball. SF was sampled with sliced polyurethane hoses (diameter: 0.04 m) as a collar sealed with a polyurethane-based glue to the bark of three trees per site at approximately 1.5 m height and connected with a polypropylene (PP) or polyethylene (PE) barrel via a PE tube. LL was collected with three zero-tension lysimeters per site (280 cm<sup>2</sup> sampling area) consisting of polyvinyl chloride plates covered with a PE net (mesh width 0.5 mm) connected with PE hoses to 2 L PE bottles stored in a box below ground. We sampled soil solution with nylon membrane (0.45 µm) suction cups (ecoTech, Germany). Three samplers were installed beneath the A horizon (Top) at approximately 10 cm depth. Another three were installed in the B horizon (Sub) in approximately 50 cm depth in the forest plots and 60 or 70 cm depth in the grassland sites. Suction cups were connected to 2 L PE bottles in an insulated aluminium box placed into a soil pit. Soil water was extracted by applying a vacuum of 50 kPa to the PE bottles with an electric pump after each sampling.

After mixing, the samples were transported on ice to the laboratory and stored overnight at 5 °C. We measured pH (Knick, Germany) and electrical conductivity (WTW, Germany) in all samples prior to filtration through ~0.7 µm glass microfiber filters (Whatman GF/F). The filters were washed with 100 mL deionized water and 10 mL of sample before sample filtration. The filtered sample was split in three aliquots for different preservation treatments: (i) no preservation (fresh) for which samples were stored at 5 °C in the dark and DOC concentrations were measured 24 h after sampling while fluorescence as well as absorbance were measured within 48 h; (ii) preservation by freezing for which the samples were stored at -18 °C for 4 weeks, and (iii) fast-freezing with liquid nitrogen (N<sub>2</sub>), for which 12 mL sample aliquots were filled in pre-rinsed 15 mL (5 mL sample) PP falcon tubes, dipped in liquid nitrogen for 30 s and then stored at -18 °C for 42 days. Fresh samples and samples frozen at -18 °C were stored in 20 mL PE scintillation vials (NeoLab) that were pre-rinsed with 5 mL sample before filling. Fluorescence, absorbance and DOC concentration from all frozen samples were measured after defrosting over night at 5 °C in the dark. For all preparation steps and treatments control samples of ultrapure water (EVOQUA, Germany) were analyzed, showing no release of DOM (DOC concentration and DOM fluorescence) from laboratory equipment.

### 2.3 Laboratory analysis

We measured the concentration of DOC as non-purgeable organic carbon on a Shimadzu TOC-5050A (Duisburg, Ger-

many) with a limit of quantification of 2 mg C L<sup>-1</sup>. Absorption spectra of DOM were scanned at wavelength from 400 to 600 nm using a Lambda 20 UV-vis spectrometer (Perkin Elmer, USA) and a 1 cm quartz cuvette. Absorbance measurements were baseline corrected using ultrapure water. All fluorescence EEMs were measured on a Hitachi F-4500 fluorescence spectrometer (Hitachi, Japan) directly after absorption measurement in the same cuvette. We measured excitation from 240 to 450 nm (5 nm steps) and emission from 300 to 600 nm (2 nm steps) with a slit width of 5 nm and scan speed 12000 nm min<sup>-1</sup>. We corrected our EEMs according to the protocol from Murphy (2010) with the *fdomcorrect* function in the *drEEM* toolbox (version 2.0) of Murphy et al. (2013) using Matlab (Version Matlab2011b, The MathWorks Inc.). We used the correction curves supplied by the manufacturer for the excitation and emission correction factors. We measured ultrapure water fluorescence spectra for blank correction and to convert EEMs to Raman units by normalizing them to the area under the Raman peak at 350 nm excitation wavelength (Lawaetz and Stedmon, 2009). In order to apply the inner-filter correction of Lakowicz (2006) integrated in the *drEEM* toolbox, all aliquots were diluted with ultrapure water to achieve an absorption of <0.3 at 254 nm (Ohno, 2002). For this reason, not all treatments of one sample were diluted with the same dilution factor. To test the possible influence of different dilutions on the pH-related changes in fluorescence (Patel-Sorrentino et al., 2002; Baker et al., 2007), dilution series with samples (*n* = 14) from the same plots and same sample types but with different sampling dates were measured for pH, absorption and fluorescence according to the protocol described above. We compared the differences of 31 dilutions and calculated the mean absolute deviation (MAD). These were compared to the MAD of measurement precision, determined by analyzing 11 samples in three replications. For the PARAFAC components %C1, %C2 and %C3 and SUVA<sub>254</sub> the MAD caused by dilution were less or equal than the precision MAD, so that there was no influence of dilution on the three humic-like components and the specific UV absorbance at 254 nm. For %C4 and HIX the effect of dilution could exceed the precision of fluorescence measurements. For detailed information see supporting information.

### 2.4 Spectroscopic indices and PARAFAC modeling

Based on the absorbance spectra, we calculated specific ultraviolet absorbance (SUVA<sub>254</sub>) as the absorbance at 254 nm divided by the DOC concentration. The SUVA<sub>254</sub> is reported in L mg<sup>-1</sup> m<sup>-1</sup>, and is associated with bulk aromaticity (Weishaar et al., 2003). Moreover, we calculated the humification index (HIX) from fluorescence EEMs (Ohno, 2002). The HIX ranges from 0 to 1 and allows characterizing samples based on their degree of DOM humification.

In addition to the calculation of indices, we used parallel factor analysis (PARAFAC) to mathematically decom-

pose the trilinear data of the EEMs into fluorescence components of DOM (Stedmon et al., 2003). Further pre-processing steps of EEMs (smoothing of Rayleigh and Raman scatter and sample normalization), as well as the PARAFAC analysis, were conducted with the drEEM toolbox (version 2.0, Murphy et al., 2013). We chose a four component PARAFAC model (components referred as C1 to C4), visually checked the randomness of residuals and the component spectral loadings, split-half validated the model and generated the best fit by random initialization. For comparison in statistical analysis we used the relative percentage distribution of the four PARAFAC components (% of the sum of total peak fluorescence of all PARAFAC components), so that percentage values for the components will be given as %C1 to %C4.

## 2.5 Statistical analysis

The DOM composition variables used for statistical analysis were the PARAFAC components %C1 to %C4, the spectroscopic indices HIX and SUVA<sub>254</sub>, as well as the DOC concentration. For all statistical analysis the variables were scaled and centered. We conducted a pair-wise (samples as strata) permutational multivariate analysis of variance (PERMANOVA) with DOC concentrations of the fresh samples as factor based on Euclidean distances in R (Oksanen et al., 2015; R core team, 2015). The adonis function was used to assess the influence of sample preparation (fresh, frozen, fast-freezing) and of the initial DOC concentration on DOM variables. To investigate preservation effects on single variables we conducted linear mixed-effect models (sometimes called multi-level models, lme function, Linear and Nonlinear Mixed Effects Models package for R, Pinheiro et al., 2015) with samples as random intercept on each of the DOM composition variables. These were used instead of simple linear models or ANOVAs, since we could not expect the same intercept for all samples due to different sample concentrations. To test the influence of the initial DOC concentration on single preservation treatments we performed Spearman Rank Order Correlation. To assess the influence of sample type (TF, SF, LL, Top or Sub) on the relative change of DOM composition due to fast-freezing with liquid nitrogen or freezing at  $-18^{\circ}\text{C}$  in relation to the measurement of fresh, cooled samples, we used an ANOVA with the sample type as fixed factor (aov function in R). To remove sample concentration-related effects and to calculate relative changes, the differences between the two preservations (either fast-freezing or freezing at  $-18^{\circ}\text{C}$ ) relative to the measurements of fresh samples were calculated for each sample before the ANOVA. This was only done for variables, for which we found strong, significant effects with the linear mixed-effect models.

## 3 Results

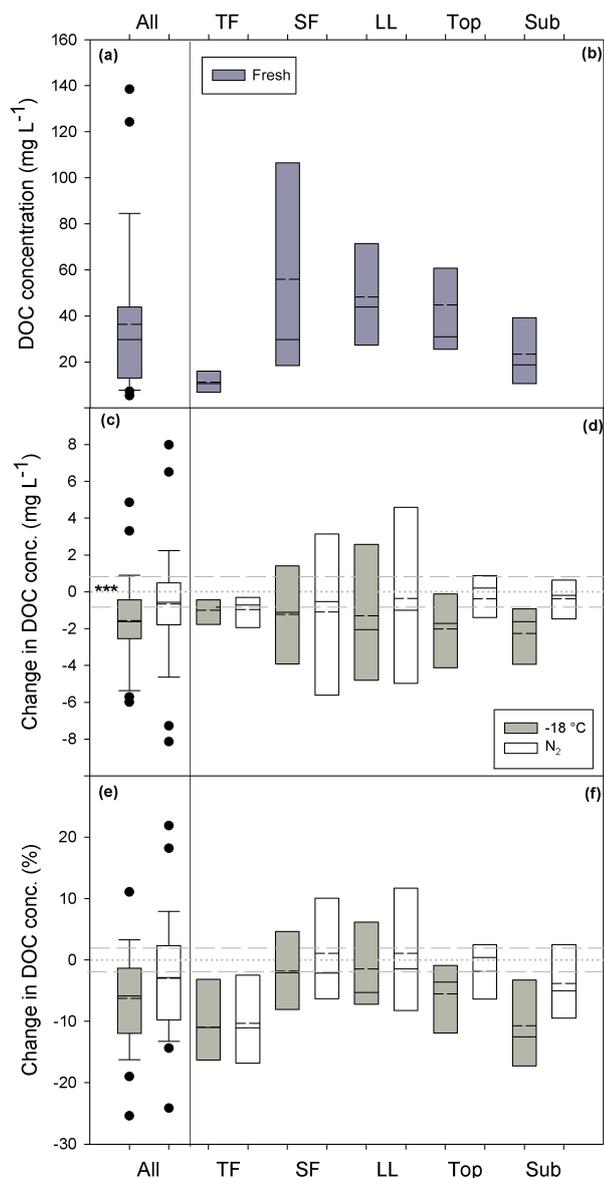
### 3.1 DOM concentrations

The samples covered a wide range of DOC concentrations (Fig. 1a, b). Fresh TF samples showed the lowest concentrations ranging from 5 to 17 mg C L<sup>-1</sup>, SF samples had the highest DOC concentrations ranging from 12 to 138 mg C L<sup>-1</sup> (Fig. 1b). High concentrations up to 75 mg C L<sup>-1</sup> were also found for LL samples, but average values were smaller than for SF (Fig. 1b). In the mineral soil, concentrations decreased from 13 to 124 mg C L<sup>-1</sup> in topsoil samples to 9 to 47 mg C L<sup>-1</sup> in subsoil samples.

We found a significant treatment effect (linear mixed-effect models (lme),  $p < 0.05$ ) on DOC concentration when comparing the fresh and frozen samples (Fig. 1c). In 24 of 27 samples DOC concentrations decreased after freezing at  $-18^{\circ}\text{C}$  and subsequent thawing, with an average change of  $-1.6$  mg C L<sup>-1</sup> or  $-6\%$  respectively. The maximum decrease that was found equalled  $-6$  mg C L<sup>-1</sup> and  $-25\%$ , respectively. In contrast to freezing at  $-18^{\circ}\text{C}$ , fast-freezing with liquid nitrogen did not result in significant changes (lme,  $p > 0.05$ ) of DOC concentrations (Fig. 1c). This different behavior between normal freezing and fast-freezing was also found for the influence of the initial DOC concentration on changes of DOM properties. Only the  $-18^{\circ}\text{C}$  treatment showed a significant correlation (Spearman's rank  $r = -0.447$ ,  $p = 0.0194$ ), indicating a larger decrease of DOC concentrations due to freezing for samples with higher initial DOC concentrations.

### 3.2 PARAFAC fluorescence components

The analysis of fluorescence spectra using PARAFAC resulted in four components that were characterized according to the review of Fellman et al. (2010) (Table 1). C1 exhibited its main excitation maximum at  $< 250$  nm, a secondary maximum at 340 nm and an emission maximum at 480 nm and was described as UVA humic-like fluorophore with a terrestrial source and a high molecular weight (Murphy et al., 2006; Stedmon et al., 2003; Shutova et al., 2014; Fellman et al., 2010). C2 had a maximum excitation at 335 nm and an emission maximum at 408 nm and was named also UVA humic-like, but associated with low molecular weight (Murphy et al., 2006; Fellman et al., 2010; Stedmon et al., 2003). C3 was defined by an excitation maximum at  $< 250$  nm, a secondary maximum at 305 nm and an emission maximum at 438 nm. This component dominated fulvic acid fractions of humic substances (Santín et al., 2009; He et al., 2006). Finally, C4 was characterized by its excitation maximum at 280 nm and an emission maximum at 328 nm and was classified as tryptophan-like, as its fluorescence resembles free tryptophan. Therefore, this component was associated with free or bound proteins (Fellman et al., 2010).

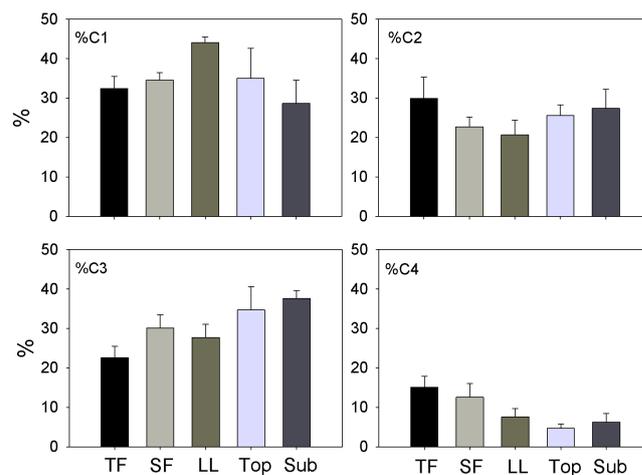


**Figure 1.** Absolute DOC concentrations (measured in fresh samples) and changes in DOC concentrations after freezing ( $-18^{\circ}\text{C}$ ) and fast-freezing with liquid nitrogen; (a, c, e) all samples ( $n = 27$ ); (b, d, f) ordered by sample type (throughfall (TF)  $n = 6$ , stem-flow (SF)  $n = 5$ , litter leachate (LL)  $n = 5$ , top soil solution (Top)  $n = 6$ , sub-soil solution (Sub)  $n = 5$ ); gray dashed line: analytical reproducibility; \*\*\* significant changes (linear mixed models (lme),  $p < 0.05$ ); boxplots: solid line: median, dashed line: mean.

We found different distributions of PARAFAC components for different sample types (Fig. 2). The contribution of %C1 to the total fluorescence increased from TF over SF to LL and then decreased again from LL to Sub (Fig. 2), while %C2 showed just the opposite trend. In contrast, %C3 tended to increase from TF to Sub, whereas %C4 showed a decreasing trend (Fig. 2).

**Table 1.** Characteristics of PARAFAC components based on Fellman et al. (2010).

Component	Maximum excitation wavelength ( $\text{EX}_{\text{max}}$ ) (nm)	Maximum emission wavelength ( $\text{EM}_{\text{max}}$ ) (nm)	Description
C1	<250 (340)	480	humic-like, terrestrial
C2	335	408	humic-like
C3	<250 (305)	438	fulvic-acid-type
C4	280	328	tryptophan-like

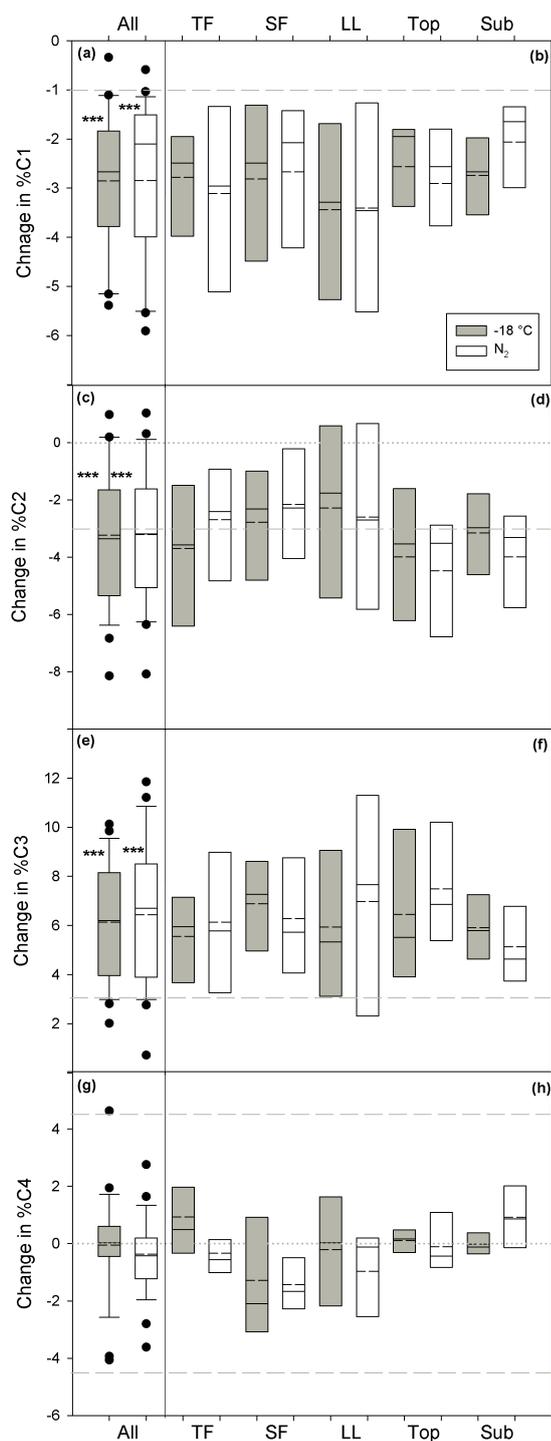


**Figure 2.** Mean distribution of PARAFAC components %C1–%C4 for different sample types.

The conducted PERMANOVA was highly significant ( $p < 0.001$ ), indicating that the preservation significantly affects the DOM composition. The interaction between treatment and initial DOC concentration of the fresh treatment explains a reasonable part of the variance ( $R^2 = 0.14$ ) and is highly significant ( $p < 0.001$ ). Therefore the original DOC concentration of the fresh sample well explains the variable strength of the treatment effect.

Similar changes in component distribution were found as a consequence of freezing at  $-18^{\circ}\text{C}$  and fast-freezing with liquid nitrogen (Fig. 3). We observed a significant (lme,  $p < 0.05$ ) decrease in all samples for the relative fraction of the humic-like components %C1 and %C2 after freezing at  $-18^{\circ}\text{C}$  and fast-freezing compared to the fresh control samples (Fig. 3a, b). The contribution of %C1 to the total fluorescence decreased on average by  $-3\%$  with maximum changes of  $-5\%$  for freezing at  $-18^{\circ}\text{C}$  and  $-6\%$  for fast-freezing with liquid nitrogen. The average decrease of %C2 was  $-3\%$  and the maximum  $-8\%$  for both treatments.

In contrast to %C1 and %C2, the share of %C3 to the total fluorescence intensity increased upon freezing (Fig. 3e, f). All samples frozen at  $-18^{\circ}\text{C}$  showed an increase in the relative intensity of the %C3 signal, with an average increase of  $+6\%$  for both treatments. The maximum increase was  $10\%$  (freezing at  $-18^{\circ}\text{C}$ ) and  $12\%$  (freezing with liquid  $\text{N}_2$ ).



**Figure 3.** Changes of relative distribution of PARAFAC components after freezing ( $-18^{\circ}\text{C}$ ) and fast-freezing with liquid nitrogen; (a, c, e, g) all samples ( $n = 27$ ); (b, d, f, h) ordered by sample type (throughfall (TF)  $n = 6$ , stemflow (SF)  $n = 5$ , litter leachate (LL)  $n = 5$ , top soil solution (Top)  $n = 6$ , sub-soil solution (Sub)  $n = 5$ ); gray dashed line: analytical reproducibility; \*\*\* significant changes (linear mixed models (lme),  $p < 0.05$ ); boxplots: solid line: median, dashed line: mean.

No significant effects of sample preservation (lme,  $p > 0.05$ ) were found for %C4, the protein-like component (Fig. 3g, h).

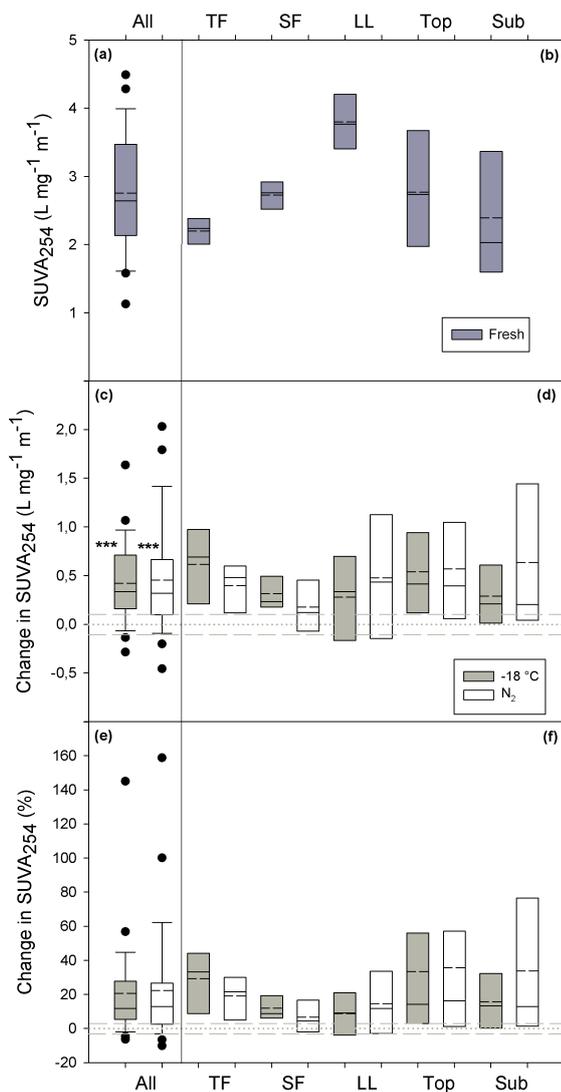
### 3.3 Aromaticity and humification index

We found  $\text{SUVA}_{254}$ -values ranging from  $1.1 \text{ L mg}^{-1} \text{ m}^{-1}$  up to  $4.5 \text{ L mg}^{-1} \text{ m}^{-1}$  for fresh samples (Fig. 4a, b). Samples frozen at  $-18^{\circ}\text{C}$  and fast-frozen samples showed a significant increase (lme,  $p < 0.05$ ) of their  $\text{SUVA}_{254}$  (Fig. 4c). The average change was  $+0.4 \text{ L mg}^{-1} \text{ m}^{-1}$  equivalent to  $+20\%$  for samples frozen at  $-18^{\circ}\text{C}$  and  $+0.5 \text{ L mg}^{-1} \text{ m}^{-1}$  equivalent to  $+24\%$  for samples that were fast-frozen with liquid nitrogen.

The humification index of the freshly measured samples ranged from 0.806 to 0.931 in TF and SF samples and from 0.849 to 0.975 for Sub, Top and LL samples (Fig. 5a, b). We found a significant decrease (lme,  $p < 0.05$ ) of the HIX when comparing the freshly measured samples with the frozen and the fast-frozen samples (Fig. 5c). The average change was  $-0.016$  or  $-2\%$  for samples frozen at  $-18^{\circ}\text{C}$  and  $-0.020$  or  $-2\%$  for samples fast-frozen with liquid nitrogen. The maximum decrease was  $-0.128$  or  $-15\%$  for  $-18^{\circ}\text{C}$  samples and  $-0.076$  or  $-8\%$  for liquid nitrogen samples (Fig. 5c, d, e, f).

## 4 Discussion

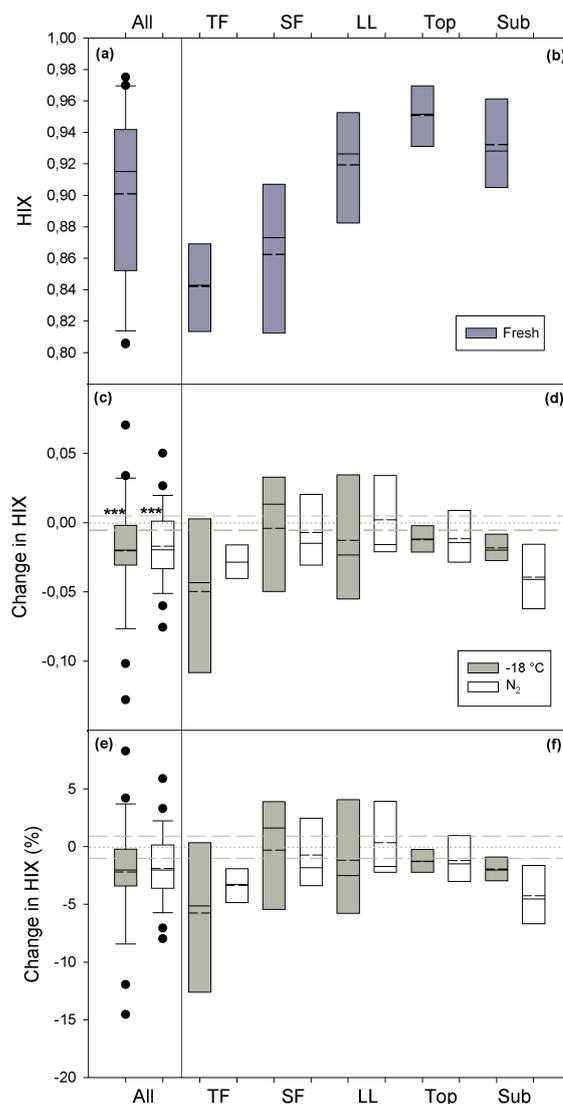
We found that freezing at  $-18^{\circ}\text{C}$  significantly reduced DOC concentrations across all sample types and that the effect is higher with higher initial DOC concentrations. This is in line with results of Fellman et al. (2008) investigating the effect of freezing and thawing on Alaskan stream water samples. This loss of DOC concentration might be due to aggregation and irreversible particle formation (Giesy and Briese, 1978) induced by partitioning and concentration effects during the freezing process (Belzile et al., 2002; Xue et al., 2015). Indeed, our results indicated that fast-freezing with liquid nitrogen can prevent significant reductions of bulk DOC for samples with a large range of DOM concentrations. In contrast to effects on DOC concentrations, we found similar significant effects of fast-freezing as well as freezing at  $-18^{\circ}\text{C}$  on the chromophoric humic fraction of DOM (PARAFAC components, HIX and  $\text{SUVA}_{254}$ ). The increase of aromaticity as indicated by higher  $\text{SUVA}_{254}$  values indicates a stronger removal of non-aromatic DOM during freezing and thawing. On the other hand, the decrease in the HIX suggests a preferential removal of humified cDOM. One potential explanation for the fact that fast-freezing in liquid nitrogen resulted in significant changes of DOM fluorescence properties, but only small changes of bulk DOC concentrations, is that cDOM reacted stronger to freezing and thawing than the remaining DOM so that spectroscopic properties were affected, but bulk DOC concentrations were not. Fast freezing may have failed to prevent changes of cDOM composition because (i)



**Figure 4.** Absolute values (measured in fresh samples) and changes of SUVA<sub>254</sub> after freezing ( $-18^{\circ}\text{C}$ ) and fast-freezing with liquid nitrogen; (a, c, e) all samples ( $n = 27$ ); (b, d, f) ordered by sample type (throughfall (TF)  $n = 6$ , stemflow (SF)  $n = 5$ , litter leachate (LL)  $n = 5$ , top soil solution (Top)  $n = 6$ , sub-soil solution (Sub)  $n = 5$ ); gray dashed line: analytical reproducibility; \*\*\* significant changes (linear mixed models (lme),  $p < 0.05$ ); boxplots: solid line: median, dashed line: mean.

cDOM changes occurred not only during the freezing process ( $-18$  or  $-196^{\circ}\text{C}$  in liquid nitrogen), but also in frozen state at  $-18^{\circ}\text{C}$  in the freezer during storage or (ii) cDOM was affected by the thawing process that was identical for both freezing treatments. The former might be supported by a re-crystallization of ice crystals in frozen state (Luyet, 1967; Meryman, 2007).

No significant changes of protein-like fluorescence (%C4) due to freezing and thawing were observed. This is in contrast to the results of Spencer et al. (2007) and Santos et



**Figure 5.** Absolute values (measured in fresh samples) and changes of HIX after freezing ( $-18^{\circ}\text{C}$ ) and fast-freezing with liquid nitrogen; (a, c, e) all samples ( $n = 27$ ); (b, d, f) ordered by sample type (throughfall (TF)  $n = 6$ , stemflow (SF)  $n = 5$ , litter leachate (LL)  $n = 5$ , top soil solution (Top)  $n = 6$ , sub-soil solution (Sub)  $n = 5$ ); gray dashed line: analytical reproducibility; \*\*\* significant changes (linear mixed models (lme),  $p < 0.05$ ); boxplots: solid line: median, dashed line: mean.

al. (2010), which could be related to similar fluorescence characteristics, but different chemical composition of proteinaceous fluorescence material from aquatic sources and the solutions from terrestrial ecosystems tested in this study.

In our experiment we used relative small sample volumes (fresh,  $-18^{\circ}\text{C}$ : 20 mL,  $\text{N}_2$ : 12 mL) because we commonly keep the volume that is stored frozen as small as possible due to space limitations in deep freezers. We think that increasing the volume of samples that are subjected to freezing

also increases the risk of artifacts, because of increasing concentration effects due to extended freezing time.

## 5 Conclusions

Freezing and thawing affected the DOC concentration, spectral absorption and fluorescence properties of water samples (throughfall, litter leachate and soil solution) from different terrestrial ecosystems (grasslands and forests). In contrast, fast-freezing with liquid nitrogen minimized the changes of bulk DOC concentrations but not the changes of spectroscopic cDOM properties. Different thawing protocols for minimizing sample storage effects on DOM should be tested in future studies. We suggest the use of fast-freezing for preservation of bulk DOC concentrations, especially for highly concentrated samples, when the increased effort and cost of using liquid nitrogen in the field is justified by advantages regarding the minimization of freezing artefacts. To preserve cDOM characteristics of samples from terrestrial sources normal freezing or fast-freezing should be avoided. Instead, filtration, cooling and measurements soon after the sampling should be the method of choice, if possible.

### Information about the Supplement

The data related to statistical analysis and for generating the figures are available in the Supplement.

**The Supplement related to this article is available online at doi:10.5194/bg-13-4697-2016-supplement.**

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