Analysis of passive-sampler monitored atmospheric ammonia at 74 sites across southern Ontario, Canada

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Received: 11 July 2013 – Published in Biogeosciences Discuss.: 5 August 2013
Revised: 27 October 2013 – Accepted: 6 November 2013 – Published: 4 December 2013

Abstract. Weekly/biweekly concentrations of atmospheric NH\(_3\) were collected using passive samplers at 74 sites across southern Ontario, Canada, during the period from June 2006 to March 2007 with eighteen sites running as early as March 2006. The annual average of NH\(_3\) (AAN) at all the sites across southwestern Ontario was over 1 µg m\(^{-3}\), a value that was recently proposed as the new critical level for protecting vegetation. High AAN values (3.6–6.1 µg m\(^{-3}\)) were observed at eight sites located inside the intensive livestock production zones. AAN values at the sites across southeastern Ontario were generally lower than 1 µg m\(^{-3}\) and the values were less than 0.4 µg m\(^{-3}\) at non-agricultural sites. Agricultural sites generally showed the highest concentrations in spring to summer and the lowest in winter, but the occurrence time of the peak concentration varied among the sites, which could be due to their different land use and thus different fertilizer application intensities or periods. Seasonal trends of NH\(_3\) concentration at some non-agricultural sites were somewhat similar to those at the agricultural sites, but the peak concentration appeared a few weeks later than at the agricultural sites. Surprisingly, at several remote non-agricultural sites, NH\(_3\) concentrations were not decreasing in winter and were sometimes increasing to a level higher than those in summer, which could be caused by different transport mechanisms in different seasons. A sharp increase in NH\(_3\) concentration was also surprisingly observed at 20 out of the 74 sites during the coldest two weeks when ambient temperature was lower than \(-7^\circ\)C, and cannot be explained by known sources or with existing knowledge. The NH\(_3\) emission inventory for the region of southern Ontario was evaluated using the measurement data and emissions within two small zones in this region were identified to be potentially underestimated.

1 Introduction

Atmospheric ammonia (NH\(_3\)) is an important reactive nitrogen compound in the geosphere and biosphere of the earth. NH\(_3\) gas has been found to cause direct adverse effects on plant community composition even at low concentration levels and the value of 1 µg m\(^{-3}\) has been proposed as a new critical level for protecting vegetation (Cape et al., 2009). NH\(_3\) gas can react with acidic gases to form ammonium salts and increase concentration of fine particles in the atmosphere (Chan and Yao, 2008; Yao and Zhang, 2012; Zhang et al., 2007, 2012). In the past several decades the decrease of ammonia emission is slower than that of acidifying sulfur and nitrogen species in many countries. Thus, the impact of ammonia on ecosystems attracts more attention (Galloway et al., 2008; Herridge et al., 2008; Sutton et al., 2009).

The origins of NH\(_3\) vary, depending on land use. For example, in intensive agricultural zones, agricultural activities such as livestock farming and N fertilizer application are generally recognized as the major contributors to NH\(_3\) (Beusen et al., 2008; Salazar et al., 2012; Sheppard et al., 2010; Sutton et al., 1998; Zbieranowski and Aherne, 2012, 2013). In remote areas, NH\(_3\) can originate from natural emissions associated with microbial activity or atmospheric transport (Sutton et al., 2009). Because NH\(_3\) is very reactive and is also removed fast through dry and wet deposition, its concentration decreases rapidly with distance from the sources (Theobald et al., 2012). However, particulate ammonium formed from NH\(_3\) can be transported far away from sources (Asman et al., 1998; Wen et al., 2013) and semi-volatile ammonium salts can be transformed back to NH\(_3\) under certain meteorological and chemical conditions (Seinfeld and Pandis, 2006; Yao et al., 2011; Zhang et al., 2008).
Assessing the impact of NH$_3$ on ecosystem health, air quality and climate requires accurate prediction of ammonia and ammonium distributions at local and regional scales. NH$_3$ emission inventories are thus needed as input for chemical transport modes. To improve the outdated NH$_3$ emission inventory in southern Ontario, Environment Canada established a high spatial-resolution monitoring network comprising 74 sites and measured weekly/biweekly NH$_3$ concentrations during the period from June 2006 to March 2007 (Lillyman et al., 2009). The measurements of NH$_3$ at eighteen of these sites were made as early as March 2006. The data set not only allows an evaluation and improvement of the Canadian NH$_3$ emission inventory for the region of southern Ontario, but also allows a better understanding NH$_3$-related sciences through in-depth data analysis and data-model comparison (e.g., Wen et al., 2013).

The purpose of the present study is to understand the geographical and seasonal patterns of NH$_3$ in southern Ontario and to identify factors controlling these patterns through comprehensive data analysis. In the following sections, the geographical distributions of NH$_3$ are first presented so the major sources of NH$_3$ can be revealed (Sect. 3.1). The study region is then split into different concentration zones and the seasonal variations of NH$_3$ in the different zones are studied in details (Sects. 3.2–3.4). To understand the phenomena identified in Sect. 3, various explanations were explored (Sect. 4). Finally, a brief evaluation of Environment Canada’s most recent NH$_3$ emission inventory was conducted (Sect. 5) using knowledge obtained in Sects. 3 and 4.

2 Experimental

Funded by the National Agri-Environmental Standards Initiative (NAESI), a project known as the Southern Ontario Ammonia Passive Sampler Survey (SOAPSS) was launched in March 2006 to monitor atmospheric NH$_3$ in southern Ontario, Canada (Lillyman et al., 2009). A total of 74 sites were selected and evenly distributed at agriculture zones and remote non-agriculture zones across southern Ontario with distances between the sites of approximately 20 km (Fig. 1a). Commercially available Ogawa passive samplers were used to monitor NH$_3$ in the high concentration areas of the network during the warm months; slightly larger custom-designed passive samplers were used to monitor NH$_3$ in the low concentration areas and throughout the winter period. The description of the Ogawa sampler and sampling methodology is documented in Roadman et al. (2003). For quality control purposes, the two types of passive samplers were compared against the NH$_3$ gas analyzer and denuder at one selected site (Egbert) for one year period before the campaign started. As shown in Fig. S1, the concentrations of NH$_3$ measured by Ogawa passive samplers and custom-designed passive samplers were both comparable to those collected by the other instruments. No systematic differences were found between the two types of passive samples used in the present study. Besides, the differences between the passive samples and the active samples were generally less than 20 %, which is within the proposed limit for passive samplers for NH$_3$ (Sutton et al., 2001). Comparability of passive samplers to complex methods has also been well established in literature (Cape et al., 2004; Puchalski et al., 2011; Sutton et al., 2001).

From March to May 2006, only eighteen sites started running for consistency check and sampler evaluation. All the 74 sites measured ambient concentration of NH$_3$ starting June 2006. Thus, only approximately 20 % of sites had the NH$_3$ data between March and May 2006 despite the fact that the strongest NH$_3$ emissions occur in this region in the month of May (Lillyman et al., 2009). Note that weekly samples were collected before 5 December 2006 while biweekly samples were collected after this date in expectation of low concentration of NH$_3$ in the cold seasons. The project provided highly spatial-resolved atmospheric NH$_3$ concentration data although with relatively low temporal resolution.

3 Results

3.1 Geographical distribution of NH$_3$ concentration

According to recent publications, more than 80 % of NH$_3$ emissions in southern Ontario are from agricultural activity while non-agricultural sources (e.g., industrial sources, vehicle emissions, open sources (except animals), non-industrial fuel combustion, miscellaneous and incineration) contribute to the remaining < 20 % (Lillyman et al., 2009; Environment Canada, 2013). For the agricultural generated emissions, livestock production accounts for approximately 80 %, with the remaining 20 % being mainly associated with fertilizer application. However, two potential sources are not included in the emission inventory of NH$_3$ in Canada (Lillyman et al., 2009). One is biogenic N fixation in soil followed by soil N cycling processes that release NH$_3$ from soil and/or plants to the atmosphere (Dabney and Bouldin, 1990). The other is atmospheric deposition of NO$_3$ to soil followed by soil N cycling processes that release NH$_3$ to the atmosphere. These emissions can be considered as natural emission or re-emission of predeposited N that is converted to NH$_3$.

Annual average NH$_3$ concentration (AAN – calculated using nine to 12 month data depending on data availability for individual sites) was first compared with the latest emission inventory for a consistency check, and was then analyzed for geographical distributions. The geographical distribution of AAN (Fig. 1b) was generally consistent with that of NH$_3$ emissions across southern Ontario (Fig. S2), as can be seen from their similar geographical patterns. Hot spots in both concentration and emission were centered in the agricultural zones (Figs. S2 and S3). A correlation analysis between AAN and the gridded emission extracted from the
grids containing the sites shows a high correlation with the square of correlation coefficient value of 0.73 (Fig. 1c).

The highest AAN was observed at eight sites (Arkona (ARK), Lynden (LYD), Mildmay (MLM), Milverton (MVT), New Dundee (NDD), St. Marys (SMY), Sweaburg (SWB) and Teviotdale (TEV)) and ranged from 3.6 to 6.1 µg m$^{-3}$ (Fig. 1b). These sites were in a region of active livestock production (Fig. S3, Zbieranowski and Aherne, 2012, 2013). AAN in agricultural zones was generally larger than 1 µg m$^{-3}$, except at seven sites (Galbraith (GAL), Holland (HOL), Lyndhurst (LYH), Merrickville (MER), Montebello (MON), Maynard (MYN) and Westport CAPMON (WPT)) where an AAN in the range of 0.4–0.7 µg m$^{-3}$ was observed. The low AAN values were likely associated with low intensity of agricultural activity. AAN lower than 1 µg m$^{-3}$ was observed at rural and remote non-agricultural-zone sites (highlighted in black in Fig. 1b).

Based on the proposed critical level of 1 µg m$^{-3}$ for protecting vulnerable vegetation such as lichens and mosses and critical level of 3 µg m$^{-3}$ for other vegetation (Cape et al., 2009), the 74 monitoring sites were classified into three concentration categories for easy interpretation in Sect. 3.2–3.4 (i.e., high (AAN > 3 µg m$^{-3}$), medium (1–3 µg m$^{-3}$) and low (< 1 µg m$^{-3}$) concentration categories). Category I (AAN > 3 µg m$^{-3}$) has eight sites as listed above. At seven of the eight sites, weekly/biweekly concentrations higher than 10 µg m$^{-3}$ were observed in a few samples. Category 2 (1 to 3 µg m$^{-3}$) has a total of 36 sites and Category 3 (< 1 µg m$^{-3}$) has the remaining 30 sites.

It is noticed that sites belonging to the same concentration category did not necessarily have the same temporal pattern. For example, even the eight sites in Category 1 which are situated close to each other had different temporal patterns. The seasonal patterns in each category are discussed in detail below. Thus, the 74 sites were further split into seven classes using hierarchical cluster analysis so the geographical distribution of sites having similar temporal patterns and concentration ranges can be easily identified (Fig. 1d). The eight sites in Category 1 took the first four classes because of their very high concentrations and different temporal patterns (Zone A in Fig. 1b and d). Class 5 consists of six sites (Inwood (INW), Mildmay (MIL), Eberts (EBT), West Lorne (WLO), Hudson (HUD), Brigden (BDG)) and had an AAN ranged from 1.8 to 2.6 µg m$^{-3}$. Four of the six sites in Class 5 (lower-left corner in Fig. 1d) were located near the strong NO$_x$ emissions (Fig. S4). NO$_2$ can be oxidized to HNO$_3$ which can then possibly form NH$_4$NO$_3$ in the atmosphere, thereby leading to depletion of atmospheric NH$_3$. Conversely, higher
atmospheric deposition of NO$_x$ in high NO$_x$ emission areas can add N to soils which could be a potential emission source of NH$_3$. In Class 6, 12 sites in southwestern Ontario and one site $\sim$ 500 km northeast were grouped together. Their AAN ranged from 1.3 to 2.9 µg m$^{-3}$. Of these, 12 sites were scattered over $\sim$ 200 km range and encircling sites in Classes 1 to 5 (Fig. 3b). Similar land use and simultaneous agricultural activity at the 12 sites probably led to their consistent temporal variations of NH$_3$ concentrations, leading to them being grouped together.

The remaining 47 sites were defined as Class 7, of which 43 sites were situated in southeastern Ontario and 4 sites were located near Lake Erie. Their AAN ranged from 0.2 to 1.8 µg m$^{-3}$. The dominant wind direction in southwestern Ontario was from the west and southwest (Fig. 2a and Bein et al., 2008). Under the dominant wind condition, most of the sites in Class 7 are located downwind of those zones with higher concentrations of NH$_3$. Regional transport of semi-volatile ammonium salts followed by evaporation or regional transport of NH$_3$ could contribute to the NH$_3$ observed at most of sites in Class 7 as discussed later, leading to temporal variations of ambient NH$_3$ concentration in Class 7 being different from those in Class 6. Note that the 36 sites in Category 2 cover all sites in Classes 5 and 6 and also include 17 out of 47 sites in Class 7, and the 30 sites in Category 3 all belonging to Class 7.

3.2 Seasonal variations of NH$_3$ in high concentration zones

Eight sites having an AAN higher than 3 µg m$^{-3}$ were distributed within intensive livestock zones (zone A in Fig. 1b). The time series of weekly/biweekly samples are shown in Fig. 2. A few samples having spikes of $>$10 µg m$^{-3}$ were observed at seven of the eight sites and two spikes of 7.7–8.3 µg m$^{-3}$ were observed at the remaining site of SMY. Based on the on-site record, approximately 50% of these spikes ($>$10 µg m$^{-3}$) can be explained by manure application. Manure application on soil was previously reported to rapidly release NH$_3$ to the atmosphere in a few days, leading to a significant N loss (Sommer and Hutchings, 2001; Misselbrook et al., 2005). During the sampling periods of another 10% of these peaks $>$10 µg m$^{-3}$, either smoke or grass cutting was recorded, but no manure application was recorded. For the remaining 40% of the spikes, the on-site record showed no manure application, no smoke, and no grass cutting. It is speculated that these unexplained spikes might be due to N fertilizer application, although the on-site record did not include the term. Like manure application on soil, N fertilizer application on soil can also cause 10–30% N loss by releasing NH$_3$ to ambient air in the initial 3–5 days after the application (Salazar et al., 2012). When the concentration of NH$_3$ in each sample was averaged over the eight sites, there was a decreasing trend from...
earl June to mid-July (Fig. 2d). The decrease was likely a decrease in emissions after early June across southern Ontario (Lillyman et al., 2009). The peaks of NH$_3$ in late July and November were probably due to occasional manure or fertilizer application.

A sharp decrease, e.g., ~50%, in NH$_3$ concentration was observed starting around mid-December. The decrease could be caused by one or more factors as follows. Emissions from livestock production in southern Ontario were estimated to decrease by ~80% in November and December compared to those in October (Lillyman et al., 2009). This does not necessarily mean that the livestock production was also reduced by 80%. Instead, the reduced emissions were due to a combination of reduced livestock production, different treatment of animal wastes and different environmental conditions. Fertilizer application in December, January and February was also negligible (Lillyman et al., 2009). Besides, the snow season in southern Ontario generally starts in November and the highest snowfall occurs in December, January and February (http://www.wunderground.com/global/stations/71265.html, Fig. S5). Thick snow cover should reduce soil emissions substantially. Daily average temperature generally decreases to below the freezing mark in December and the surface soil $T$ (in the top 1–2 cm depth) usually has a rapid response to the change of ambient $T$ in a few hours (Marshall and Holmes, 1988). Low soil $T$ lowers microbial activity and thus soil emissions of NH$_3$ (Misselbrook et al., 2001; Sutton et al., 2009). Moreover, the lower ambient $T$ could favor formation of NH$_3$NO$_3$ in the atmosphere, which has been observed at a few places of southern Ontario (Yao et al., 2011; Zhang et al., 2008), and thus led the depletion of NH$_3$ gas.

However, a biweekly peak of NH$_3$ concentration was observed in the early of February 2007 at MVT (8.8 ppb), ARK (6.4 ppb) and SWB (6.2 ppb) when the lowest ambient $T$ occurred. The peak concentration of NH$_3$ at each site was 1–2 times higher than the site’s average value during the period from December 2006 to March 2007. This phenomenon was not detected at the other five sites. The peaks in the coldest weeks were also observed at some sites belonging to the other two concentration categories and are discussed in detail in Sect. 4.

3.3 Seasonal variations of NH$_3$ in medium concentration zones

Thirty-six sites had an AAN in the range of 1 to 3 µg m$^{-3}$ and are further classified into two sub-Categories here, i.e., Category 2A for sites having no peaks occurred under the lowest $T$ ($\leq -7 ^{\circ}C$) and Category 2B for sites having peaks under the lowest $T$. In the cold seasons, weather conditions in the southern Ontario were generally controlled by synoptic system (Bein et al., 2008). Thus, temporal trends of ambient temperature were all similar at the 74 sites (Figs. 2–4). Apparently, ambient weather conditions alone cannot explain the difference between Categories 2A and 2B.

Category 2A includes 22 sites. The concentration of NH$_3$ in each weekly/biweekly sample was averaged over 22 sites and the time series of the spatially averaged value was obtained for analysis. The simple approach can make the common trends among sites more obvious. The disadvantage of this approach is that it smoothes the distinctive trends associated with specific sites. The spatially averaged value of NH$_3$ over the 22 sites in each weekly/biweekly sample fluctuated at 1.8 ± 0.9 µg m$^{-3}$ (mean ± standard deviation, Fig. 3a) during the period from 1 June to 7 November 2006. The spatially averaged value decreased by approximately 20% and fluctuated at 1.4 ± 0.7 µg m$^{-3}$ during the period from 7 November to 19 December 2006. The spatially averaged value sharply decreased down to 0.8 ± 0.4 µg m$^{-3}$ in mid-December 2006 and then slowly increased and reached 1.3 ± 0.7 µg m$^{-3}$ at the end of March 2007. Similar to those in the high concentration zones, the elevated concentrations before or during early June were only observed at some agricultural-zone sites, such as at West Lorne (WLO) and Egbert (EGC), and were absent at other agricultural-zone sites, such as Caistorville (CSV), Welland (WEL) and Aylmer (AYL) (Fig. 3b, c). This indicated that not all agricultural-zone sites had enhanced emissions of NH$_3$ before or during early June, although the period was the traditional manure and fertilizer application season. This is likely due to the different land use since different crops require different fertilization periods (Fig. S3).

Category 2B includes 14 sites (13 sites in Fig. 4a and one site in Fig. S6). The spatially averaged value of NH$_3$ over 13 sites in each weekly/biweekly sample of Category 2B was approximately 10% higher than that in Category 2A. The seasonal trend of the spatially averaged value in Category 2B was generally similar to that in Category 2A, except that there was a peak of NH$_3$ concentration at low ambient $T$ ($\leq -7 ^{\circ}C$) in Category 2B. Again, an increase of NH$_3$ concentration was observed before or during early June at several agricultural-zone sites, such as Inwood (INW), Comber (COM) and Brigden (BDG) (Fig. 4b), but was absent at other sites, such as Essex (ESS), Teeterville (TTV) and Longwoods (LON). Moreover, one agriculture-zone site (Clear Creek, CLC) had much high concentrations of NH$_3$ in the weekly samples collected between 13–20 June (19.3 µg m$^{-3}$) and 20–27 June (13.0 µg m$^{-3}$) 2006 (Fig. S6). The peak concentrations were 7–10 times higher than the AAN of 1.7 µg m$^{-3}$ and were probably due to fertilizer application (no manure application was recorded during these two weeks).

3.4 Seasonal variations of NH$_3$ in low concentration zones

Thirty sites had an AAN < 1 µg m$^{-3}$. The spatially averaged value of NH$_3$ over the 30 sites in each weekly/biweekly
Fig. 4. Temporal variations of NH$_3$ in 2006–2007 at sites located in the medium concentration zones but with spikes during the coldest temperature periods: (a) the spatially averaged value over 13 sites, (b) three sites having spikes before the early of June, and (c) three sites having no spikes before the early of June.

Fig. 5. Temporal variations of NH$_3$ in 2006–2007 at sites located in the low concentration zones: (a) the spatially averaged value over 30 sites, (b) three sites having data before June, and (c) three sites having no data before June.

Fig. 6. Correlation of NH$_3$ between sites TEV and DDK.

However, the enhanced emissions of NH$_3$ from agricultural and non-agricultural activity before or during early June cannot explain the broad peak during the period from 13 June to 1 August 2006 in Category 3 (Fig. 5a). The lifetime of atmospheric ammonia is relatively short (e.g., a few days, Wen et al., 2013). It was unlikely to take such a long time (e.g., more than a week) to transport NH$_3$ from higher concentration zones to these low concentration sites with known wind conditions, especially considering that dominant wind was from the west and the southwest (Fig. 1b). As discussed later, negative correlations were obtained for NH$_3$ concentration measured at tens of sites between southeastern Ontario and DDK.

Sample fluctuated around 0.4 ± 0.3 µg m$^{-3}$ during the period from August 2006 to March 2007, except during the period from 13 June to 1 August 2006 (Fig. 5a) when the spatially averaged value increased by 50–100 % and reached a maximum of 0.8 µg m$^{-3}$ during mid-July 2006. Twenty out of the 30 sites were situated at the remote non-agriculture zone where no manure and fertilizer application occurred. The temperature effect cannot explain the high concentration levels during the six-week period because the highest ambient temperature occurred in the early of August (Fig. 5a, b). Regional transport of NH$_3$ from the high concentration zones may have led to the higher concentration of NH$_3$ in the low concentration zones. However, the broad peak of NH$_3$ in the Category 3 had a few weeks delay when compared to the peaks observed at the sites in high concentration zones in southwestern Ontario (Figs. 2–4), the latter cases being consistent with seasonal variations of NH$_3$ emissions associated with agricultural activities. For example, as reported by Lillyman et al. (2009), the agricultural emissions of NH$_3$ in June or in July across southern Ontario were only approximately one third of those in May. Zbieranowski and Aherne (2012, 2013) also pointed out that the highest concentration of NH$_3$ across agricultural zones of southern Ontario occurred in May. As presented above, a few peaks of NH$_3$ in the high concentration zones indeed occurred before or during early June 2006.
and southwestern Ontario. Various possible transport mechanisms are discussed in Sect. 4.

Again, at three sites (RBY, Vansickle (VAL) and GAL), a peak of NH$_3$ was observed in January–February 2007 when $T$ was as low as $-10^\circ$C (Fig. 7c). The peak concentration of NH$_3$ observed at each site was 2–3 times higher than the site’s average value during the period from December 2006 to March 2007.

4 Discussion

4.1 Outstanding issues and proposed explanations

The geographical and seasonal patterns of NH$_3$ concentrations presented above can largely be explained by known emission sources and meteorological factors. There are, however, some phenomena that need more explanations. These include (1) peak concentrations at remote non-agricultural sites appeared a few weeks later than at the majority of agricultural sites; (2) NH$_3$ concentrations were higher in winter than in summer at some remote non-agricultural sites; and (3) the sharp increase in NH$_3$ concentration at 20 out of the 74 sites during the coldest two weeks when ambient temperature was lower than $-7^\circ$C.

Differences in local meteorology, for example, the frequency of stable atmospheric condition which typically causes concentration accumulation, might play a small role, however, not enough to explain the phenomena (1) and (2) listed above. Local emissions, which were assumed to be the highest in summer, cannot explain the even higher observed NH$_3$ concentrations in winter at remote sites. Manure dumping in mid-winter by farmers, while forbidden by the law, could be a reason for phenomena (3) listed above. This would explain peaks of NH$_3$ concentrations in the agricultural zones, but cannot explain the peaks in the non-agricultural zones which were several hundred kilometers away from the potential sources.

One possibility causing phenomena (1) and (2) above could be due to the transport mechanisms that are unique for NH$_3$ than for other pollutants. Here we proposed two transport mechanisms based on existing knowledge. The first one is the indirect transport of NH$_3$. Air–soil and air–plant exchanges of NH$_3$ are bi-directional, especially in agricultural areas (Walker et al., 2013). Although NH$_3$ concentration decreased rapidly with distance from sources, the deposited NH$_3$ could be emitted back into the atmosphere and underwent further transportation. Thus, NH$_3$ emitted from intensive agricultural zones travels a short distance every day and takes much longer time to reach remote areas several hundred kilometers away. The second one is the transport of semi-volatile ammonium salts which were formed in NH$_3$-rich area and evaporated in low NH$_3$ area. An earlier study showed that NH$_3$ was enough to neutralize both sulfate and nitrate at agricultural areas, but not enough to neutralize nitrate at remote areas (Zhang et al., 2008). This provides some evidences of the proposed mechanism. These proposed transport mechanisms can best be examined using high spatial resolution concentration measurements with a reasonable time resolution combined with high-resolution model simulation, which unfortunately cannot be accommodated in the present study. However, we managed to look more closely at our data and found some evidences that seem to agree, if not to verify, the above proposed mechanisms, as discussed below.

4.2 Regional transport of NH$_3$ to low concentration zones at $T > 0^\circ$C

Direct transport of NH$_3$ was likely unimportant from the agricultural areas to the remote areas. To estimate the maximum contributions of direct atmospheric transport between these two regions, we chose two sites that had a good correlation in NH$_3$ concentration, but were much closer with each other than the distance between the two regions. The relationship between the site TEV, which had the highest AAN (6.1 µg m$^{-3}$) among all of the 74 sites, and the site DDK, which was 30 km downwind of TEV (noting that the dominant wind was from the west and the southwest) and had an AAN of 1.7 µg m$^{-3}$, was investigated as an example. A moderately good correlation, i.e., $[\text{NH}_3]_{\text{DDK}} = 0.36 \times [\text{NH}_3]_{\text{TEV}} + 0.13$ with $R^2 = 0.58$ ($P < 0.01$), was obtained for NH$_3$ concentration between the two sites when six NH$_3$ spikes $> 9 \mu m^{-3}$ at TEV were excluded (Fig. 6). The good correlation was not necessarily a result of direct transport. Similar meteorological conditions coupled with similar land use and soil nitrogen content could also result in a good correlation if local emissions were the main contributor. Theoretical modeling results by Theobald et al. (2012) showed a two-order of magnitude decrease in NH$_3$ concentration in just 1000 m distance from agriculture sources. The direct transport to 30 km downwind could be much smaller. During five of the six spikes at TEV, the weekly concentrations of NH$_3$ at TEV were approximately one order of magnitude higher than those at DDK during the same period. This suggested that at most 10% of NH$_3$ at TEV can be transported to DDK. Thus, the slope of 0.36 was overwhelmingly caused by similar meteorological conditions. Moreover, as discussed above, the concentrations of NH$_3$ at most of the agriculture sites were higher than 1 µg m$^{-3}$. It is thus reasonable to assume that the local emission of NH$_3$ at DDK contributed 1 µg m$^{-3}$ to the total of 1.7 µg m$^{-3}$ (AAN). This suggests that NH$_3$ from atmospheric transport only contributed 0.7 µg m$^{-3}$ to the AAN at DDK, which was only one tenth of the AAN at TEV.

However, when the concentrations of NH$_3$ in each pair of 74 sites were correlated to each other, a few negative correlations were obtained between southwestern sites and southeastern sites. Table 1 and Fig. S7 showed these pairs in which negative correlations were obtained with $R^2 > 0.4$ and $P < 0.01$. 

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Fig. 7. Comparison of time series of NH$_3$ concentration between southwestern and southeastern sites (the straight lines are trend lines artificially drawn).

Fig. 8. (a) distribution of sites with elevated NH$_3$ in the coldest weeks, and (b) correlations between peak (or net increase) concentrations and annual averages for these sites.

The dominant wind from the west and the southwest likely played an important role on regional transport of NH$_3$ from southwestern Ontario to southeastern Ontario (Wen et al., 2013). In addition, land-lake breeze (land wind is from the south in daytime while lake wind is from the north at night) might have influenced the regional transport of NH$_3$ to the southeastern sites adjacent to Lake Ontario. Regional transport signal of NH$_3$ at the southeastern sites adjacent to Lake Ontario could be different from that at sites dozens or hundreds kilometers north of Lake Ontario. To explore the role of regional transport in the negative correlations of NH$_3$ identified between southwestern and southeastern sites, NH$_3$ concentration at four selected sites in southeastern Ontario were analyzed in detail. The four selected sites, Whitney (WTY) → Hartsmer (HTS) → Fernleigh (FNL) → Crow Lake (CWL), follow a northwest to southeast transect, and are all located in remote non-agricultural zones. Considering that there was no agricultural activity in the remote non-agricultural zones, ammonia collected there should be from a combination of natural emissions and regional transport. Natural emissions of NH$_3$ were found to increase exponentially with air temperature (Zhang et al., 2010; Flechard et al., 2013). It is thus reasonable to assume that the observed summer NH$_3$ level represents the maximum contributions from natural emissions.
Table 1. Negative correlation between each pair of the sites.

<table>
<thead>
<tr>
<th>Southwestern site</th>
<th>Southeastern site</th>
<th>$R^2$ ($P &lt; 0.01$)</th>
</tr>
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<tr>
<td>EGC</td>
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<td>TEV</td>
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At WTY (a remote non-agriculture site), NH$_3$ concentrations were negatively correlated to those at Tobermory (TOB) and at EGC (Table 1 and Fig. 7a). NH$_3$ at TOB fluctuated and increased during the period from 6 June to 29 August 2006 and then fluctuated and decreased till the end of December 2006. The temporal trend matched that of $T$ well ($R^2 = 0.44$ with $P < 0.01$). Higher ambient $T$ usually disfavors formation of ammonium nitrate which is the sink of NH$_3$ gas (Seinfeld and Pandis, 2006). In contrast, higher soil $T$ favors decomposition of organics, but inhibits the nitrification process (Grunditz and Dalhammar, 2001; Pagans et al., 2006). Thus, higher soil $T$ usually results in higher emissions of NH$_3$ (Lillyman et al., 2009). The temporal variation of NH$_3$ at EGC was generally similar to that at TOB. However, during the period from 6 June to 29 August 2006, NH$_3$ at WTY fluctuated and decreased and reached a minimum of 0.2 µg m$^{-3}$ in August, which probably represented the maximum contribution from local emissions at this site because $T$ was the highest during this period. Higher concentrations of NH$_4$ at WTY in early June were probably from external transport of NH$_3$ associated with enhanced agricultural emissions in May. With the external transport signal decreasing from June to August, the concentrations of NH$_3$ at WTY decreased accordingly. From the end of September to the end of October, NH$_3$ at WTY fluctuated and increased while $T$ decreased. The increase of NH$_3$ in concentration probably reflected the enhancement of new external transport signal.

At HTS (a remote non-agriculture site), concentrations of NH$_3$ were negatively correlated with that at AYL. The concentration of NH$_3$ at HTS increased from < 0.1 µg m$^{-3}$ in early June to the maximum value of 0.6 µg m$^{-3}$ at the end of June (Fig. 7b). From early July to the end of July, the concentration of NH$_3$ at HTS decreased to ~ 0.1 µg m$^{-3}$ and remained at 0.1–0.2 µg m$^{-3}$ in August and September. The 0.2 µg m$^{-3}$ should represent the maximum contribution from local emissions at this site because $T$ was the highest in the early of August. The increase of NH$_3$ at HTS in June 2006 was probably due to transport of NH$_3$ from the southwestern agriculture zones with higher emissions of NH$_3$. The maximum value at HTS had more than two weeks delay when compared to the spikes of NH$_3$ observed in the southwestern Ontario (Figs. 2–4). The proposed transport mechanisms might have played a role, and the first mechanism is likely more important than the second one during hot seasons. The decrease of NH$_3$ at HTS from early July 2006 was probably due to a decreased external transport signal. In early October, NH$_3$ rapidly increased and then fluctuated at a higher value, which should attributed to a new external transport signal arriving at HTS.

At FNL (a remote non-agriculture site), NH$_3$ concentrations were negatively correlated with those at Kincardine (KIN) and Zurich (ZUR) (Table 1 and Fig. 7c). A broad peak of NH$_3$ at FNL occurred during the period from the mid-June to the end of July. NH$_3$ rapidly decreased to ~ 0.2 µg m$^{-3}$ in early August and stayed low through the rest of August. Similar to discussions for the above two sites, 0.2 µg m$^{-3}$ was proposed to be the maximum contribution from local emissions at FNL. The broad peak between June and July was also likely related to regional transport. The delayed peak at FNL compared to those at KIN and ZUR seems to agree with the first proposed mechanisms above. NH$_3$ gradually increased from early September to mid-October, suggesting a new transport signal appeared. After mid-October, NH$_3$ fluctuated at higher concentration levels. The reverse temporal trend was generally observed at KIN and ZUR, leading to their concentrations of NH$_3$ being negatively correlated to that at FNL.

At CWL (a remote non-agriculture site), NH$_3$ concentrations were negatively correlated with those at COM and WLO (Fig. 7d). The concentrations of NH$_3$ at CWL increased from 0.2 µg m$^{-3}$ in early June to the maximum concentration of 1.4 µg m$^{-3}$ during 11–18 July and then decreased. The reverse trend was observed at COM and WLO. The concentration peak of NH$_3$ at CWL in mid-July was also likely associated with regional transport of NH$_3$. The delayed peak seems to agree with the first proposed mechanisms above. This is because the peak has over one month delay, compared to high agricultural emissions of NH$_3$ in May and the early of June at sites of southwestern Ontario. The concentrations of NH$_3$ at CWL reached a lower value (~ 0.2 µg m$^{-3}$) in September and then fluctuated between September and November. 0.2 µg m$^{-3}$ could also represent...
the maximum contribution from local sources to the observed NH$_3$ concentration. The concentrations of NH$_3$ at CWL gradually increased after the end of November, suggesting the arrival of a new transport signal.

### 4.3 Regional transport of NH$_3$ to low concentration zones at $T < 0 \degree C$

As shown in Figs. 5 and 7, no decreases of NH$_3$ concentrations were observed at these low concentration sites during the period from December 2006 through March 2007. At a few sites, the concentrations of NH$_3$ in the winter even increased to some extent and were higher than their respective concentrations in June and July 2006. For example, NH$_3$ at WTY and FNL gradually increased from the lowest concentration at the end of August to a higher value in mid-October and then fluctuated at the higher value through the winter. The average value of NH$_3$ (0.5 ± 0.2 µg m$^{-3}$) at WTY during the period from the mid-October 2006 through March 2007 was significantly higher than the average (0.4 ± 0.1 µg m$^{-3}$) during the period from June through July 2006 with 95% confidence. The same phenomenon was also observed at FNL.

The maximum contribution from local emissions to NH$_3$ measured at WTY and FNL was likely ≤ 0.2 µg m$^{-3}$ as discussed in Sect. 4.2. The proposed transport mechanisms should also apply in cold season, which might have increased NH$_3$ at WTY and FNL. Due to the very low $T$ and thus low natural and re-emissions of NH$_3$ along the transport paths, the first mechanism should play a much smaller role in cold than in hot season. Thus, the second proposed transport mechanism should be more important. In the cold seasons, NH$_4$NO$_3$ can be formed in the polluted environments of the southern Ontario (Zhang et al., 2008, Yao et al., 2011, Yao and Zhang, 2012). The chemically formed NH$_4$NO$_3$ in the atmosphere was mainly present in the particles < 2.5 µm and can be transported over dozens to hundreds of kilometers, depending on ambient conditions. When the NH$_4$NO$_3$ arrived at receptor sites with low concentrations of NH$_3$, the product of [HNO$_3$]$^+$[NH$_3$] was less than the equilibrium constant of NH$_4$NO$_3$ and resulted in the release of NH$_3$ gas from the particle phase. In the warm seasons from June to August, the chemically formed NH$_4$NO$_3$ in the atmosphere was often thermodynamically disfavored in southern Ontario (Zhang et al., 2008) and ammonium was mainly associated with sulfate in atmospheric particles.

### 4.4 Analysis of elevated NH$_3$ concentration at $T < -7 \degree C$

As mentioned above, the elevated concentrations were observed at 20 out of the 74 total sites when the biweekly averaged ambient $T$ sharply decreased to below $-7 \degree C$. In southwestern Ontario, these sites were situated either at Algonquin Arch or evenly distributed at the two sides of the Arch (Fig. 8a). Note that Algonquin Arch was formed in Paleozoic time (Johnson et al., 1992). However, the sites in southeastern Ontario were located on a much narrow line, which almost overlapped the demarcation line between non-agriculture and agriculture zones in this region.

Among these sites, a good correlation existed between the elevated concentration and associated AAN; so was the case between the elevated concentration and the net increase of NH$_3$ (calculated as the difference between the elevated concentration and the minimum concentration in January and February) (Fig. 8b). At the southwestern sites, the good correlation implies that the elevated NH$_3$ concentration were mainly due to local emissions. However, agricultural activity was limited in January and February. Emissions from accident outdoor agricultural activity at the lowest ambient $T$ were unlikely. Indoor agricultural activity should neither be the main cause for the elevated concentration. This is because indoor activity occurred throughout the whole winter and should not cause elevated concentration only in the coldest two weeks and only at specific sites. In southeastern Ontario, the elevated NH$_3$ supposed to be related to regional transport of NH$_3$. However, regional transport should cause an increase of NH$_3$ in a broad band zone downwind of source zones instead of only on a narrow line along ~300 km length in the middle of agricultural and non-agricultural areas. Recently, Yao et al. (2013) found that traffic emissions of NH$_3$ yielded a negligible contribution to atmospheric ammonia in southern Ontario. Thus, the elevated NH$_3$ under extremely low $T$ at sites in both agricultural and non-agricultural regions cannot be explained by existing knowledge.

Satellite data during that period showed that the percentage of snow-covered soil across the southern Ontario was approximately 100%, except over lakes, rivers and traffic roads, etc. (Fig. S5). The emission of NH$_3$ from snow-covered soil under harsh low-temperature conditions were traditionally considered to be very low (Sommer and Hutchings, 2001; Sutton et al., 2009; Sheppard et al., 2010). However, a few studies had demonstrated that microbial processes continued in snow-covered soils, and a significant portion of decomposition and production of microbiologically derived trace gases can occur in snow-covered soils (Groffman et al., 2001; Nemergut et al., 2005; Edwards et al., 2006). This is because snow can insulate soil and protect the microbial community from freezing ambient conditions (Nemergut et al., 2005). These previous studies and observations in this study suggested that emissions of NH$_3$ from snow-covered soil need to be re-examined.

### 5 Missing emission sources of NH$_3$ in two agricultural zones

When the geographical distribution of NH$_3$ concentration was compared with that of NH$_3$ emission inventory recently developed by Environment Canada, inconsistency was found
at two zones (B and D) as outlined in Fig. 1b. Three sites (TOB, WIA and MEA) between Lake Huron and Georgian Bay (circled B zone in Fig. 1b) were approximately 100–150 km distance from the high NH₃ emission zone (circled A zone in Fig. 1b). As shown in Fig. 1c, AAN at the three sites, which was ∼ 1.1 µg m⁻³, was situated above the regression line between AAN vs. the emission density, suggesting the potential underestimation of NH₃ emission in B zone. Direct regional transport of NH₃ from A zone to B zone was likely unimportant because of the 100–150 km distance between the two zones. In fact, AAN at MOO, a site between A zone and B zone and approximately 30 km from A zone, was only 1.0 µg m⁻³. NH₃ transported from A zone to B zone should be diluted to a much higher extent than that from A zone to the site of MOO due to further distance to B zone. Thus, NH₃ emission from B zone was thought to be underestimated by a factor of 1 to 3. Another zone which NH₃ emission was also potentially underestimated was D zone. As shown in Fig. 1c, AAN at six sites was also situated above the regression line between AAN vs. the emission density. NH₃ emission from D zone seemed to be underestimated by 10–50%.

Ammonia in soils can be constantly formed from biological degradation of organic compounds (Sommer and Hutchings, 2001; Beusen et al., 2008). In agricultural zones, external N input to soil mainly includes manure and N fertilizer application, atmospheric deposition of NOₓ and biogenic N fixation. Atmospheric deposition of NOₓ and biogenic N fixation could release NH₃ gas through soil N cycling processes, but were not considered in the emission inventory of NH₃. Atmospheric deposition of NOₓ might be a negligible source to N budget in soil at B zone because of low NOₓ concentration (Fig. S4). Globally, the estimated N fixed biologically in agricultural systems was 50–70 Tg N per year. Soybean was the dominant crop legume and fixed 16.4 Tg N globally and 5.7 Tg N in the US (Galloway et al., 2008; Herdige et al., 2008). Biogenic N fixation in agriculture systems followed by soil N cycling in B agriculture zones should be investigated in the future in order to identify missing sources.

D zone is located immediately downwind of the zone with intensive anthropogenic emission of NOₓ (Fig. S4). High atmospheric deposition of NO_xy was expected in D zone. The deposited NOₓ can infiltrate to sub-surface soil and convert into NH₃ through soil N cycling processes. Atmospheric deposition of NOₓ could be one of the important sources for local emissions of NH₃ in D zone and might have contributed to the underestimation of NH₃ emissions. Moreover, N fixed biologically by agriculture systems could also be a potential source at D zone.

6 Conclusions

High spatial-resolution weekly/biweekly ambient NH₃ concentration data measured mainly from June 2006 to March 2007 at 74 sites across southern Ontario were analyzed in detail. In the remote and rural non-agricultural zones, the AAN was less than 1 µg m⁻³ and the same was true for a few sites in low intensity agricultural zones. Approximately 50% of the sites had an AAN in the range of 1–3 µg m⁻³ and exceeded the new critical level value. There were eight sites situated inside the intensive livestock production zones having an AAN in the range of 3.6 to 6.1 µg m⁻³. These high NH₃ concentration levels could have a potential damage to vegetation.

At remote non-agricultural sites with low NH₃ concentrations, local emissions were estimated to contribute a maximum of 0.2 µg m⁻³ to the overall measured concentration, the rest was likely from regional transport which was the dominant contributor during most of the time. Two transport mechanisms were proposed, but needed further validation using high temporal resolution field data and/or high spatial and temporal resolution air quality model simulations. The first mechanism is a combination of bi-directional air–soil and air–plant exchanges of NH₃ and transport of air masses containing the so-produced NH₃ along the trajectories, and the second mechanism is transport of NH₄NO₃ produced in upstream followed by the release of NH₃ from NH₄NO₃ at receptor sites. The first mechanism generally took a few weeks transporting NH₃ from emission sources in southwestern Ontario to receptor sites in southeastern Ontario, which seems to be consistent with the time-lag in peak concentrations between the two regions. Direct transport of NH₄NO₃ should only take a few hours to a few days, but the time of releasing NH₃ from NH₄NO₃ could be much longer depending on the temperature and humidity conditions. It is likely that the first mechanism is more important than the second one in hot seasons while the opposite should be the case in cold seasons. One unexplained phenomenon was the increase of ambient NH₃ level observed in the coldest two weeks at 20 out of 74 sites. The 20 sites were distributed along a banana-shaped curve in ∼ 600 km length extending from southwestern to southeastern Ontario. The mechanisms causing the high NH₃ at these sites are not clear and require further studies.

In two agriculture zones, the observed high concentrations of NH₃ were contradictory to the low emissions in the recently developed emission inventory. Emissions of NH₃ at the two zones were likely underestimated. N fixed biologically in agricultural systems and/or atmospheric deposition of NOₓ followed by soil N cycling processes should be considered in future development of NH₃ emission inventory to explain the missing sources.

Supplementary material related to this article is available online at http://www.biogeosciences.net/10/7913/2013/bg-10-7913-2013-supplement.pdf.

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Acknowledgements. We greatly appreciate Q. Zheng for proving emission inventory data, J. Narayan for providing land use map, and all the people who contributed to the field data collection and data quality control.

Edited by: X. Wang

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