Export of $^{134}$Cs and $^{137}$Cs in the Fukushima river systems at heavy rains by Typhoon Roke in September 2011

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Abstract. At stations on the Natsui River and the Same River in Fukushima Prefecture, Japan, effects of a heavy rain event on radiocesium export were studied after Typhoon Roke during 21–22 September 2011, six months after the Fukushima Dai-ichi Nuclear Power Plant accident. Radioactivity of $^{134}$Cs and $^{137}$Cs in river waters was 0.009–0.098 Bq L$^{-1}$ in normal flow conditions during July–September 2011, but it increased to 0.85 Bq L$^{-1}$ in high flow conditions because of heavy rains occurring with the typhoon. The particulate fractions of $^{134}$Cs and $^{137}$Cs were 21–56 % of total radiocesium in the normal flow condition, but were close to 100 % after the typhoon. These results indicate that the pulse input of radiocesium associated with suspended particles from land to coastal ocean occurred because of the heavy rain event. Export flux of $^{134}$Cs and $^{137}$Cs attributable to the heavy rain accounts for 30–50 % of the annual radiocesium flux from inland to coastal ocean region in 2011. Results show that rain events are one factor contributing to the transport and dispersion of radiocesium in river watersheds and coastal marine environments.

1 Introduction

A nuclear accident at the Fukushima Dai-ichi Nuclear Power Plant (NPP) occurred after the 2011 Tōhoku earthquake and tsunami. About 15 PBq of both $^{134}$Cs and $^{137}$Cs was released from the NPP as a result of venting operations and hydrogen explosions (Japanese Government, 2011; Chino et al., 2011). Surface deposition of $^{134}$Cs and $^{137}$Cs shows considerable external radioactivity in a zone extending northwest from the NPP, about 20 km wide and 50–70 km long inside the 80 km zone of the NPP (MEXT, 2011; Yoshida and Takahashi, 2012). Moderate radioactivity (100–600 kBq m$^{-2}$) was also detected in the Naka-dori region. The deposition pattern is explained by emission rates of $^{134}$Cs and $^{137}$Cs coupled with wind direction and precipitation (Morino et al., 2011).

A major part of radiocesium deposited on the ground surface is present at the surface to 5 cm depth (MEXT, 2012a; Koarashi et al., 2012). Chemical extraction of $^{134}$Cs and $^{137}$Cs from selected soil samples has revealed that both radionuclides in the soil are only slightly water-soluble. Even the fraction extracted with 1 M ammonium acetate was only approximately 10 % (Matsunaga et al., 2013). However, $^{134}$Cs and $^{137}$Cs have been transported from contaminated watersheds to rivers in Fukushima Prefecture since the Fukushima Dai-ichi NPP accident (MEXT, 2012a; Sakaguchi et al., 2012). Similar outcomes were observed for the Pripyat River and Dnieper River in Ukraine after the Chernobyl accident in 1986 (IAEA, 2006a). The migration of $^{137}$Cs has decreased markedly over time in river waters from Ukraine (UHMI, 2004; IAEA, 2006b) and Finland (Saxén and Ilus, 2001). The radioactivity of $^{137}$Cs shows little change from upstream to downstream of the exclusion zone in the Pripyat River of the Chernobyl area (IAEA, 2006a). An increase in radioactivity of $^{134}$Cs and $^{137}$Cs in...
river waters was also found in the Chernobyl area during a spring flood event (IAEA, 2006a, b) and in northwestern Italy from a delayed release in summer during ice and snow melting in mountain areas (Spezzano et al., 1994). The transport of materials generally depends on watershed conditions such as vegetation, slope, soil types, and spring snow-melting. Clarifying the migration behavior of radiocesium and its controlling factors is important for future prediction of its dispersion in Fukushima Prefecture, Japan.

To elucidate the short-term to long-term impacts of the Fukushima Dai-ichi NPP accident on riverine and coastal marine ecosystems, the Japanese Government has been monitoring the radioactivity of $^{134}$Cs and $^{137}$Cs in river systems in Fukushima Prefecture (MEXT, 2012b). Japanese rivers have short lengths, high riverbed slopes, and high river regime coefficients (ratio of maximum/minimum discharge; Suetsugu, 2005). Annual mean precipitation is generally high (e.g., 1718 mm during 1971–2000) because of the rainy season in Japan, typhoons, and snow-melting events in spring (MLIT, 2012). Matsunaga et al. (1991) reported that the radioactivity of $^{137}$Cs derived from fallout increased in the Kuji River during high flow conditions caused by rain events. Nagano et al. (2003) pointed out that variations of suspended form and dissolved form concentrations of elements in the Kuji River waters occurred as a function of the water discharge rates. It is important that we evaluate rain event effects on radiocesium export from land to ocean in the Fukushima area.

This study was conducted to investigate the transport of $^{134}$Cs and $^{137}$Cs in river systems in Fukushima Prefecture after rain events. This report describes monitoring results of radioactivity of $^{134}$Cs and $^{137}$Cs in river waters at two rivers after a heavy rain event, a typhoon, in September 2011. Field experiments were conducted at the Natsui River and the Same River in the southern part of Fukushima Prefecture, Japan. We examined the transport behavior of radiocesium and estimated its export flux from inland to coastal areas.

2 Materials and methods

Typhoon Roke (T1115) struck Japan on 21 September and subsequently weakened to an extra-strong tropical cyclone on 22 September 2011 (JMA, 2011). The typhoon precipitated more than 400 mm of rain daily in parts of eastern and western Japan (JMA, 2011). Fukushima Prefecture recorded rainfall of 100–200 mm during 15–22 September. The daily rainfall on 21 September was 137 mm because of the impact of Typhoon Roke at Onahama of Iwaki city, located in the southern coastal region of Fukushima and in a watershed of the Natsui River. This value is about one-tenth of the annual mean rainfall (1409 mm for 1981–2010; JMA, 2012).

This study investigated Natsui River and Same River, each flowing to the Pacific coast through less-contaminated areas (<100 kBq m$^{-2}$; MEXT, 2011). The Natsui River watershed area is 749 km$^2$. That of the Same River is 600 km$^2$. The Natsui River length is 67 km. That of the Same River is 58 km. The annual mean water discharge data in 2011 were 17.6 m$^3$ s$^{-1}$ for the Natsui River and 21.4 m$^3$ s$^{-1}$ for the Same River (Fukushima Prefectural Government, 2012). The water discharge data are presented in Fig. 2. River water samples (10–20 L) were collected at normal flow conditions on 12 and 27 July, 13 September, 24 November, 6 December, and 22 September in high-flow conditions after Typhoon Roke in 2011. Sampling was conducted at the Iwaki-bashi bridge over the Natsui River and the Eguri-Ohashi bridge over the Same River.

In river waters after the heavy rain event with the typhoon, particles were separated using centrifugation and filtration with No. 5A (approx. pore size of 7 µm) filters and 0.45 µm pore size membrane filters. In this study, suspended solids using centrifugation are designated as “fraction 1” (FR1). The suspended solids on the filters are designated as “fraction 2” (FR2) for those filtered with No. 5A filters and “fraction 3” (FR3) for those collected with 0.45 µm filters.

The radioactivity of $^{134}$Cs and $^{137}$Cs in the filtered river waters was measured as dissolved forms of radiocesium with gamma-ray spectrometry using ammonium
molybdophosphate (AMP)/Cs compound method (Tanaka et al., 2006). The $^{134}$Cs and $^{137}$Cs were measured using gamma-ray spectrometry with a low background Ge detector at the Low Level Radioactivity Laboratory and the Ogoya Underground Laboratory of Kanazawa University for 1–3 days (Hamajima and Komura, 2010). The gamma lines were used for the activity calculation at 605 keV and 795 keV for $^{134}$Cs and 661 keV for $^{137}$Cs. The cascade summing effect was corrected for $^{134}$Cs using a Fukushima contaminated soil sample. The decay correction of radioactive concentration for $^{134}$Cs and $^{137}$Cs was done at each sampling date. The deposited solids and suspended solids on the filters were also measured using gamma-ray spectrometry after drying them at room temperature. The total radioactivity of $^{134}$Cs and $^{137}$Cs is the sum of radioactivity for the suspended solids and filtered river waters.

In normal flow conditions on 12 July and 13 September, dissolved and particulate forms of radiocesium were separated sequentially using cartridge filters with pore sizes of 10 µm, 1 µm, and 0.45 µm. The radioactivity of $^{134}$Cs and $^{137}$Cs was recorded as total (dissolved and particulate forms) for the raw river waters and as dissolved for the filtered waters. The $^{134}$Cs and $^{137}$Cs particulate phases were estimated by subtracting the radioactivity for the dissolved phase from the total radioactivity.

The mineral composition of riverine suspended solids on the filters and of deposited solids after the heavy rain event was analyzed using X-ray diffraction (XRD). The XRD analysis of powdered samples mounted on glass slides was conducted using an Ultima IV (Rigaku Corp.) diffractometer using CuKα radiation operated at 40 kV and 30 mA from 2° to 65°. Suspended solids on GF/F and membrane filters were also mounted on glass slides and were measured for clay mineral composition.

### 3 Results and discussion

#### 3.1 Radioactivity of $^{134}$Cs and $^{137}$Cs

Total radioactivity measurements of $^{134}$Cs and $^{137}$Cs in the river waters are shown in Table 1 and Fig. 3. The respective radioactivity measurements of $^{134}$Cs and $^{137}$Cs were 0.009 Bq L$^{-1}$ to 0.089 Bq L$^{-1}$ and 0.011 Bq L$^{-1}$ to 0.098 Bq L$^{-1}$ in the normal flow condition. The radioactivity decreased concomitantly with increasing time after the Fukushima Dai-ichi NPP accident. However, the radioactivity of $^{134}$Cs and $^{137}$Cs in the high flow condition after the typhoon was about one order higher than that in normal flow conditions, which indicates that high export of $^{134}$Cs and $^{137}$Cs occurred after the heavy rain event. The $^{134}$Cs/$^{137}$Cs activity ratio for all samples is about 1.0, corrected to 11 March 2011. These values are consistent with those of surface soil samples at 0–5 cm depth (Yamamoto et al., 2012; Matsunaga et al., 2013). Therefore, radiocesium derived from the Fukushima Dai-ichi NPP was transported from the deposited surface to rivers. The highest radioactivity is 2–3 orders higher than that collected in 1985–1988 from Japanese rivers (Hirose et al., 1980; Matsunaga et al., 1991).

#### 3.2 Existing forms of $^{134}$Cs and $^{137}$Cs

Figure 4 shows the total radioactivity of $^{137}$Cs versus radioactivity of dissolved forms of $^{137}$Cs. In normal water flow conditions, dissolved forms of $^{137}$Cs were 0.011–0.064 Bq L$^{-1}$, but in the high flow conditions, after the heavy rain event, dissolved $^{137}$Cs was about 0.005 Bq L$^{-1}$. These results indicate that heavy rains affect the export of radiocesium from the Fukushima Dai-ichi NPP accident that had
Table 1. Water quality and radioactivity of $^{134}$Cs and $^{137}$Cs in water samples from the Natsui River and the Same River.

<table>
<thead>
<tr>
<th>Sampling date</th>
<th>pH</th>
<th>WT (°C)</th>
<th>$^{134}$Cs ($\times 10^{-3}$ Bq L$^{-1}$)</th>
<th>$^{137}$Cs ($\times 10^{-3}$ Bq L$^{-1}$)</th>
<th>$^{134}$Cs/$^{137}$Cs$^a$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Natsui River</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>12 Jul 2011</td>
<td>7.6</td>
<td>29.5</td>
<td>49.0 ± 1.3</td>
<td>52.0 ± 2.5</td>
<td>1.06 ± 0.06</td>
</tr>
<tr>
<td>27 Jul 2011</td>
<td>7.3</td>
<td>25.5</td>
<td>89.0 ± 1.6</td>
<td>98.0 ± 2.5</td>
<td>1.02 ± 0.03</td>
</tr>
<tr>
<td>13 Sep 2011</td>
<td>7.4</td>
<td>25.2</td>
<td>25.2 ± 1.2</td>
<td>26.0 ± 1.2</td>
<td>1.14 ± 0.07</td>
</tr>
<tr>
<td>22 Sep 2011</td>
<td>7.5</td>
<td>18.6</td>
<td>673.0 ± 4.0</td>
<td>853.0 ± 4.0</td>
<td>1.00 ± 0.01</td>
</tr>
<tr>
<td>24 Nov 2011</td>
<td>7.6</td>
<td>6.5</td>
<td>61.5 ± 1.8</td>
<td>78.7 ± 1.2</td>
<td>0.97 ± 0.03</td>
</tr>
<tr>
<td>6 Dec 2011</td>
<td>7.4</td>
<td>2.4</td>
<td>22.0 ± 1.4</td>
<td>27.7 ± 1.6</td>
<td>1.04 ± 0.06</td>
</tr>
<tr>
<td>Same River</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>12 Jul 2011</td>
<td>7.7</td>
<td>25.4</td>
<td>74.5 ± 2.4</td>
<td>81.0 ± 2.2</td>
<td>1.03 ± 0.04</td>
</tr>
<tr>
<td>27 Jul 2011</td>
<td>7.4</td>
<td>23.4</td>
<td>47.5 ± 5.9</td>
<td>52.0 ± 5.7</td>
<td>0.99 ± 0.16</td>
</tr>
<tr>
<td>13 Sep 2011</td>
<td>7.5</td>
<td>22.9</td>
<td>15.9 ± 1.3</td>
<td>18.9 ± 1.2</td>
<td>1.00 ± 0.10</td>
</tr>
<tr>
<td>22 Sep 2011</td>
<td>7.7</td>
<td>18.6</td>
<td>360.0 ± 3.0</td>
<td>424.0 ± 3.0</td>
<td>1.01 ± 0.01</td>
</tr>
<tr>
<td>24 Nov 2011</td>
<td>7.3</td>
<td>11.8</td>
<td>13.2 ± 0.7</td>
<td>15.3 ± 0.6</td>
<td>1.00 ± 0.05</td>
</tr>
<tr>
<td>6 Dec 2011</td>
<td>7.5</td>
<td>8.5</td>
<td>8.9 ± 1.0</td>
<td>11.4 ± 1.5</td>
<td>1.10 ± 0.13</td>
</tr>
</tbody>
</table>

$^a$ $^{134}$Cs/$^{137}$Cs activity ratio was calculated after decay correction of $^{134}$Cs and $^{137}$Cs radioactivity on 11 March 2011.

![Fig. 4. Total and dissolved radioactivity of $^{137}$Cs in river waters from the Natsui River (●) and the Same River (○).](image)

been deposited on the ground surface. Similar results were reported for Ukraine river systems after the Chernobyl accident (Matsunaga et al., 1998).

Table 2 presents percentages of particulate phase $^{137}$Cs at normal and high flow conditions in the Natsui River and the Same River after the Fukushima Dai-ichi NPP accident, together with results for the Kuji River before the Fukushima Dai-ichi NPP accident. At normal flow conditions, the percentage of particulate $^{137}$Cs is 21–56 % on average of 40 ± 17 %. The other rivers in central to northern Japan were 11–47 % (Hirose et al., 1990; Matsunaga et al., 1991). The Pripyat River had approximately 40–60 % of radiocesium in the particulate phase during the decade after the Chernobyl accident (Voitsekhovich et al., 1997). The effects of heavy rains reveal that a major part of $^{137}$Cs is present as particulate phase in the Natsui River and the Same River. The particulate form of $^{137}$Cs is predominant in river waters after the rain event in the Kuji River (Matsunaga et al., 1991).

As Table 3 shows, radioactivity of suspended solids is about 2000 Bq kg$^{-1}$-dried suspended solids (ss). This value corresponds to that of river-bottom sediment (about 2000 Bq kg$^{-1}$: Ministry of the Environment, 2012) and soil in watershed (230–2400 Bq kg$^{-1}$: MEXT, 2012b). $^{134}$Cs and $^{137}$Cs of fine particle fraction trapped on the filters in the river waters also show higher radioactivity content of 1548–2336 Bq kg$^{-1}$-ss for the Natsui River and 7161–8649 Bq kg$^{-1}$-ss for the Same River. Radiocesium associated with suspended solids suggests an important pathway from watersheds to rivers.

Figure 5 shows X-ray diffraction analysis for the suspended solids filtered with Advantec no. 5A filters (FR2), membrane filters with pore size of 0.45 µm (FR3), and deposited suspended solids (FR1). Blank GF/F and membrane filters have their own characteristics such as a broad peak between 13 and 35 ° 2θ and three peaks of 13–25 ° 2θ, respectively. All samples include clay minerals such as chlorite, vermiculite, smectite, mica, and kaolin. Clay mineral types in the fixation of radiocesium have been widely claimed (e.g.,
Facchinelli et al., 2001; Korobova et al., 2007). For example, an increase in $^{137}$Cs-specific activity was observed in floodplain soil with increased smectite content in clay fractions (Korobova et al., 2007). Selective sorption of $^{137}$Cs has been reported for illite and mica at laboratory experimental systems (Brouwer et al., 1983; Staunton and Roubaud, 1997). Therefore, clay minerals in suspended solids from the Natsui River and the Same River appear to have fixation and/or association capability for $^{134}$Cs and $^{137}$Cs. However, X-ray diffraction data for size-fractionated samples can not estimate the percentage of each clay mineral because of the small amounts of samples and their crystallinity. Future studies are expected to provide detailed characterization of suspended solids in the Same River and Natsui River.

### 3.3 Migration behavior of $^{134}$Cs and $^{137}$Cs in river systems

Distribution coefficient ($K_d$) between suspended solids and river water is defined as

$$K_d = \frac{C_{\text{solid}}}{C_{\text{dissolved}}},$$  

(1)

where $C_{\text{solid}}$ and $C_{\text{dissolved}}$ respectively denote the $^{137}$Cs concentrations in the suspended solids (Bq g$^{-1}$) and dissolved phase (Bq mL$^{-1}$). The fate and bioavailability depend strongly on the $K_d$ and strength of the particle–contaminant association. Estimation of $K_d$ values was conducted using data from measurements presented in Tables 1 and 3. The $K_d$ is 0.43–0.55 $\times 10^6$ mL g$^{-1}$ for the Natsui River and 4.1–5.0 $\times 10^6$ mL g$^{-1}$ for the Same River. These values are 1–2 orders higher than those of other Japanese rivers such as the Tone River and the Ishikari River (Hirose et al., 1990) and the Kujiriver (Matsunaga et al., 1991) before the Fukushima Daiichi NPP accident. The $K_d$ values of the Fukushima rivers are

<table>
<thead>
<tr>
<th>River</th>
<th>Flow condition</th>
<th>$^{137}$C $\times 10^{-3}$ Bq L$^{-1}$</th>
<th>Particulate $^{137}$Cs (%)</th>
<th>No. sample</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>Natsui$^a$</td>
<td>High</td>
<td>49.0 ± 1.3</td>
<td>~ 100</td>
<td>1</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>Normal</td>
<td>89.5 ± 1.6</td>
<td>40 ± 17</td>
<td>2</td>
<td>This study</td>
</tr>
<tr>
<td>Same$^a$</td>
<td>High</td>
<td>25.2 ± 1.2</td>
<td>99</td>
<td>1</td>
<td>This study</td>
</tr>
<tr>
<td></td>
<td>Normal</td>
<td>673.0 ± 4.0</td>
<td>26 ± 5</td>
<td>2</td>
<td>This study</td>
</tr>
<tr>
<td>Kuji$^b$</td>
<td>High</td>
<td>1.0$^c$</td>
<td>77</td>
<td>1</td>
<td>Matsunaga et al. (1991)</td>
</tr>
<tr>
<td></td>
<td>Normal</td>
<td>0.12 ± 0.06</td>
<td>41 ± 17</td>
<td>12</td>
<td>Matsunaga et al. (1991)</td>
</tr>
<tr>
<td>Kitakami</td>
<td></td>
<td>0.14</td>
<td>39</td>
<td>1</td>
<td>Hirose et al. (1990)</td>
</tr>
<tr>
<td>Tone</td>
<td></td>
<td>0.77 ± 0.59</td>
<td>21 ± 9</td>
<td>6</td>
<td>Hirose et al. (1990)</td>
</tr>
<tr>
<td>Ishikari</td>
<td></td>
<td>0.37 ± 0.12</td>
<td>47 ± 31</td>
<td>2</td>
<td>Hirose et al. (1990)</td>
</tr>
<tr>
<td>Kuzuryu</td>
<td></td>
<td>1.17</td>
<td>11</td>
<td>1</td>
<td>Hirose et al. (1990)</td>
</tr>
</tbody>
</table>

$^a$ Water discharge was 238–350 m$^3$ s$^{-1}$ in a high flow condition and 11.0–23.7 m$^3$ s$^{-1}$ in a normal flow condition.

$^b$ Water discharge was 75.4 m$^3$ s$^{-1}$ in a high flow condition and 8.1–59.4 m$^3$ s$^{-1}$ in a normal flow condition.

$^c$ Precipitation was above 30 mm day$^{-1}$.

<table>
<thead>
<tr>
<th>Sample</th>
<th>Suspended solid (g)</th>
<th>$^{134}$C (Bq kg$^{-1}$-ss)</th>
<th>$^{137}$Cs (Bq kg$^{-1}$-ss)</th>
<th>$^{134}$Cs/$^{137}$Cs$^*$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Natsui</td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>FR1</td>
<td>6.88</td>
<td>1980 ± 11</td>
<td>2336 ± 11</td>
<td>1.00 ± 0.01</td>
</tr>
<tr>
<td>FR2</td>
<td>0.096</td>
<td>1548 ± 58</td>
<td>1817 ± 64</td>
<td>1.03 ± 0.04</td>
</tr>
<tr>
<td>FR3</td>
<td>0.362</td>
<td>1934 ± 70</td>
<td>2303 ± 76</td>
<td>0.97 ± 0.04</td>
</tr>
<tr>
<td>Same</td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>FR1</td>
<td>3.42</td>
<td>1569 ± 14</td>
<td>1865 ± 14</td>
<td>1.01 ± 0.01</td>
</tr>
<tr>
<td>FR2</td>
<td>0.062</td>
<td>7176 ± 113</td>
<td>8268 ± 127</td>
<td>1.03 ± 0.02</td>
</tr>
<tr>
<td>FR3</td>
<td>0.182</td>
<td>7691 ± 81</td>
<td>8649 ± 93</td>
<td>1.03 ± 0.01</td>
</tr>
</tbody>
</table>

ss is suspended solids in river water.

Decay correction of radioactivity for $^{134}$Cs and $^{137}$Cs was done each sampling date.

$^*$ $^{134}$Cs/$^{137}$Cs activity ratio was calculated after decay correction of $^{134}$Cs and $^{137}$Cs radioactivity on 11 March 2011.
also two orders higher than that of Ukraine after the Chernobyl accident (Matsunaga et al., 1998), which is regarded as supplying suspended solids from the watershed and resuspension of river bottom sediments by rain events.

The cumulative $^{134}\text{Cs}$ and $^{137}\text{Cs}$ inventory from the surface down to depth in undisturbed soils in Fukushima Prefecture confirms that >90% of the total $^{134}\text{Cs}$ and $^{137}\text{Cs}$ in the soil profile was found within the upper 5 cm layer at cropland and grassland sites (Koarashi et al., 2012). Surface erosion processes in watersheds have been studied by many researchers using $^{137}\text{Cs}$ derived from fallout as a tracer of suspended solids. Surface runoff generally does not occur in forested areas, but unmanaged Japanese cypress plantations often have little surface cover. For that reason, surface runoff is generated in such areas during large rainstorms (Miyata et al., 2007; Gomi et al., 2008). Fukuyama et al. (2010) reported for different stand species that surface coverage and forest management practices affect the runoff of the surface-derived suspended solids at the catchment scale. To elucidate the effects of input of suspended solids on $K_d$ value, we compared the $K_d$ values after heavy rain events caused by a typhoon in September 2011 and July 2012. The suspended solid concentration was 0.20–0.41 g l$^{-1}$ for the September samples and 0.26–0.34 g l$^{-1}$ (Nagao unpublished data) for the July samples. The dissolved radioactivity in the activity was almost equal for both sampling dates: 0.0025–0.0046 Bq L$^{-1}$. The $K_d$ values of $^{134}\text{Cs}$ and $^{137}\text{Cs}$ for the July samples were 0.20–0.43 ×10$^3$ mL g$^{-1}$ and still high, although the values were 1/2–1/10 lower than that of 2011 September samples. These results suggest that the source of particulate forms of radioactivity is an important factor controlling $K_d$ values related to secondary dispersion after the nuclear accident. Suspended solids are derived primarily from the erosion of riverbanks and hill slope soils by overland flow (Knighton, 1998; Ritter et al., 2002). The riverbank erosion and resuspension of river bottom sediments are important sources of particulate $^{137}\text{Cs}$ in river waters after rain events (Matsunaga et al., 1991). Direct input of suspended solids eroded from the ground surface might be reflected in the higher values found for the Natsui River and the Same River after heavy rains at the initial stage of the Fukushima Dai-ichi NPP accident.

The $^{134}\text{Cs}$ and $^{137}\text{Cs}$ radioactivity in the riverside surface soil at Samegawa-Ohashi, located downstream along the Same River, was 22–970 Bq kg$^{-1}$ during September 2011–November 2012 (Ministry of Environment in Japan, 2012). The radioactivity at Idosawa-Ohashi, an upper downstream location, was 70–900 Bq kg$^{-1}$ during September 2011–July 2012, but 1500–7100 Bq kg$^{-1}$ on 2 September and 9 November 2012 at high water level conditions. These results suggest that $^{134}\text{Cs}$ and $^{137}\text{Cs}$ were deposited on the riverside surface soil as a result of rain events.

The radioactivity of $^{134}\text{Cs}$ and $^{137}\text{Cs}$ deposited in river bottom sediments was approx. 770 Bq kg$^{-1}$-sediment for the Natsui River and approx. 2000 Bq kg$^{-1}$ for the Same River during May–July 2011, but 10–250 Bq kg$^{-1}$ during September–November 2011–November 2012 (Ministry of the Environment, 2012). The radioactivity varied with sampling, although the samplings were conducted at fixed stations in each river. The river bottom sediments are sandy, so the apparent residence time of fine particles might be short in the Natsui River and the Same River because of flushing out that occurs concomitantly with rain events. Therefore, the contaminated area around the river basin and river bottom sediments plays an important role as a source of particulate phase $^{134}\text{Cs}$ and $^{137}\text{Cs}$ in river waters.

Figure 6 presents a schematic illustration showing the export of radioactivity from the watershed to the river. Erosion of the riverbank and ground surfaces, along with re-suspension of river bottom sediments, occurs during rain events, thereby increasing the amounts of suspended solids and the radioactivity of $^{134}\text{Cs}$ and $^{137}\text{Cs}$ associated with riverine suspended solids. Similar results have been reported for the Kuji River, running through Ibaraki and Fukushima prefectures (Matsunaga et al. 1998). Increased erosion and radioactivity of $^{137}\text{Cs}$ derived from global fallout have also been observed at the Kuzuryuu River in 2009 after rain events (Nagao, unpublished data).

### 3.4 Estimation of radioactivity flux from watersheds

To ascertain the impacts of heavy rains on the transport of radioactivity and its dispersion to coastal ocean, flux of radioactivity from land areas through rivers to the ocean was estimated simply based on our measurements, as shown in Table 1, and from water discharge data (Fukushima Prefectural Government, 2011). In the Natsui River, the averaged radioactivity was 0.0493 ± 0.0276 Bq L$^{-1}$ for $^{134}\text{Cs}$ and 0.0566 ± 0.0315 Bq L$^{-1}$ for $^{137}\text{Cs}$, except for the sample taken after the heavy rain. Mean monthly water discharge was 17.6 ± 30.8 m$^3$ s$^{-1}$ during March–December 2011. The fluxes ($F$) of $^{134}\text{Cs}$ and $^{137}\text{Cs}$ are defined as

$$F = C \times WD \times T, \quad (2)$$

where $C$ stands for the mean radioactivity of $^{137}\text{Cs}$ except for the heavy rain sample, WD signifies the mean monthly water discharge, and $T$ denotes the period of 11 March–31 December 2011. The fluxes of radioactivity from land to ocean are estimated as 2.3 ± 4.2 Bq/305 days for $^{134}\text{Cs}$ and 2.6 ± 4.8 × 10$^4$ Bq/305 days for $^{137}\text{Cs}$. The radioactivity of $^{137}\text{Cs}$ during 1 January–10 March might resemble that before the Fukushima Dai-ichi NPP accident and 2–3 order lower values. Therefore, the estimated value corresponds to the annual export flux.

Regarding the heavy rain event, the mean daily water discharge was 350 ± 264 m$^3$ s$^{-1}$, as calculated from hourly water discharge on 22 September 2011. The radioactivity levels of $^{134}\text{Cs}$ and $^{137}\text{Cs}$ were, respectively, 0.673 ± 0.004 Bq L$^{-1}$ and 0.853 ± 0.004 Bq L$^{-1}$. Therefore the fluxes from river to ocean were 2.0 ± 1.5 × 10$^4$ Bq day$^{-1}$ for $^{134}\text{Cs}$ and 2.6 ± 1.9 × 10$^4$ Bq day$^{-1}$ for $^{137}\text{Cs}$ on 22 September 2011. This value,
Fig. 5. X-ray diffraction patterns of suspended solids in the Natsui River (a) and Same River (b) after the heavy rain event (Typhoon Roke) in 2011. Legend for abbreviations of minerals: 14 Å – chlorite, vermiculite, smectite; M – mica; Am – amphibole; F – feldspar; Q – quartz; Cl/K – chlorite and/or kaolin.

Fig. 6. Schematic illustrations of export of $^{134}$Cs and $^{137}$Cs from watersheds to rivers.

which corresponds to that for March–December 2011, represents a rough estimation because of a lack of consideration for the six rain events with precipitation greater than 50 mm monitored at the Onahama site (JMA, 2012). The monitoring sites in the Natsui River and Same River are located at ca. 4–11 km from the river mouths. From the present dataset, we can not ascertain the transport behavior of radiocesium at the river mouth. However, Fujita et al. (2001) observed that turbidity in river water at the mouth of Same River increased to 150 ppm after the rain event with 47 mm day$^{-1}$ (14.4 mm hr$^{-1}$) in September 2000, which suggests that some suspended solids were transported from the lower river to the river mouth. Consequently, we can infer the importance of rain events for radiocesium transport from watersheds to rivers during the year.

Table 4 shows flux data of $^{137}$Cs for the Natsui River and the Same River together with those from other river systems. The flux at the Same River was estimated using a simple method similar to that used in this study. As Table 4 shows, the export flux in Fukushima rivers in 2011 is one order higher than that of the Tone River in 1985–1986 and the Kuji River in 1988–1989. The heavy rain event by Typhoon Roke on 22 September occupied 30–50 % of the annual export flux. These results indicate that the heavy rain event is one factor determining the export of $^{134}$Cs and $^{137}$Cs in the Natsui River and the Same River.

4 Summary

During July–December, 2011, research was conducted for the Natsui River and the Same River, which run through a contaminated watershed in Fukushima Prefecture, to elucidate the transport of $^{134}$Cs and $^{137}$Cs from land to coastal ocean area after the Fukushima Dai-ichi NPP accident. The levels of $^{134}$Cs and $^{137}$Cs radioactivity were measured for river water samples in normal flow conditions and high flow conditions caused by a heavy rain event with daily precipitation of 137 mm (Typhoon Roke) in September 2011. The $^{134}$Cs and $^{137}$Cs radioactivity levels in a normal flow condition were 0.009–0.098 Bq L$^{-1}$, but at a high flow condition, they were 0.36–0.85 Bq L$^{-1}$. Particulate phase of $^{134}$Cs and $^{137}$Cs was 21–56 % in a normal flow condition and ca. 100 % in a high flow condition. Therefore, radiocesium associated with suspended solids is exported in large quantities from river to coastal ocean areas by heavy rains. Export of $^{134}$Cs and $^{137}$Cs after the heavy rain event is estimated as roughly $2.0-2.6 \times 10^{10}$ Bq day$^{-1}$ for the Natsui River and $0.74-0.87 \times 10^{10}$ Bq day$^{-1}$ for the Same River. These values account for 30–50 % of the export of radiocesium for the 10 months of 11 March–31 December 2011 in both rivers. Therefore, results show that pulse input by heavy rain events is one important pathway of radiocesium from inland to coastal ocean areas in the southern part of Fukushima Prefecture, Japan.

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Table 4. Export fluxes of $^{137}$Cs from land to ocean in the Natsui, Same, Kuji and Tone rivers.

<table>
<thead>
<tr>
<th>River</th>
<th>Catchment area (km$^2$)</th>
<th>Export flux of $^{137}$Cs Annual (x10$^{10}$ Bq yr$^{-1}$)</th>
<th>Heavy rain (x10$^{10}$ Bq day$^{-1}$)</th>
<th>Contribution of heavy rain (%)</th>
<th>Year</th>
<th>Major origin of $^{137}$Cs</th>
</tr>
</thead>
<tbody>
<tr>
<td>Natsui</td>
<td>749</td>
<td>$5.2 \pm 1.3$</td>
<td>$2.6 \pm 1.9$</td>
<td>50</td>
<td>2011</td>
<td>NA</td>
</tr>
<tr>
<td>Same</td>
<td>600</td>
<td>$2.9 \pm 1.2$</td>
<td>$0.87 \pm 0.92$</td>
<td>30</td>
<td>2011</td>
<td>NA</td>
</tr>
<tr>
<td>Kuji*</td>
<td>1490</td>
<td>$0.26$</td>
<td>-</td>
<td>-</td>
<td>1988</td>
<td>GF</td>
</tr>
<tr>
<td>Tone**</td>
<td>16840</td>
<td>$0.50 \pm 0.38$</td>
<td>-</td>
<td>-</td>
<td>1985</td>
<td>GF</td>
</tr>
</tbody>
</table>

NA indicates nuclear accident, and GF is global fallout.

* Flux estimates from $^{137}$Cs annual load (Matsunaga et al., 1991) and watershed area of the Kuji River.

** Flux estimates from the average radioactivity of $^{137}$Cs (Hirose et al., 1990) and mean annual water discharge in 1985 (MLIT, 2012).

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