



## Biomass burning fuel consumption rates: a field measurement database

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**Abstract.** Landscape fires show large variability in the amount of biomass or fuel consumed per unit area burned. Fuel consumption (FC) depends on the biomass available to burn and the fraction of the biomass that is actually combusted, and can be combined with estimates of area burned to assess emissions. While burned area can be detected from space and estimates are becoming more reliable due to improved algorithms and sensors, FC is usually modeled or taken selectively from the literature. We compiled the peer-reviewed literature on FC for various biomes and fuel categories to understand FC and its variability better, and to provide a database that can be used to constrain biogeochemical models with fire modules. We compiled in total 77 studies covering 11 biomes including savanna (15 studies, average FC of 4.6 t DM (dry matter) ha<sup>-1</sup> with a standard deviation of 2.2), tropical forest ( $n = 19$ , FC = 126 ± 77), temperate forest ( $n = 12$ , FC = 58 ± 72), boreal forest ( $n = 16$ , FC = 35 ± 24), pasture ( $n = 4$ , FC = 28 ± 9.3), shifting cultivation ( $n = 2$ , FC = 23, with a range of 4.0–43), crop residue ( $n = 4$ , FC = 6.5 ± 9.0), chaparral ( $n = 3$ , FC = 27 ± 19), tropical peatland ( $n = 4$ , FC = 314 ± 196), boreal peatland ( $n = 2$ , FC = 42 [42–43]), and tundra ( $n = 1$ , FC = 40). Within biomes the regional variability in the number of measurements was sometimes large, with e.g. only three measurement locations in boreal Russia and 35 sites in North America. Substantial regional differences in FC were found within the defined biomes: for example, FC of temperate pine forests in the USA was 37 % lower than Australian forests dominated by eucalypt trees. Besides showing the differences between biomes, FC estimates were also grouped into different fuel classes. Our results highlight the large variability in FC, not only between biomes but also within biomes and fuel classes. This implies that substantial uncertainties are associated with using biome-averaged values to represent FC for whole biomes. Comparing the compiled FC values with co-located Global Fire Emissions Database version 3 (GFED3) FC indicates that modeling studies that aim to represent variability in FC also within biomes, still require improvements as they have difficulty in representing the dynamics governing FC.

## 1 Introduction

Landscape fires occur worldwide in all biomes except deserts, with frequencies depending mostly on type of vegetation, climate, and human activities (Crutzen, 1990; Cooke and Wilson, 1996; Andreae and Merlet, 2001; Bowman et al., 2009). The amount of fire-related research is increasing, partly due to improved abilities to monitor fires around the world using satellite data and appreciation of the important role of fires in the climate system and in air quality (Bowman et al., 2009; Johnston et al., 2012). Studies focusing on the effects of fires on the atmosphere require accurate trace

gas and particle emissions estimates. Historically, these are based on the Seiler and Crutzen (1980) equation, multiplying burned area, fuel loads (abbreviated as “FL” in the remainder of the paper), combustion completeness (abbreviated as “CC” in the remainder of the paper), and emission factors over time and space of interest.

These four properties are obtained in different ways and, generally, uncertainties are substantial (van der Werf et al., 2010). The burned area may be estimated directly from satellite observations, with the MODerate resolution Imaging Spectroradiometer (MODIS) 500 m maps (Roy et al., 2005; Giglio et al., 2009) being currently the most commonly used products for large-scale assessments. Although small fires and fires obscured by forest canopies escape detection with this method (Randerson et al., 2012), the extent of most larger fires can be relatively well constrained in this way.

With burned area estimates improving, the other parameters may become the most uncertain component when estimating emissions (French et al., 2004) as they are less easily observed from space. In general, the FL is equivalent to the total biomass available. New studies do provide estimates of standing biomass (e.g., Baccini et al., 2012). However, fires do not necessarily affect standing biomass. Especially in savannas the trees are usually protected from burning by a thick barch, and in some of the literature the FL therefore has a more restrictive definition, referring to only that portion of the total available biomass that normally burns under specified fire conditions, which is often only the fine surface fuels. In both definitions, the FL is typically expressed as the mass of fuel per unit area on a dry weight basis. CC corresponds to the fraction of fuel exposed to a fire that was actually consumed or volatilized. Just like total FL, CC cannot be directly derived from satellite observations. Instead, these quantities are usually based on look-up tables of biome-averaged values, or are calculated from global vegetation models including dynamic global vegetation models (DGVM, e.g., Kloster et al., 2010) and biogeochemical models (e.g., Hély et al., 2007; van der Werf et al., 2010).

Another approach that has been developed over the past decade is the measurement of fire radiative power (FRP) (Wooster et al., 2003, 2005; Kaiser et al., 2012). FRP per unit area relates directly to the fuel consumption (abbreviated as “FC” in the remainder of the paper) rate, which again is proportional to the fire emissions. The FRP method has several advantages compared to the Seiler and Crutzen (1980) approach, such as the ability to detect smaller fires and the fact that the fire emissions estimates derived this way do not rely on FL or CC. One disadvantage is that the presence of clouds and smoke can prevent the detection of a fire, and the poor temporal resolution of polar orbiting satellites hampers the detection of fast-moving or short-lived fires (which still can show a burn scar in the burned area method), and makes the conversion of FRP to fire radiative energy (FRE, time-integrated FRP) difficult.

Finally, emission factors, relating the consumption of dry matter to trace gas and aerosol emissions of interest, are obtained by averaging field measurements for the different biomes. Andreae and Merlet (2001) have compiled these measurements into a database that is updated annually, while Akagi et al. (2011) used a similar approach to derive biome-averaged emission factors but focused on measurement of fresh plumes only and provided more biome-specific information. The accompanying database is updated frequently and is online.

To improve and validate fire emissions models, it is crucial to gain a better overview of available FC measurements, as well as of the FL and CC components that together govern FC. This is obviously the case for emissions estimates based on burned area, but FRP estimates could also benefit from this information because one way to constrain these estimates is by dividing the fire-integrated FRE by the fire-integrated burned area, which in principle should equal FC.

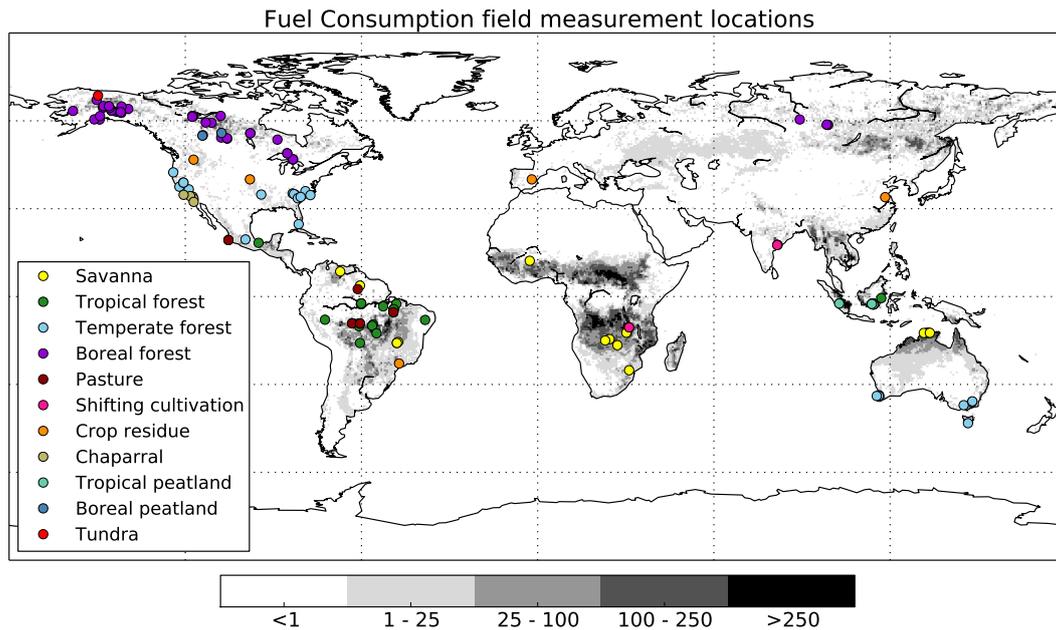
Over the last few decades, many field measurements of FL and CC have been made over a range of biomes and geographical locations. An examination of these studies revealed several generalities: forested ecosystems in general show relatively little variability in FL over time for a given location, but CC can vary due to weather conditions. Fine fuels usually burn more completely than coarser fuels, and therefore CC in grassland savannas is often higher than in forested ecosystems. While CC in the savanna biome shows relatively little variability over time, the FL can vary on monthly timescales depending on season, time since fire, and grazing rates. Another generalization that can be made is that FL in boreal and tropical forests are relatively similar, but the distribution into components (organic soil, boles, peat) is very different, with FL in tropical forests being mostly composed of above-ground biomass, while in the boreal region the organic soil (including fermentation and humus layers) represents a large part of the FL. Overall CC is often higher in tropical forests though, leading to higher FC values.

While these findings are relatively easy to extract from the body of literature, what is lacking is a universal database listing all the available measurements so that they can be compared in a systematic way, used to constrain models, and to identify gaps in our knowledge with regard to spatial representativeness. Building on Akagi et al. (2011), who listed 47 measurements for nine fuel types, this paper is a first attempt to establish a complete database listing all the available FC field measurements for the different biomes that were found in the peer-reviewed literature. We focus on FC estimates, but if FL and/or CC were reported separately, these were included as well. The database, available at <http://www.falw.vu/~gwerf/FC>, will be updated when new information becomes available. In follow-up papers we aim to provide more in-depth analyses of the variability we found; the goal of this paper is to give a quantitative overview of FC measurements made around the world to improve large-scale fire emissions assessments. This paper is organized as follows: in Sect. 2,

we list all the measurements and divide them into 11 different biomes. In that section we also provide a short summary of the methods used during the field campaigns, give a brief introduction to fire processes in each biome, and present data for different fuel classes (ground, surface, and crown fuels). Our findings are discussed in Sect. 3, and in addition a comparison between the FC field measurements and (1) the values used in the GFED3 (van der Werf et al., 2010) modeling framework, and (2) several FRE-derived estimates, is given. Finally, our results are summarized in Sect. 4.

## 2 Measurements

Figure 1 provides an overview of the locations where peer-reviewed FC was measured in the field, overlaid on mean annual fire carbon (C) emissions (van der Werf et al., 2010). Field measurements of FC were conducted in most fire-prone regions in the world, including the “arc of deforestation” in Amazonia, the boreal regions of North America, and savannas and woodlands in Africa, South America and Australia. Due to ecological, technical, and logistical reasons (e.g., wildfire versus prescribed fire), the FL and FC sampling procedures at these measurement locations have ranged in scope from simple and rapid visual assessment (e.g., Maxwell, 1976; Sandberg et al., 2001) to highly detailed measurements of complex fuel beds along lines (line transect method: van Wagner, 1968) or in fixed areas (planar intersect method; Brown, 1971) that take considerable time and effort. Most of the studies we found in the literature rely on the planar intersect method (PIM), where fuel measurement plots are typically divided into multiple, randomized smaller subplots. The (small-size) biomass in these subplots is oven dried and weighed both pre- and post-fire to estimate the CC and to determine the FC. The consumption of larger-size material (diameter > 10 cm) is often estimated based on experimental observations of randomly selected trunks and branches that were identified before the fire (Araújo et al., 1999). The PIM is mainly applied in prescribed burns, and obtaining FC measurements for large wildfires is logistically more challenging but can be based on comparing burned and adjacent unburned patches. Usually, the total FC of a fire is presented, but some studies also include separate values for different fuel categories of the total below-ground biomass (peat, organic soils, and roots) and total above-ground biomass (above-ground litter and live biomass). Diameters of woody fuels have been classified according to their “time-lag”, which refers to the length of time that a fuel element takes to respond to a new moisture content equilibrium (Bradshaw et al., 1983). The time lag categories traditionally used for fire behavior are specified as 1, 10, 100, and 1000 h, and correspond to round woody fuels in the size ranges of 0–0.635, 0.635–2.54, 2.54–7.62, and 7.62–20.32 cm, respectively. In this study, we used US fire management standards to classify fuels into three different categories: (1) ground (all materials lying beneath



**Figure 1.** Fuel consumption field measurement locations for different biomes. Background map shows annual GFED3 fire C emissions in  $\text{g C m}^{-2} \text{ year}^{-1}$ , averaged over 1997–2009.

the surface, including organic soil, roots, rotten buried logs, and other woody fuels), (2) surface (all materials lying on or immediately above the ground, including needles or leaves, grass, small dead wood, downed logs, stumps, large limbs, low brush, and reproduction) and (3) crown (aerial) fuel (all green and dead materials located in the upper forest canopy, including tree branches and crowns, snags, moss, and high brush).

Although a substantial body of grey literature of FC measurements is available, we focused on peer-reviewed studies. An exception was made for a few reports that focus on measurements conducted in the boreal forest and chaparral biome, because these reports were extensive and cited in peer-reviewed literature. Because the available data from peer-reviewed literature were obtained from a wide variety of sources spanning multiple decades, the reported FC data needed to be standardized. We converted all FC measurements to units of tons dry matter per hectare ( $\text{t ha}^{-1}$ ), which is the most commonly used unit. A carbon to dry matter conversion factor of 2 was used to convert carbon FC values to dry matter FC values. We note though that this conversion factor is not always representative of all biomes. Especially in the boreal regions – having a relative large contribution of organic soil fuels – but also in other biomes, this factor is sometimes lower, and therefore our approach may slightly overestimate FL and FC.

In Table 1 we present the FL, CC, and FC data compiled for 11 different biomes that are frequently used in global fire emissions assessments (e.g., van der Werf et al., 2010; Wiedinmeyer et al., 2011; Kaiser et al., 2012; Randerson

et al., 2012). Some studies provided data for specific fuel classes (e.g., ground fuels) only, while others estimated a total FC for both the below- and above-ground biomass. The data presented in Table 1 focused on FC. Additional studies on FL measurements exist and were not included here, but are listed in a spreadsheet that is available online at <http://www.falw.vu/~gwerf/FC>. These estimates were extensive, mostly for southern Africa (e.g., Scholes et al., 2011) and Australia (e.g., Rossiter et al., 2003). Including these additional field measurements may change regional FL averages. More specific details on the measurements and different fuel categories for each biome are listed in Sects. 2.1–2.11.

## 2.1 Savanna

Savanna fires in the tropics can occur frequently, in some cases annually. Their FL consists mainly of surface fuels (like grass and litter from trees), and is influenced both by rainfall of the previous years and the time since the last fire (Gill and Allan, 2008). Traditionally, (African) savannas are split into dry and wet forms (Menaut et al., 1995). This split occurs at a precipitation rate of about  $900 \text{ mm year}^{-1}$ . Most savanna fires burn due to human ignition, but it is believed that these systems are seldom ignition limited, and more often limited by available fuel (Archibald et al., 2010). Fire incidence generally increases after years of above-average rainfall, especially in dry savannas with low population densities (van Wilgen et al., 2004; Russell-Smith et al., 2007). As these systems are generally fuel limited, grass production and consumption by herbivores are very important factors controlling the extent of area burned, particularly in drier regions

**Table 1.** Location, fuel load (FL), combustion completeness (CC) and fuel consumption (FC) for field measurements conducted in the savanna (a), tropical forest (b), temperate forest (c), boreal forest (d), pasture (e), shifting cultivation (f), crop residue (g), chaparral (h), tropical peat (i), boreal peat (j), and tundra biome (k). Standard deviation (SD) is shown in parenthesis, and values indicated in bold were used to calculate the biome average.

Ref <sup>a</sup>	Lat. (°)	Lon. (°)	Location	FL (tha <sup>-1</sup> )	CC (%)	FC (tha <sup>-1</sup> )	Description
<b>(a)</b>							
1	25.15 S	31.14 E	Kruger Park, South Africa	<b>4.4 (1.4)</b>	<b>80 (16)</b>	<b>3.5 (1.4)</b>	Lowveld sour bushveld savanna
1	12.35 S	30.21 E	Kasanka National Park, Zambia	<b>5.4 (2.1)</b>	<b>81 (15)</b>	<b>4.2 (1.0)</b>	Dambo, miombo, chitemene
1	16.60 S	27.15 E	Choma, Zambia	<b>5.1 (0.4)</b>	<b>88 (2)</b>	<b>4.5 (-)</b>	Semi-arid miombo
2	14.52 S	24.49 E	Kaoma local forest, Zambia	<b>5.8 (3.8)</b>	<b>53 (32)</b>	<b>2.2 (1.2)</b>	Dambo and miombo
3	15.00 S	23.00 E	Mongu region, Zambia	<b>4.2 (0.8)</b>	<b>69 (21)</b>	<b>2.9 (0.9)</b>	Dambo and floodplain
4	12.22 N	2.70 W	Tiogo state forest, Senegal	<b>5.8 (1.6)</b>	<b>75 (15)</b>	<b>4.2 (0.7)</b>	Grazing and no grazing
5	15.84 S	47.95 W	Brasilia, Brazil	<b>8.3 (1.3)</b>	<b>88 (13)</b>	<b>7.2 (0.9)</b>	Different types of cerrado
6	8.56 N	67.25 W	Calaboza, Venezuela	<b>6.9 (2.3)</b>	<b>82 (17)</b>	<b>5.5 (1.9)</b>	Protected savanna for 27 years
7	15.51 S	47.53 W	Brasilia, Brazil	<b>8.3 (-)</b>	<b>90 (-)</b>	<b>7.5 (-)</b>	Campo limpo and campo sujo
8	15.84 S	47.95 W	Brasilia, Brazil	<b>8.9 (3.1)</b>	<b>92 (4.1)</b>	<b>8.2 (2.8)</b>	Different types of cerrado
9	3.75 N	60.50 W	Roraima, Brazil	<b>6.1 (3.6)</b>	<b>56 (27)</b>	<b>2.6 (0.9)</b>	Different types of cerrado
10	12.40 S	132.50 E	Kapalga, Kakadu, Australia	<b>4.8 (1.3)</b>	<b>94 (0.6)</b>	<b>4.5 (1.3)</b>	Woodland
11	12.30 S	133.00 E	Kakadu National Park, Australia	<b>5.6 (0.9)</b>	<b>91 (-)</b>	<b>5.1 (-)</b>	Tropical savanna
12	12.43 S	131.49 E	Wildman Reserve, Australia	<b>2.9 (1.8)</b>	<b>91 (14)</b>	<b>2.4 (1.1)</b>	Grass and woody litter
13	12.38 S	133.55 E	Arnhem Plateau, Australia	<b>3.6 (3.1)</b>	<b>44 (35)</b>	<b>1.4 (1.6)</b>	Early and late season fires
14	12.38 S	133.55 E	Arnhem Plateau, Australia	<b>8.5 (-)</b>	<b>39 (-)</b>	<b>4.8 (-)</b>	Grass and open woodland
15	17.65 N	81.75 E	Kortha Valasa and Kudura, India	<b>35 (6.4)</b>	<b>22 (7.7)</b>	<b>7.7 (2.6)</b>	Woodland
<b>(b)</b>							
5	4.30 S	49.03 W	Marabá, Pará, Brazil	<b>207 (-)</b>	<b>48 (-)</b>	<b>103 (-)</b>	Primary and secondary forest
16	2.29 S	60.09 W	Fazenda Dimona, Manaus, Brazil	<b>265 (-)</b>	<b>29 (-)</b>	<b>77 (-)</b>	200 ha clearing for pasture
17	7.98 S	38.32 W	Serra Talh., Pernambuco, Brazil	<b>74 (0.2)</b>	<b>87 (8.6)</b>	<b>64 (6.4)</b>	Second-growth tropical dry forest
18	4.50 S	49.01 W	Marabá, Pará, Brazil	<b>364 (-)</b>	<b>52 (-)</b>	<b>190 (-)</b>	Cleared for pastures
18	15.85 S	60.52 W	Santa Barbara, Rondônia, Brazil	<b>326 (-)</b>	<b>50 (-)</b>	<b>166 (-)</b>	Cleared for shifting cultivation
19	2.61 S	60.17 W	Manaus, Brazil	<b>425 (-)</b>	<b>25 (-)</b>	<b>107 (-)</b>	Tropical dense rainforest
20	9.11 S	63.16 W	Jamari, Rondônia, Brazil	<b>377 (31)</b>	<b>50 (4.5)</b>	<b>191 (33)</b>	Primary forest slash
21	2.61 S	60.17 W	Manaus, Brazil	<b>402 (-)</b>	<b>20 (-)</b>	<b>82 (-)</b>	Humid dense tropical forest
22	10.16 S	60.81 W	Ariguimes, Rondônia, Brazil	<b>307 (49)</b>	<b>36 (-)</b>	<b>110 (-)</b>	Open tropical forest
23	3.37 S	52.62 W	Altamira, Pará, Brazil	<b>263 (-)</b>	<b>42 (-)</b>	<b>110 (-)</b>	Lowland amazonian dense forest
24	2.50 S	48.12 W	Igarape do vinagre, Pará, Brazil	<b>214 (-)</b>	<b>20 (-)</b>	<b>43 (-)</b>	Tropical dense rainforest
25	5.35 S	49.15 W	Djair, Pará, Brazil	<b>121 (17)</b>	<b>43 (-)</b>	<b>52 (-)</b>	Slashed second-growth forest
25	9.20 S	60.50 W	Rondônia, Brazil	<b>118 (45)</b>	<b>56 (7.7)</b>	<b>65 (21)</b>	Second-, third-growth forest
25	4.30 S	49.03 W	José, Pará, Brazil	<b>64 (4.0)</b>	<b>86 (-)</b>	<b>55 (-)</b>	Third-growth forest
26	2.34 S	60.09 W	Fazenda dimona, Manaus, Brazil	<b>369 (187)</b>	<b>30 (-)</b>	<b>111 (-)</b>	Lowland amazonian dense forest
27	9.52 S	56.06 W	Alta floresta, Mato Grosso, Brazil	<b>496 (-)</b>	<b>39 (18)</b>	<b>192 (87)</b>	1, 4, and 9 ha clearings
28	9.97 S	56.34 W	Alta floresta, Mato Grosso, Brazil	<b>306 (-)</b>	<b>24 (-)</b>	<b>73 (-)</b>	Primary forest, 4 ha
29	12.53 S	54.88 W	Feliz Natal, Mato Grosso, Brazil	<b>219 (-)</b>	<b>71 (-)</b>	<b>155 (-)</b>	Seasonal semi-deciduous forest
30	7.90 S	72.44 W	Cruzeiro do Sul, Acre, Brazil	<b>583 (-)</b>	<b>39 (-)</b>	<b>226 (-)</b>	Primary forest, 4 ha clearing
31	18.35 N	95.05 W	Los Tuxtlas, Mexico	<b>403 (-)</b>	<b>95 (-)</b>	<b>380 (-)</b>	Evergreen tropical forest
32	19.30 N	105.3 W	San Mateo, Jalisco, Mexico	<b>127 (-)</b>	<b>71 (-)</b>	<b>91 (-)</b>	Tropical dry forest
33	0.52 S	117.01 E	East Kalimantan, Indonesia	<b>237 (106)</b>	<b>56 (24)</b>	<b>120 (47)</b>	Lightly and heavily disturbed stands
<b>(c)</b>							
34	34.80 N	82.60 W	Southern Appalachians, USA	<b>110 (-)</b>	<b>59 (-)</b>	<b>65 (-)</b>	Mixed pine hardwoods
34	35.21 N	83.48 W	Nantahala, N. Carolina, USA	<b>177 (49)</b>	<b>52 (5.5)</b>	<b>93 (34)</b>	Pine: jacob w. and e., devil den
34	36.00 S	79.10 W	Hillsborough, N. Carolina, USA	<b>21 (1.2)</b>	<b>11 (-)</b>	<b>2.3 (-)</b>	Loblolly pine forest floor
34	34.80 N	82.60 W	South East Piedmont, USA	-	-	<b>5.2 (3.4)</b>	Pinus taeda plantation, forest floor
34	37.50 N	122.00 W	South East Coastal plain, USA	-	-	<b>15 (9.1)</b>	Pine forest floor
35	34.82 N	94.13 W	Scott County, Arkansas, USA	<b>10 (-)</b>	<b>45 (-)</b>	<b>4.7 (-)</b>	Shortleaf pine grassland
36	36.60 N	118.81 W	Sequoia National Park, USA	<b>230 (-)</b>	<b>92 (-)</b>	<b>212 (-)</b>	Mixed conifer trees
37	38.90 N	120.67 W	Dark Canyon Creek, USA	<b>141 (49)</b>	<b>79 (-)</b>	<b>111 (-)</b>	Two week post fire
38	38.90 N	120.62 W	Blodgett Forest, California, USA	<b>154 (-)</b>	<b>70 (-)</b>	<b>108 (-)</b>	Mixed conifer: moist and dry burn
39	24.73 N	81.40 W	National Key Deer Refuge, USA	<b>23 (5.9)</b>	<b>57 (11)</b>	<b>13 (4.3)</b>	Pine forest, potential fuels
40	42.40 N	124.10 W	Southwestern Oregon, USA	-	-	<b>39 (-)</b>	Mixed conifer forest
41	33.56 N	81.70 W	Savannah River, USA	<b>19 (-)</b>	<b>55 (-)</b>	<b>11 (-)</b>	Mature loblolly, old longleaf pine
42	34.63 N	77.40 W	Camp Lejeune, N. Carolina, USA	<b>11 (3.8)</b>	<b>45 (29)</b>	<b>5.6 (4.7)</b>	Pine understory
42	34.01 N	80.72 W	Fort Jackson, S. Carolina, USA	<b>10 (-)</b>	<b>54 (-)</b>	<b>6.3 (2.0)</b>	Pine understory
34	36.00 S	148.00 E	Southeastern Australia	<b>79 (-)</b>	<b>84 (-)</b>	<b>67 (-)</b>	27 year old pine plantation
43	33.68 S	116.25 E	Wilga, Australia	<b>48 (-)</b>	<b>76 (-)</b>	<b>28 (-)</b>	Eucalypt forest
43	34.20 S	116.34 E	Quillben, Australia	<b>183 (-)</b>	<b>46 (-)</b>	<b>58 (-)</b>	Eucalypt forest
43	33.91 S	116.16 E	Hester, Australia	<b>101 (-)</b>	<b>68 (-)</b>	<b>53 (-)</b>	Eucalypt forest
43	37.09 S	145.08 E	Tallarook, Victoria, Australia	<b>60 (-)</b>	<b>61 (-)</b>	<b>27 (-)</b>	Eucalypt forest
43	33.93 S	115.46 E	McCorkhill, Australia	<b>70 (-)</b>	<b>78 (-)</b>	<b>43 (-)</b>	Eucalypt forest
43	43.22 S	146.54 E	Warra, Tasmania	<b>644 (-)</b>	<b>62 (-)</b>	<b>299 (-)</b>	Eucalypt forest
43	35.77 S	148.03 E	Tumbarumba, Australia	<b>99 (-)</b>	<b>70 (-)</b>	<b>47 (-)</b>	Eucalypt forest
44	19.50 N	99.50 W	Mexico City, Mexico	-	-	<b>17 (12)</b>	Pine-dominated forest

Table 1. Continued.

Ref <sup>a</sup>	Lat. (°)	Lon. (°)	Location	FL (tha <sup>-1</sup> )	CC (%)	FC (tha <sup>-1</sup> )	Description
<b>(d)</b>							
45	46.98 N	83.43 W	Aubinadong River, ON, Canada	<b>99 (4.2)</b>	<b>66 (5.4)</b>	<b>34 (6.6)</b>	Different depth classes used
46	46.78 N	83.33 W	Sharpsand Creek, ON, Canada	<b>48 (10)</b>	<b>49 (18)</b>	<b>23 (7.6)</b>	Immature jack pine
47	48.92 N	85.29 W	Kenshoe Lake, ON, Canada	<b>332 (-)</b>	<b>7.5 (-)</b>	<b>24 (-)</b>	Surface and crown
48	63.38 N	158.25 W	Innoko, Alaska, USA	–	–	<b>37 (7.0)</b>	Black spruce forest/shrub/bog
49	64.45 N	148.05 W	Rosie Creek, Alaska, USA	–	–	<b>83</b>	Ground fuels
49	60.43 N	149.17 W	Granite Creek, Alaska, USA	–	–	<b>30</b>	Ground fuels
49	67.14 N	150.18 W	Porcupine, Alaska, USA	–	–	<b>25</b>	Ground fuels
49	63.12 N	143.59 W	Tok River, Alaska, USA	–	–	<b>51</b>	Ground fuels
49	63.45 N	145.12 W	Dry Creek, Alaska, USA	–	–	<b>41</b>	Ground fuels
49	63.08 N	142.30 W	Tetlin, Alaska, USA	–	–	<b>56</b>	Ground fuels
49	63.50 N	145.15 W	Hajdukovich Creek, Alaska, USA	–	–	<b>129</b>	Ground fuels
50	61.60 N	117.20 W	Fort Providence, NT, Canada	<b>83 (10)</b>	<b>44 (7.6)</b>	<b>36 (5.8)</b>	Jack pine and black spruce
51	65.10 N	147.30 W	Alaska, USA	–	–	<b>19 (1.7)</b>	Forest floor
52	64.40 N	145.74 W	Delta Junction, Alaska, USA	<b>75 (-)</b>	<b>48 (-)</b>	<b>35 (-)</b>	Ground fuels: (non)-permafrost
53	53.92 N	105.70 W	Montreal Lake, SK, Canada	<b>43 (4.0)</b>	<b>62 (7.5)</b>	<b>26 (3.9)</b>	Spruce, pine, mixed wood
54	65.03 N	147.85 W	Fairbanks, Alaska, USA	<b>95 (17)</b>	<b>61 (17)</b>	<b>57 (19)</b>	Differently facing slopes
55	46.87 N	83.33 W	Sharpsand Creek, ON, Canada	<b>13 (2.0)</b>	<b>69 (32)</b>	<b>9 (4.0)</b>	Experimental fire: forest floor
55	48.87 N	85.28 W	Kenshoe Lake, ON, Canada	<b>17 (3.0)</b>	<b>35 (13)</b>	<b>6 (2.0)</b>	Experimental fire: forest floor
55	61.37 N	117.63 W	Fort Providence, NT, Canada	<b>47 (9.0)</b>	<b>36 (9.0)</b>	<b>17 (3.0)</b>	Experimental fire: forest floor
55	61.69 N	107.94 W	Porter Lake, NT, Canada	<b>15 (0.0)</b>	<b>60 (20)</b>	<b>9 (3.0)</b>	Experimental fire: forest floor
55	55.07 N	114.03 W	Hondo, AB, Canada	<b>3 (1.0)</b>	<b>33 (35)</b>	<b>1 (1.0)</b>	Experimental fire: forest floor
55	59.31 N	111.02 W	Darwin Lake, NT, Canada	<b>18 (3)</b>	<b>72 (20)</b>	<b>13 (3.0)</b>	Experimental fire: forest floor
55	55.74 N	97.91 W	Burntwood River, MB, Canada	<b>72 (12)</b>	<b>26 (8.0)</b>	<b>19 (5.0)</b>	Wildfire: forest floor
55	54.29 N	107.78 W	Green Lake, SK, Canada	<b>36 (13)</b>	<b>86 (54)</b>	<b>31 (16)</b>	Wildfire: forest floor
55	53.57 N	88.62 W	Kasabonika, ON, Canada	<b>69 (19)</b>	<b>55 (46)</b>	<b>38 (30)</b>	Wildfire: forest floor
55	55.74 N	97.85 W	Thompson, MB, Canada	<b>23 (14)</b>	<b>87 (63)</b>	<b>20 (8.0)</b>	Wildfire: forest floor
55	54.05 N	105.81 W	Montreal Lake, SK, Canada	<b>61 (41)</b>	<b>57 (47)</b>	<b>35 (17)</b>	Wildfire: forest floor
55	64.06 N	139.43 W	Dawson City, YT, Canada	<b>84 (30)</b>	<b>46 (31)</b>	<b>39 (22)</b>	Wildfire: forest floor
55	59.40 N	113.03 W	Wood Buffalo Nat. Pk., Canada	<b>37 (9.0)</b>	<b>59 (35)</b>	<b>22 (12)</b>	Wildfire: forest floor
56	60.49 N	150.98 W	Soldotna, Alaska, USA	<b>91 (22)</b>	<b>37 (5.2)</b>	<b>33 (4.8)</b>	Mystery creek 1–3
56	61.61 N	149.04 W	Palmer, Alaska, USA	<b>84 (4.2)</b>	<b>61 (3.5)</b>	<b>51 (5.7)</b>	Deshka 1–2
56	62.69 N	141.77 W	Tetlin Refuge, Alaska, USA	<b>105 (16)</b>	<b>45 (15)</b>	<b>49 (20)</b>	Tetlin, chisana 1–4
56	64.87 N	147.71 W	Fairbanks, Alaska, USA	<b>86 (16)</b>	<b>36 (23)</b>	<b>32 (22)</b>	Bonanza creek, frostfire
57	63.00 N	142.00 W	Alaska, USA	<b>152 (-)</b>	<b>59 (-)</b>	<b>90 (-)</b>	Black spruce forest
58	65.00 N	146.00 W	Alaska, USA	<b>72 (-)</b>	<b>58 (-)</b>	<b>40 (-)</b>	Black spruce forest
59	60.45 N	89.25 E	Bor, Krasnoyarsk, Russia	<b>34 (-)</b>	<b>50 (-)</b>	<b>17 (-)</b>	Pine-lichen forest and litter
60	58.58 N	98.92 E	Lower Angara, Russia	<b>54 (12)</b>	<b>31 (15)</b>	<b>17 (8.6)</b>	Scots pine, larch mixed wood
60	58.70 N	98.42 E	Lower Angara, Russia	<b>43 (-)</b>	<b>42 (-)</b>	<b>18 (-)</b>	Scots pine, larch mixed wood
<b>(e)</b>							
20	9.17 S	63.18 W	Jamari, Rondônia, Brazil	<b>66 (13)</b>	<b>31 (10)</b>	<b>21 (17)</b>	12 year old pasture site
61	5.30 S	49.15 W	Francisco, Pará, Brazil	<b>53 (4.8)</b>	<b>83 (-)</b>	<b>44 (-)</b>	Two slash fires prior to burning
61	9.20 S	60.50 W	João and Durval, Rondônia, Brazil	<b>96 (-)</b>	<b>34 (-)</b>	<b>30 (-)</b>	4 year old pasture site
62	2.54 N	61.28 W	Vila de Apiau, Roraima, Brazil	<b>119 (-)</b>	<b>20 (-)</b>	<b>24 (-)</b>	Pasture and forest
32	19.30 N	105.3 W	San Mateo, Jalisco, Mexico	<b>35 (-)</b>	<b>69 (-)</b>	<b>23 (-)</b>	High and low severity
<b>(f)</b>							
63	10.53 S	31.14 E	Kasama, Zambia	<b>75 (-)</b>	<b>64 (-)</b>	<b>43 (-)</b>	Shifting cultivation
64	17.59 N	81.55 E	Damanapalli and Velegapalli, India	<b>14 (-)</b>	<b>30 (-)</b>	<b>4.0 (-)</b>	Shifting cultivation in dry forest
<b>(g)</b>							
65	40.00 N	2.00 W	Spain, Europe	<b>1.4 (-)</b>	<b>80 (-)</b>	<b>1.1 (-)</b>	Cereal crops
66	22.85 S	47.60 W	Piracicaba, Sao Paulo, Brazil	–	–	<b>20 (-)</b>	Sugar cane
67	33.94 N	118.33 E	Suqian, China	<b>6.7 (1.2)</b>	<b>44 (4.6)</b>	<b>2.9 (0.5)</b>	Mix (wheat, rice, corn, potato)
68	40.00 N	98.00 E	North America	<b>2.4 (3.6)</b>	<b>86 (6.0)</b>	<b>2.1 (3.2)</b>	Mix of crop types
68	46.73 N	117.18 E	North America	<b>12 (-)</b>	<b>90 (-)</b>	<b>11 (-)</b>	Seedgrass
<b>(h)</b>							
69	34.10 N	117.47 W	Lodi Canyon, California, USA	–	–	<b>45 (-)</b>	Prescribed chaparral fire
70	33.33 N	117.16 W	Bear Creek, California, USA	<b>60 (5.9)</b>	<b>83 (6.0)</b>	<b>50 (8.4)</b>	Mature caenothus and chamise
70	34.29 N	118.33 W	Newhall, California, USA	<b>20 (6.7)</b>	<b>75 (4.0)</b>	<b>15 (5.4)</b>	Mature chamise
70	32.32 N	117.15 W	TNC, California, USA	<b>21 (-)</b>	<b>77 (-)</b>	<b>16 (-)</b>	Young and healthy chamise
42	34.73 N	120.57 W	Vandenberg, California, USA	<b>14 (-)</b>	<b>68 (-)</b>	<b>10 (-)</b>	Coastal sage and maritime chaparral

Table 1. Continued.

Ref <sup>a</sup>	Lat. (°)	Lon. (°)	Location	FL (t ha <sup>-1</sup> )	CC (%)	FC (t ha <sup>-1</sup> )	Description
<b>(i)</b>							
71	2.52 S	113.79 E	Kalimantan, Indonesia	–	–	<b>500 (–)</b>	Peat and overstory
72	2.50 S	114.17 E	Palangka Raya, Indonesia	<b>399 (11)</b>	<b>27 (4.7)</b>	<b>109 (19)</b>	Various peat fire fuels
73	2.37 S	102.68 E	Pelawan, Riau, Indonesia	45 (6.1)	81 (10)	37 (8.2)	Litter and branches
74	2.52 S	113.79 E	Kalimantan, Indonesia	–	–	<b>332 (6.4)</b>	Measured by LIDAR
<b>(j)</b>							
75	55.85 N	107.67 W	Patuanak, Canada	–	–	<b>42 (25)</b>	Continental and permafrost bogs
76	54.93 N	114.17 W	Chisholm, Canada	–	–	<b>43 (–)</b>	Hummocks and hollows
<b>(k)</b>							
77	68.58 N	149.72 W	Anaktuvuk River, Alaska, USA	<b>165 (15)</b>	<b>24 (5.0)</b>	<b>40 (9.0)</b>	Soil and plants

<sup>a</sup> References: (1) Shea et al. (1996)/Ward et al. (1996); (2) Hoffa et al. (1999); (3) Hély et al. (2003b); (4) Savadogo et al. (2007); (5) Ward et al. (1992); (6) Bilbao and Medina (1996); (7) Miranda et al. (1996); (8) De Castro and Kauffman (1998); (9) Barbosa and Fearnside (2005); (10) Cook et al. (1994); (11) Hurst et al. (1994); (12) Rossiter-Rachor et al. (2007); (13) Russell-Smith et al. (2009); (14) Meyer et al. (2012); (15) Prasad et al. (2001); (16) Fearnside et al. (1993); (17) Kauffman et al. (1993); (18) Kauffman et al. (1995); (19) Carvalho et al. (1995); (20) Guild et al. (1998); (21) Carvalho et al. (1998); (22) Graça et al. (1999); (23) Fearnside et al. (1999); (24) Araújo et al. (1999); (25) Hughes et al. (2000a); (26) Fearnside et al. (2001); (27) Carvalho et al. (2001); (28) Christian et al., 2007/Soares Neto et al. (2009); (29) Righi et al. (2009); (30) Carvalho Jr. et al. (2014); (31) Hughes et al. (2000b); (32) Kauffman et al. (2003); (33) Toma et al. (2005); (34) Carier et al. (2004); (35) Sparks et al. (2002); (36) Stephens and Finney (2002); (37) Béche et al. (2005); (38) Hille and Stephens (2005); (39) Sah et al. (2006); (40) Campbell et al. (2007); (41) Goodrick et al. (2010); (42) Yokelson et al. (2013); (43) Hollis et al. (2010); (44) Yokelson et al. (2007); (45) Stocks et al. (1987a); (46) Stocks et al. (1987b); (47) Stocks (1989); (48) Goode et al. (2000); (49) Kasichke et al. (2000); (50) Stocks et al. (2004); (51) Harden et al. (2004); (52) Harden et al. (2006); (53) de Groot et al. (2007); (54) Kane et al. (2007); (55) de Groot et al. (2009); (56) Ottmar and Sandberg (2010); (57) Turetsky et al. (2011); (58) Boby et al. (2010); (59) FIRESCAN Science Team (1996); (60) Ivanova et al. (2011); (61) Kauffman et al. (1998); (62) Barbosa and Fearnside (1996); (63) Stromgaard, 1985; (64) Prasad et al. (2000); (65) Zarate et al. (2005); (66) Lara et al. (2005); (67) Yang et al. (2008); (68) McCarty et al. (2011); (69) Cofer III et al. (1988); (70) Hardy et al. (1996); (71) Page et al. (2002); (72) Usup et al. (2004); (73) Saharjo and Nurhayati (2006); (74) Ballhorn et al. (2009); (75) Turetsky and Wieder (2001); (76) Benscoter and Wieder (2003); (77) Mack et al. (2011).

where rainfall can vary strongly between years (Menaut et al., 1991; Cheney and Sullivan, 1997; Russell-Smith et al., 2007). Grass production controls fire spread because low-biomass grasslands have less continuous fuel swards and also because they burn at lower intensities, which reduces the probability of spread. In wet savannas, the grass production is poorly correlated with rainfall and much higher than in dry savannas (10 to 20 t ha<sup>-1</sup> year<sup>-1</sup>, Gignoux et al., 2006). This results in higher-intensity fires, keeping the landscape relatively open. In Australia, the division into dry and wet savannas is less clear. Annual grass production is typically low (less than 3 t ha<sup>-1</sup> year<sup>-1</sup>), even for precipitation rates of 2000 mm year<sup>-1</sup>. This difference is mostly due to the fact that Australia's native grasses are limited by nitrogen availability at high rainfalls, something African grasses such as *Andropogon gayanus* overcome through various mechanisms (Rossiter-Rachor et al., 2009).

Miombo woodlands in Africa are high-rainfall savannas where up to 40% of the fuel can be provided by litter from trees (Frost et al., 1996). A similar type of vegetation can be found in Brazil, mainly consisting of woodlands with a closed canopy of tall shrubs and scattered trees (Cerrado denso). We found several measurements conducted in Miombo woodlands, as well as field measurements in the Brazilian Cerrado denso. Moreover, one study was found for an Indian deciduous forest, which can be classified as wooded savanna and thus the savanna biome (Ratnam et al., 2011).

To calculate averages, we divided the savanna biome into grassland savanna and wooded savanna by using the fuel type description that was provided in each study. The savanna measurements presented in Table 1a were taken between

1990 and 2009, and represent 17 unique measurement locations (Fig. 1) taken from 15 different studies. For all measurements conducted, we found an average FL of  $7.6 \pm 6.5$  t ha<sup>-1</sup> and FC of  $4.6 \pm 2.2$  t ha<sup>-1</sup>. The average of the CC values as presented in the different studies indicated a value of  $71 \pm 26$ %, higher than the ratio derived from the average FL and FC (61%) above. This difference is because not all FC measurements reported FL. Within the savanna biome, regional differences were found (Fig. 2): FL and FC for South American savannas,  $8.2 \pm 1.6$  and  $6.0 \pm 2.4$  t ha<sup>-1</sup>, respectively, were nominally higher than the ones measured in the savannas of Australia ( $5.1 \pm 2.2$  and  $3.6 \pm 1.6$  t ha<sup>-1</sup>). Measurements conducted in Africa, contributing to roughly 40% of all measurements in the biome, showed the lowest FC ( $3.4 \pm 1.0$  t ha<sup>-1</sup>) of all regions. A larger number of measurements are required to say conclusively whether these differences are statistically significant. To show the difference between grassland savannas and wooded savannas, data of both types of savanna are also provided in Fig. 2. For grassland savannas the average FL was relatively low ( $5.3 \pm 2.0$  t ha<sup>-1</sup>) and the CC high ( $81 \pm 16$ %), yielding an average FC of  $4.3 \pm 2.2$  t ha<sup>-1</sup>. Wooded savannas, on the other hand, had a higher FL ( $11 \pm 9.1$  t ha<sup>-1</sup>) but lower CC ( $58 \pm 32$ %), and therefore the average FC of  $5.1 \pm 2.2$  t ha<sup>-1</sup> was only slightly higher than the one found for grasslands.

In Table 2 these values are given for different fuel categories. For the savanna biome most of the fuels were classified as surface fuels (Table 2a). In general, fuels with a large surface-area-to-volume ratio (like litter, grass and dicots) had a high CC of at least 88%. CC values were significantly lower for the woody debris classes, with a minimum of  $21 \pm 12$ % found for woody fuels with a diameter larger than

2.54 cm (100 h fuel). FC for the different fuel types was between 0.3 and 1.9 t ha<sup>-1</sup>, with litter having the highest values. In general, the total sum of different fuel categories agrees well with the biome-averaged values presented. However, not all measurements distinguished between fuel categories and therefore small discrepancies were sometimes found: for FC in the savanna biome, for example, the sum of different fuel categories is 5.3 t ha<sup>-1</sup>, slightly higher than the biome average of 4.6 ± 2.2 t ha<sup>-1</sup>.

## 2.2 Tropical forest

Tropical evergreen forests are generally not susceptible to fire except during extreme drought periods (e.g., Field et al., 2009; Marengo et al., 2011; Tomasella et al., 2013) due to their dense canopy cover keeping humidity high and wind speed low, and also because the amount of fuel on the surface is low due to rapid decomposition. Human activities have resulted in fire activity in tropical forests, often with the goal to clear biomass and establish pasture or cropland. These deforestation fires can be small-scale (e.g., shifting cultivation, discussed in Sect. 2.6) or on a large scale with the aid of heavy machinery. In the latter case, biomass is often piled in windrows after the first burn and subject to additional fires during the same dry season to remove the biomass more completely. In large-scale deforestation regions like the state of Mato Grosso in the Brazilian Amazon, the expansion of mechanized agriculture could result in increased fuel consumed per unit area (Cardille and Foley, 2003; Yokelson et al., 2007a). All these fires, but also selective logging, may lead to more frequent accidental fires as fragmented forests are more vulnerable to fire (Nepstad et al., 1999; Siegert et al., 2001; Pivello, 2011).

The total FL in tropical forests is mostly determined by the tree biomass (surface and canopy fuels) and generally on the order of a few hundred t ha<sup>-1</sup>. CC depends partly on the size of the clearing and on the curing period. In general, the CC for tropical forest clearings is lower than 50 % (Balch et al., 2008), but when there is a long (more than a year) lag between slash and burning the CC might increase to 60 % and more (Carvalho et al., 2001). The El Niño–Southern Oscillation (ENSO) phenomenon may also have a large effect on fuel conditions over tropical regions. Large-scale fires have been shown to occur in South America, Southeast Asia, and Africa in ENSO years, thereby likely increasing CC due to drought conditions (Chen et al., 2011; Field et al., 2009; Hély et al., 2003a).

The 22 unique measurement locations shown in Table 1b cover Brazil (19), Mexico (2), and Indonesia (1). In general, measurement sites were divided into several smaller subplots and the forest was slashed at the beginning of the dry season. The biomass was then weighed using the PIM. After about 2 months the plots were set on fire and the remaining biomass was weighed within 1 week after the burn. The average FL for the whole biome was 285 ± 137 t ha<sup>-1</sup>, CC averaged

49 ± 22 %, and total FC was 126 ± 77 t ha<sup>-1</sup>. Since more than 90 % of all measurements were conducted in Brazil (Fig. 3), the biome-averaged values are biased towards measurements conducted there. Studies conducted in Mexican and Indonesian tropical forests reported average FL of 265 and 237 t ha<sup>-1</sup>, respectively. Surprisingly, the CC of tropical forest in Mexico was the highest of all the studies (83 % on average), resulting in an average FC of 236 t ha<sup>-1</sup>, which was significantly higher than values found for both Brazil (117 ± 56 t ha<sup>-1</sup>) and Indonesia (120 ± 47 t ha<sup>-1</sup>). However, due to the small number of measurements conducted in Mexico and Indonesia, these findings are not conclusive.

Different forest types may partly explain the discrepancy found, and therefore we distinguished between measurements conducted in primary tropical evergreen forest, secondary tropical evergreen forest, and tropical dry forest (Fig. 3). To distinguish between tropical dry forests and wooded savannas (Sect. 2.1), we harmonized with the emission factor compilation of Akagi et al. (2011) in which 60 % canopy cover (Hansen et al., 2003) was the delineation between both ecosystems. FL and FC were largest for primary tropical evergreen forests, with average values of 339 ± 104 and 143 ± 79 t ha<sup>-1</sup>, respectively. For secondary tropical evergreen forests these values were substantially lower (101 ± 32 and 57 ± 7.0 t ha<sup>-1</sup>) and comparable with tropical dry forests in South America and Mexico, where the average FL was 100 and the FC 78 t ha<sup>-1</sup>.

Different fuel categories for the tropical forest biome are presented in Table 2b and can be mainly classified as surface fuels, except for the attached foliage (crown fuels) and rootmat category (ground fuels). Large woody debris (diameter > 20.5 cm) and trunks – although not always taken into account in certain studies – correspond to a large part of the above-ground biomass (FL = 147 ± 83 t ha<sup>-1</sup>) but are usually only slightly burned during a forest clearing process (Carvalho et al., 1995), as shown by an average CC of 32 ± 23 % leading to an FC of this category of only 37 ± 32 t ha<sup>-1</sup>. Similar to the savanna biome, we found a high CC of at least 73 % for surface fuels with a large surface area to volume ratio (litter, leaves, and dicots). The small woody fuels (1 and 10 h) also had a high CC, and the CC of the woody debris generally decreased with increasing diameter. From an FC perspective, the most important fuel types in the tropical forest biome were litter (14 ± 8.4 t ha<sup>-1</sup>) and woody debris size classes with a diameter larger than 0.64 cm (15–37 t ha<sup>-1</sup>).

## 2.3 Temperate forest

Although accounting for only a small part of the global emissions, temperate forest fires frequently occur nearby the wildland–urban interface with important consequences for human safety and air quality. While tropical fires are largely intentionally ignited to pursue land management goals, the temperate forest is also subject to wildfires. Obtaining FC

**Table 2.** Fuel load (FL), combustion completeness (CC) and fuel consumption (FC) field measurements for different fuel categories within the savanna (a), tropical forest (b), temperate forest (c), and boreal forest (d) biomes. Standard deviation (SD) is shown in parentheses.

Cl <sup>a</sup>	Fuel category	FL (tha <sup>-1</sup> )	CC (%)	FC (tha <sup>-1</sup> )	References <sup>b</sup>
<b>(a)</b>					
S	Dicots	0.4 (0.5)	91 (12)	0.3 (0.3)	1, 2, 5
S	Grass-dormant	1.9 (1.4)	93 (14)	1.3 (0.5)	1, 2, 5
C	Grass-green	0.4 (0.2)	88 (23)	0.3 (0.1)	1, 2, 5
S	Litter	2.1 (0.5)	88 (13)	1.9 (0.5)	1, 2, 5, 8, 12, 15
S	Tree/shrub leaves	0.4 (0.8)	64 (12)	0.3 (0.6)	1, 2, 5
S	Woody debris (0–0.64 cm)	0.6 (0.7)	65 (16)	0.4 (0.5)	1, 2, 5, 8
S	Woody debris (0.64–2.54 cm)	0.9 (1.0)	39 (25)	0.5 (0.7)	1, 2, 5, 8
S	Woody debris (> 2.54 cm)	1.0 (1.1)	21 (12)	0.3 (0.3)	1, 2, 5, 8
<b>(b)</b>					
C	Attached foliage	3.8 (3.0)	94 (5.1)	3.6 (2.8)	5, 18, 20, 25, 32
S	Dicots	0.5 (0.3)	89 (23)	0.5 (0.3)	5, 18, 20, 25, 32
S	Leaves	13 (8.8)	73 (38)	11 (9.8)	16, 17, 19, 21, 24, 27, 28, 29
S	Litter	18 (9.9)	85 (30)	14 (8.4)	5, 17–29, 32
S	Liana	5.2 (0.8)	21 (35)	0.9 (1.4)	19, 21, 24
G	Rootmat	5.2 (2.7)	87 (13)	4.4 (2.2)	18, 20, 25
S	Woody debris (0–0.64 cm)	4.6 (2.8)	94 (4.8)	6.4 (8.6)	5, 17, 18, 20, 25, 32
S	Woody debris (0.65–2.54 cm)	17 (3.9)	87 (7.9)	15 (4.0)	5, 17, 18, 20, 25, 32
S	Woody debris (2.55–7.6 cm)	27 (15)	65 (19)	18 (13)	5, 17, 18, 20, 25, 32
S	Woody debris (7.6–20.5 cm)	45 (29)	41 (18)	18 (9.3)	5, 17, 18, 20, 25, 32
S	Woody debris (> 20.5 cm), trunks	147 (83)	32 (23)	37 (32)	5, 16, 18–23, 26–30
<b>(c)</b>					
G	Organic soil	49 (37)	48 (44)	25 (31)	34, 37, 38
S	Litter	20 (11)	81 (8.9)	17 (9.9)	34, 37, 38
S	Woody debris (0–0.64 cm)	1.2 (0.8)	87 (11)	1.0 (0.6)	36, 37, 38
S	Woody debris (0.65–2.54 cm)	5.2 (1.9)	79 (11)	4.0 (1.2)	36, 37, 38
S	Woody debris (2.55–7.6 cm)	6.0 (0.9)	73 (14)	4.3 (0.2)	36, 37, 38
S	Woody debris (7.6–20.5 cm sound)	16 (9.6)	38 (42)	6.2 (8.2)	36, 37, 38
G	Woody debris (7.6–20.5 cm rotten)	20 (4.1)	96 (5.4)	20 (4.8)	36, 37, 38
<b>(d)</b>					
G	Ground fuels (soil, forest floor)	50 (29)	51 (18)	32 (26)	44, 48, 49, 50, 51, 53, 54, 55, 57, 58
S	Surface fuels	44 (49)	52 (25)	12 (8.1)	44, 46, 49, 52, 55, 58, 59
C	Crown fuels	37 (70)	71 (29)	8.1 (6.9)	44, 46, 49, 57, 59

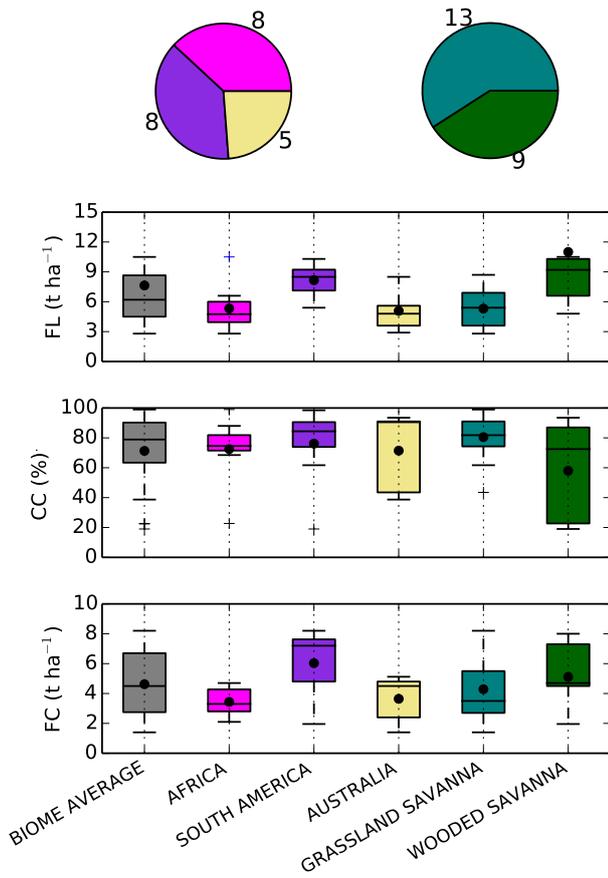
<sup>a</sup> Fuel category classification: S = surface fuels, G = ground fuels, C = crown fuels.

<sup>b</sup> References: (1) Shea et al. (1996)/Ward et al. (1996); (2) Hoffa et al. (1999); (5) Ward et al. (1992); (8) De Castro and Kauffman (1998); (12) Rossiter-Rachor et al. (2007); (15) Prasad et al. (2001); (16) Fearnside et al. (1993); (17) Kauffman et al. (1993); (18) Kauffman et al. (1995); (19) Carvalho et al. (1995); (20) Guild et al. (1998); (21) Carvalho et al. (1998); (22) Graça et al. (1999); (23) Fearnside et al. (1999); (24) Araújo et al. (1999); (25) Hughes et al. (2000a); (26) Fearnside et al. (2001); (27) Carvalho et al. (2001); (28) Christian et al. (2007)/Soares Neto et al. (2009); (29) Righi et al. (2009); (30) Carvalho Jr. et al. (2014); (32) Kauffman et al. (2003); (34) Carter et al. (2004); (36) Stephens and Finney (2002); (37) Bêche et al. (2005); (38) Hille and Stephens (2005); (45) Stocks et al. (1987a); (47) Stocks (1989); (49) Kasischke et al. (2000); (50) Stocks et al. (2004); (51) Harden et al. (2004); (52) Harden et al. (2006); (53) de Groot et al. (2007); (54) Kane et al. (2007); (55) de Groot et al. (2009); (56) Ottmar and Sandberg (2010); (58) Boby et al. (2010); (59) FIRESCAN Science Team (1996); (60) Ivanova et al. (2011).

measurements for wildfires is obviously challenging, so most information is derived from prescribed fires that allow researchers to measure pre-fire conditions. However, these fires may not always be a good proxy for wildfires. For example, wildfires in western conifer forests of the US are often crown fires (while prescribed fires usually only burn surface fuels).

Due to potential discrepancies with respect to FC, we distinguished between these fire types in Sect. 3.2.

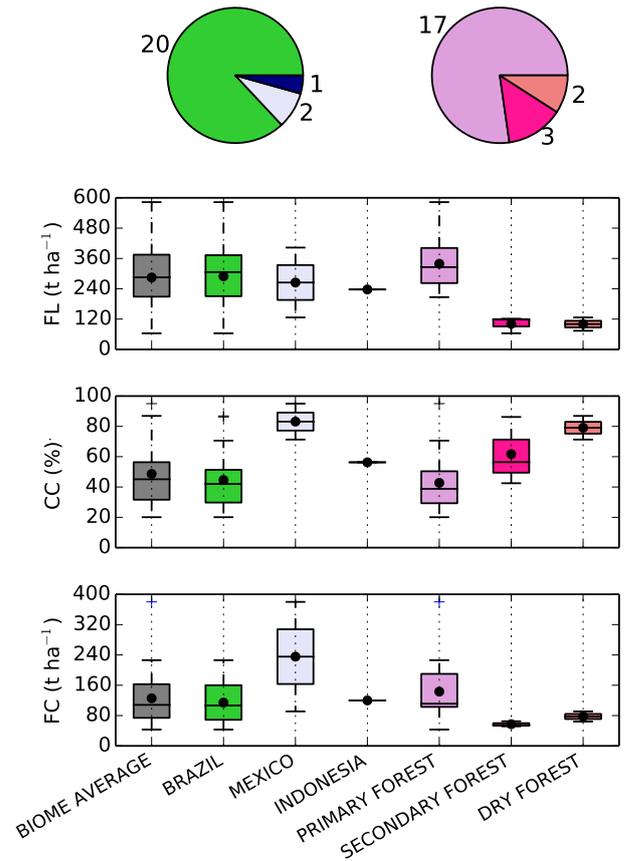
The 23 unique FC measurement locations for the temperate forest are from sites in North America (14), Australia (7), Tasmania (1) and Mexico (1), and were taken between 1983 and 2011 (Fig. 1). In general, measurements were conducted on sites that were divided into multiple, randomized subplots



**Figure 2.** Overview of field measurements of fuel load (FL), combustion completeness (CC), and fuel consumption (FC) in the savanna biome. The pie charts on top correspond to the number of unique measurement locations for different geographical regions (left) and savanna types (right), and in the box plots below field averages of FL, CC, and FC are presented. The boxes extend from the lower to upper quartile values of the measurement data, with a line at the median and a black filled circle at the mean. The whiskers extend from the box to show the range of the data, and outliers are indicated with pluses.

on which the pre-fire biomass was weighed according to the PIM. The sites were then burned and within a few days after the burn the post-fire biomass was gathered, dried and weighed.

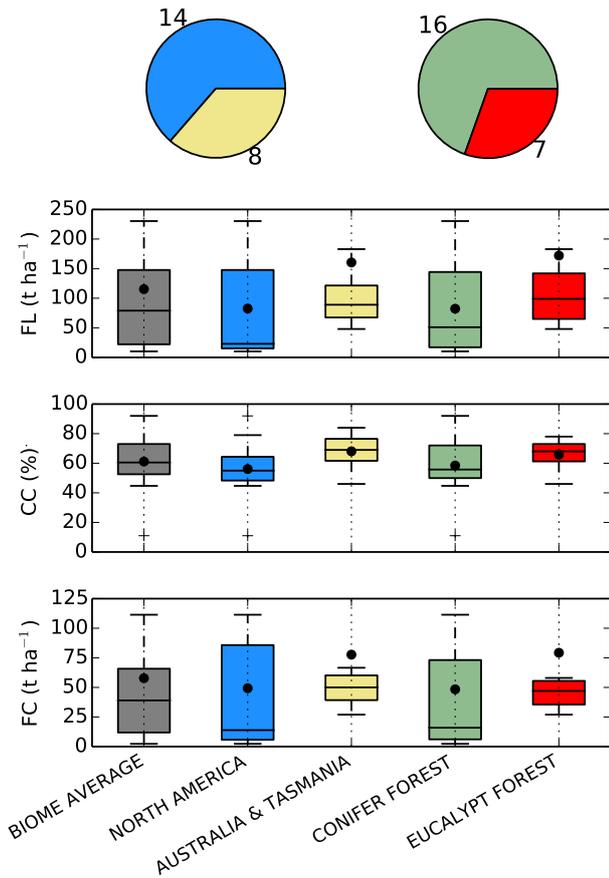
The biome-averaged FL for the temperate forest biome was  $115 \pm 144 \text{ t ha}^{-1}$ , the CC equaled  $61 \pm 18 \%$ , and fuel consumed by the fire was  $58 \pm 72 \text{ t ha}^{-1}$ . Note that we focused on all the measurements presented in Table 1c, so studies that provide information on one specific fuel class only (e.g., ground fuels, Goodrick et al., 2010) were also included to calculate biome-averaged values. Although CC for North America, Australia and Tasmania were comparable ( $\sim 60 \%$ ), the FC showed lower values for North America ( $49 \pm 62 \text{ t ha}^{-1}$ ) than Australia and Tasmania ( $78 \pm 91 \text{ t ha}^{-1}$ ). One possible cause of this discrepancy



**Figure 3.** Overview of field measurements of fuel load (FL), combustion completeness (CC), and fuel consumption (FC) in the tropical forest biome. The pie charts on top correspond to the number of unique measurement locations for different geographical regions (left) and forest types (right), and in the box plots below field averages of FL, CC, and FC are presented. The boxes extend from the lower to upper quartile values of the measurement data, with a line at the median and a black filled circle at the mean. The whiskers extend from the box to show the range of the data, and outliers are indicated with pluses.

is the contribution of different vegetation types, as elaborated in Fig. 4. Measurements in North America were mainly conducted in conifer forest, while eucalypt was the more dominant forest type for Australia and Tasmania. FC for both forest types compares fairly well with the regional averages found, and equaled  $48 \pm 58 \text{ t ha}^{-1}$  for conifers and  $79 \pm 98 \text{ t ha}^{-1}$  for eucalypt forest.

Table 2c shows that litter in the temperate forest had a higher FL and FC than in the tropical forest biome, and the average FC for this surface fuel category equaled  $17 \pm 9.9 \text{ t ha}^{-1}$ . The different woody debris classes had a similar pattern to that found for the savanna and tropical forest biome, with decreasing CC for categories with increasing fuel diameters. However, an interesting difference was found in the biggest size class: sound woody debris had a low CC ( $38 \pm 42 \%$ ), while the fraction of rotten woody

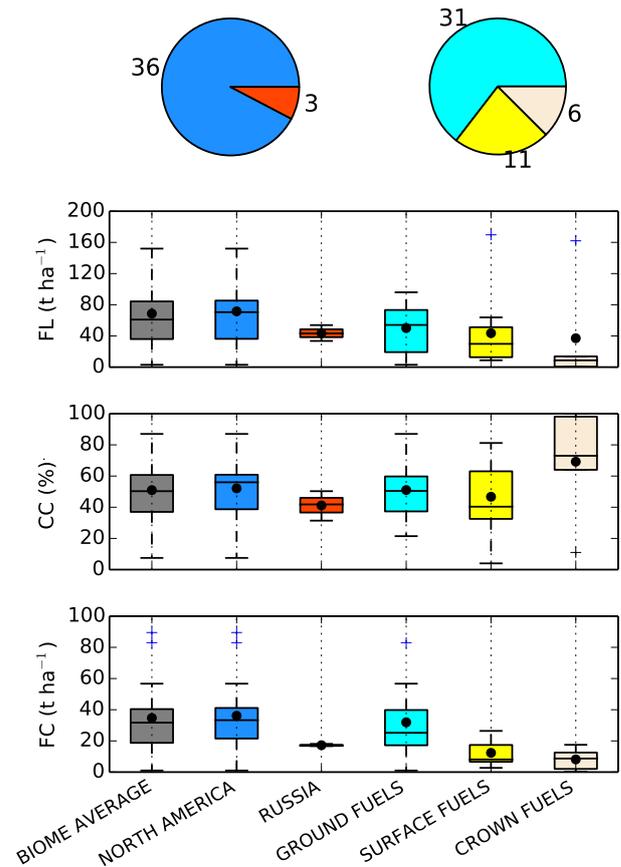


**Figure 4.** Overview of field measurements of fuel load (FL), combustion completeness (CC), and fuel consumption (FC) in the temperate forest biome. The pie charts on top correspond to the number of unique measurement locations for different geographical regions (left) and forest types (right), and in the box plots below field averages of FL, CC, and FC are presented. The boxes extend from the lower to upper quartile values of the measurement data, with a line at the median and a black filled circle at the mean. The whiskers extend from the box to show the range of the data, and outliers are indicated with pluses.

debris consumed by the fire was very high ( $96 \pm 5.4\%$ ), resulting in an average FC of  $20 \pm 4.8 \text{ t ha}^{-1}$  for this category. Although this difference was observed in a few other studies as well, little research is available on comparing the physical and chemical properties of sound and rotten woody debris, which is likely to affect the FC (Hyde et al., 2011). The most important fuel category from an FC perspective was organic soil, with an average value of  $25 \pm 31 \text{ t ha}^{-1}$ . For the same reason as explained in Sect. 2.1, a small discrepancy was found between the total FC sum of different fuel categories ( $77 \text{ t ha}^{-1}$ ) and the biome average ( $58 \pm 72 \text{ t ha}^{-1}$ ).

## 2.4 Boreal forest

Fires in the boreal (high latitudes of about  $50$  to  $70^\circ$ ) forest are thought to be mostly natural (wildfires) due to the vast



**Figure 5.** Overview of field measurements of fuel load (FL), combustion completeness (CC), and fuel consumption (FC) in the boreal forest biome. The pie charts on top correspond to the number of unique measurement locations for different geographical regions (left) and fuel classes (right), and in the box plots below field averages of FL, CC, and FC are presented. The boxes extend from the lower to upper quartile values of the measurement data, with a line at the median and a black filled circle at the mean. The whiskers extend from the box to show the range of the data, and outliers are indicated with pluses.

size of the forest region, the low population densities and the difficult accessibility. However, much of the Asian boreal forests are disturbed by (il)legal logging activities (Vandergert and Newel, 2003), which can increase fire activity in more remote regions (Mollicone et al., 2006). Approximately two thirds of the boreal forests are located in northern Eurasia, while the remainder is in North America. The circumpolar boreal fire regime is characterized by large forest fires, although fires in North America are in general larger and less frequent than the ones in Eurasia (de Groot et al., 2013a). North American boreal fires are characterized by high intensity crown fires, while fires in boreal Russia are more often surface fires of lower intensity (Amiro et al., 2001; Soja et al., 2004; Wooster et al., 2004; de Groot et al., 2013a). Canada has a very long fire record, starting in 1959, while the record

for Alaska starts in 1950 (Kasischke et al., 2002). Since 1990, 2.65 million ha year<sup>-1</sup> burned in the North American boreal forest, with high year-to-year variability (Kasischke et al., 2011). FL in the boreal forests depends to a large extent on tree species, stand density, climate, topography, moisture, seasonal thawing of permafrost and the time since the last burn. In many forest types, dead material accumulates in deep organic soil horizons due to the slow decomposition rates. CC in organic soils is mostly controlled by conditions that control surface soil moisture, including topography, seasonal thawing of permafrost, and antecedent weather conditions. When dry conditions prevail, such as during a high-pressure blocking event that can last for few days to several weeks over North America (Nash and Johnson, 1996), much of the forest floor can burn, and depths of 30 cm or more can be reached. There is a strong relation between moisture content and fuel bed depth on the one hand and forest floor consumption on the other hand (e.g., de Groot et al., 2009). Of all global fire regimes, the boreal forest is most susceptible to climate change due to polar amplification of temperature increase (Flannigan et al., 2013; de Groot et al. 2013b). For example, the area burned by lightning fires in the North American boreal region doubled between 1960 and 1990 (Kasischke and Turetsky, 2006).

Field measurements described in the literature were almost all conducted in boreal North America (35 in total), except for three measurement sets that came from boreal Asia (Fig. 1, Table 1d). The general method for determining FL and FC was to apply the PIM. Approaches have also been developed to estimate consumption of surface organic layer fuels by estimating the pre- and post-fire thicknesses and densities of surface organic horizons (de Groot et al., 2009; Turetsky et al., 2011).

We estimated a biome-averaged FL of  $69 \pm 61 \text{ t ha}^{-1}$ , substantially lower than the average FL for the temperate forests. The average FL for this biome is for upland forest types. However, deep peatland deposits (see Sect. 2.10) cover about 107 M ha (Zoltai et al., 1998) or 18 % of the North American boreal forest zone (Brandt, 2009) and 16 % of the northern circumpolar permafrost soil area (Tarnocai et al., 2009). By contrast, peatlands only cover about 0.07 M ha in the temperate zone, which has higher FL overall. Despite low decomposition rates due to a cold, moist climate, the lower FL in the boreal forest region is primarily a result of slower tree growth rates (biomass accumulation) and frequent to infrequent fire disturbance that can remove substantial amounts of fuel. The average CC was  $51 \pm 17 \%$ , and the FC equaled  $35 \pm 24 \text{ t ha}^{-1}$ . Similar as for the temperate forest, we included all measurements (presented in Table 1d) to calculate the biome-averaged values. The representativeness of these values for wildfires and prescribed fires is discussed in Sect. 3.2. Differences between boreal North America and Siberia were observed, but it should be noted that only three studies provided an FC estimate for Russia. Values on FL,

CC, and FC were overall higher for boreal fires in North America than the field studies in Russia (Fig. 5).

Information on fuel categories is presented in Table 2d, as well as in Fig. 5. Different classification systems were sometimes used for boreal fuels, and therefore it was difficult to extract the right information for ground, surface and crown fuels (further discussed in Sect. 3.4). Moreover, it was not always clear in which class certain fuels are consumed: e.g., organic material can be consumed on the ground but also in a crown fire (Hille and Stephens, 2005). The highest FL ( $50 \pm 29 \text{ t ha}^{-1}$ ) and FC ( $32 \pm 26 \text{ t ha}^{-1}$ ) in the boreal forest biome was found for ground fuels, mainly consisting of organic soils. Furthermore, a difference in organic matter FL in permafrost and non-permafrost regions was found ( $56$  and  $86 \text{ t ha}^{-1}$ , respectively). However, due to a CC of 62 and 41 % for permafrost and non-permafrost regions, the FC for both regions was equal ( $35 \text{ t ha}^{-1}$ ). Finally, slope aspect has been shown to have an effect as well, with the south facing slopes having the highest FL and FC due to warmer and drier conditions that better favour plant growth and fire intensity than shadowed north faces (Viereck et al., 1986; Turetsky et al., 2011). As with most of our findings, however, the number of studies is far too low to evaluate whether this is also the case in general.

## 2.5 Pasture

Fires related to agricultural practices were divided into shifting cultivation (Sect. 2.6), the burning of crop residues (Sect. 2.7) and pasture burning. The latter type of burning often follows tropical deforestation fires and is used to convert land into pasture. Prior to this conversion, lands can be used in shifting cultivation as well. Typically, landowners set fires every 2–3 years to prevent re-establishment of forests (Kauffman et al., 1998) and to enhance the growth of certain grasses (Fearnside, 1992). In general, these fires mostly consume grass and residual wood from the original forest. Pasture fires are most common in the Brazilian Amazon where many cattle ranches have been established in areas that were previously tropical forest. Although less abundant, these “maintenance” fires also occur in tropical regions of Africa, central America and Asia.

The pasture measurements presented in Table 1e represent five unique measurement locations and cover two different continents (Fig. 1). Pasture had an average FL, CC, and FC of  $74 \pm 34 \text{ t ha}^{-1}$ ,  $47 \pm 27 \%$ , and  $28 \pm 9.3 \text{ t ha}^{-1}$ , respectively. Regional discrepancies for FC were found though, with FL for Brazilian pastures ( $84 \pm 29 \text{ t ha}^{-1}$ ) being substantially higher than those found in Mexico ( $35 \text{ t ha}^{-1}$ ). However, FC values compared reasonably well for both regions ( $30 \pm 10$  and  $24 \text{ t ha}^{-1}$  for Brazil and Mexico, respectively).

## 2.6 Shifting cultivation

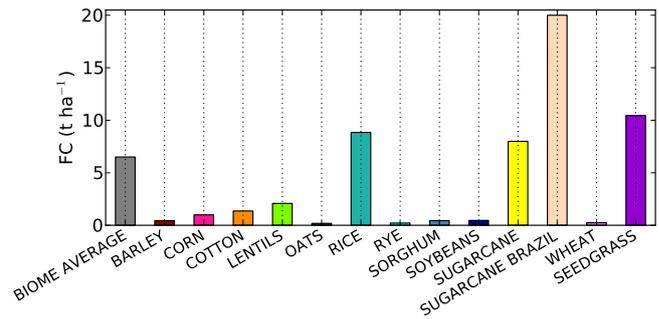
Shifting cultivation is commonly practiced in Africa, central America, South America and Asia. In general, lands are cultivated temporarily (often for only a few years) before soil fertility is exhausted or weed growth overwhelms the crops. The lands are then abandoned and may revert to their natural vegetation, while the farmers move on to clear a new fields elsewhere. The land is slashed and burned, which leaves only stumps and large trees in the field after the fire (Stromgaard, 1985). Apart from the fact that fire is an easy and cheap tool to clear the land, it has the further advantage that the ashes will also (temporarily) enrich the soil.

For shifting cultivation fires the average FL was 44 with a range of 14–75 t ha<sup>-1</sup>, the CC equaled 47 [30–64] %, and FC was 23 [4–43] t ha<sup>-1</sup>. Note that these values are based on the measurements of two studies only (Fig. 1, Table 1f). The two shifting cultivation studies showed a remarkable difference: FC of Indian tropical dry deciduous forest (4.0 t ha<sup>-1</sup>; Prasad et al., 2000) was 1 order of magnitude lower than for shifting cultivation practices in the wooded savanna of Zambia (43 t ha<sup>-1</sup>; Stromgaard, 1985). Due to the relatively small number of measurements, these findings are not conclusive.

## 2.7 Crop residue

Crop residue burning is a common practice to recycle nutrients, control pests, diseases, and weeds and, in general, to prepare fields for planting and harvesting. The main crop residue types that burn are rice, grains (i.e., wheat) and sugarcane, but burning is not limited to these crop types. FL is highly variable, as it depends on both the type of crop burned and the method used for harvesting the crop (mechanized, manual, etc.). Detecting these fires using global burned area products is difficult, as in general, cropland fires are small and the land can be tilled and replanted quickly after burning (making it difficult to observe the latency of burned ground as is common in less managed and/or more natural landscapes). Moreover, the fuel geometry varies globally from short-lived burning of loose residue in the field to long-lasting smoldering combustion of small hand piles of residue, and both are hard to detect from space. Traditional methods of obtaining estimates for agricultural fires are the use of governmental statistics on crop yield (e.g., Yevich and Logan, 2003), residue usage for cooking and livestock (the leftovers are assumed to be burned), field measurements, or the use of agronomic data (e.g., Jenkins et al., 1992).

On average, crop residue burning had FL of  $8.3 \pm 9.9$  t ha<sup>-1</sup>, CC of  $75 \pm 21$  % and FC of  $6.5 \pm 9.0$  t ha<sup>-1</sup> (Table 1g). We estimated an average FL of 23 t ha<sup>-1</sup> for Brazilian sugarcane (Lara et al., 2005) by using a CC of 88 % as reported by McCarty et al. (2011). FC values for different US crop types (McCarty et al., 2011) were used to derive crop-specific FL data (French et al., 2013) and CC values were taken from expert knowledge from agriculture



**Figure 6.** Fuel consumption (FC) for different US crop types as reported by McCarty et al. (2011), and Brazilian sugarcane (Lara et al., 2005). The grey bar corresponds to the biome-averaged FC value for crop residue burning as presented in this study.

extension agents in Arkansas, Louisiana, Florida, Kansas, and Washington during field campaigns in 2004, 2005, and 2006, as well as from the scientific literature (Dennis et al., 2002; Johnston and Golob, 2004). CC variables ranged from 65 % for cotton and sugarcane and 85 % for wheat and bluegrass, which are lower but within the range of the CC value (−23 to −3 % less than CC of 88 %) used by the Environmental Protection Agency (EPA) of 88 % (EPA 2008 GHG).

FC values varied between different crop types, as shown in Fig. 6. For US crops the highest FC was found for seedgrass (10 t ha<sup>-1</sup>) and rice (8.8 t ha<sup>-1</sup>), while values for soybeans (0.5 t ha<sup>-1</sup>) and corn (1.0 t ha<sup>-1</sup>) were lower. In general, US crop values are assumed in the analysis to be approximately representative of other developed agricultural areas like Brazil and Russia (McCarty et al., 2012), but uncertainty increases for less industrialized agricultural areas in Africa and Asia. However, Brazilian sugarcane (20 t ha<sup>-1</sup>) was found to have an FC that is more than twice as high as sugarcane in the US (8.0 t ha<sup>-1</sup>). More measurements are needed to confirm this discrepancy.

## 2.8 Chaparral

Chaparral vegetation is a type of shrubland that is primarily found in the southwestern US and in the northern portion of Baja California (Mexico), but similar plant communities are found in other Mediterranean climate regions around the world like Europe, Australia and South Africa. Typically, the Mediterranean climate is characterized by a moderate winter and dry summer, which makes the chaparral biome most vulnerable to fires in summer and fall (Jin et al., 2014). In California, the combination of human ignition, the large wildland–urban interface, and extreme fire weather characterized by high temperatures, low humidities, and high offshore Santa Ana winds (Moritz et al., 2010) may lead to large and costly wildfires (Keeley et al., 2009).

We found two studies covering five different measurement locations in the southwestern US (Table 1, Fig. 1h). Since

Cofer III et al. (1988) only provided an FC for chaparral burning, we used a CC of 76 % (average CC from the studies of Hardy et al. (1996) and Yokelson et al., 2013) to derive an FL estimate for the Cofer et al. (1988) study. We then used the FL values of all three studies to estimate the biome-averaged FL of  $40 \pm 23 \text{ t ha}^{-1}$ . The CC equaled 76 %, yielding an average FC of  $27 \pm 19 \text{ t ha}^{-1}$ .

## 2.9 Tropical peat

Tropical peatland has only recently been recognized as an important source of biomass burning emissions. Roughly 60 % of the worldwide tropical peatland is located in Southeast Asia, and more specifically in Indonesia (Rieley et al., 1996; Page et al., 2007). Peat depth is an indicator for the total biomass stored in peatland, but only the peat layer above the water table can burn. Drainage and droughts lower the water table, adding to the total FL. On top of that, living biomass and dead above ground organic matter also contribute to the FLs in these peatlands. The bulk density and carbon content of peat are of importance to determine the amount of carbon stored. The average density is around  $0.1 \text{ g cm}^{-3}$  and the carbon content ranges between 54–60 % (Page et al., 2002; Riely et al., 2008; Ballhorn et al., 2009; Stockwell et al., 2014). The depth of burning is the key factor that determines the total FC, but information about it is scarce. Results from several field measurements indicate a link between this burning depth and the depth of drainage (Ballhorn et al., 2009). Commercial logging in drained peat swamps has increased their susceptibility to fire, especially during droughts (such as during an ENSO event).

In total four studies provided data on tropical peatland measurements in Indonesia (Table 1i). In general, post-fire observations of the average burn depth were combined with pre-fire conditions reconstructed from adjacent unburned patches to determine the FC.

Tropical peatland (including peat soils and overstory) had the highest FC of all biomes, with an average of  $314 \pm 196 \text{ t ha}^{-1}$ . Only two studies provided data on FL and CC, and since the study of Saharjo and Nurhayati (2006) focused on litter and branches only, a CC of 27 % (Usup et al., 2004) was found to be representative for the tropical peat biome. Taking a CC of 27 %, the biome-averaged FL equaled  $1056 \pm 876 \text{ t ha}^{-1}$ , thereby having the highest FL of all biomes. However, due to limited information on CC measured in the field there is no clear definition of the average FL for tropical peat. Note that the measurements taken by Ballhorn et al. (2009) were using Laser Imaging, Detection And Ranging (LIDAR) aerial remote sensing, and the study of Page et al. (2002) relied on field measurements combined with information obtained from Landsat Thematic Mapper (TM) images.

## 2.10 Boreal peat

The northern peatlands are a result of the slow decomposition of organic material over thousands of years. Traditionally, northern peatlands have been considered a slow, continuous carbon sink. However, the vulnerability of this region to global warming and the resulting increase in wildland fires have challenged this idea (Zoltai et al., 1998; Harden et al., 2000; Turetsky, 2002). There are still large uncertainties associated with the FL and CC of peat fires. The depth of fires is not well documented, leading to large uncertainties in the total FC estimates. In some cases, water table depth may serve as a proxy for determining the depth of burning. However, the susceptibility of peatlands to fire under different moisture conditions is also poorly documented at best. This makes modeling peat fires very difficult and stresses the importance of more field measurements.

Two measurements were taken between 1999 and 2001 in boreal Canada (Table 1j). On each burn site, multiple plots were established and information on the peat density (which is assumed to increase nonlinearly with depth) was used in combination with the burn depth to determine the FC. No data on FL and CC were provided, but the average FC of the two studies is  $43 [42\text{--}43] \text{ t ha}^{-1}$ . A standard deviation of  $25 \text{ t ha}^{-1}$  (Turetsky and Wieder, 2001) can be used as the average uncertainty for the boreal peat biome. Turetsky and Wieder (2001) showed that FC of permafrost bogs ( $58 [43\text{--}72] \text{ t ha}^{-1}$ ) is more than twice as high as that of continental bogs ( $27 [11\text{--}42] \text{ t ha}^{-1}$ ). A similar difference was found for hummocks and hollows, which are raised peat bogs and lows, respectively: FC for hummocks was  $29 \pm 2.0 \text{ t ha}^{-1}$ , while fires in hollows consumed on average  $56 \pm 6.0 \text{ t ha}^{-1}$  (Benscoter and Wieder, 2003).

## 2.11 Tundra

The Arctic tundra stores large amounts of carbon in its organic soil layers that insulate and maintain permafrost soils, although these soil layers are shallower than those found in peatlands and boreal forests. While the region is treeless, some vegetation types include a substantial shrub component where additional carbon is available for burning. On Alaska's North Slope approximately 10 % of the land cover is shrub dominated ( $> 50 \%$  shrub cover), while the remainder is dominated by herbaceous vegetation types (Raynolds et al., 2006). Fire regime in the Arctic is largely unknown, but historically fire is generally absent in the tundra biome compared to other biomes. However, the evidence of increasing fire frequency and larger extent of the fires in the arctic (Hu et al., 2010) may represent a positive feedback effect of global warming, so in the future more fires may occur in this biome (Higuera et al., 2011). There are still large unknowns of the impacts that fires have on the carbon stocks of the tundra ecosystems. Even the topsoil layers in the tundra store large pools of carbon in organic-rich material. This removal

of the topsoil may also expose the permafrost layers to heating by the warm summer temperatures, thawing the ground and destabilizing the tundra carbon balance.

The only measurements found in the literature of FC in the tundra biome are from the Anaktuvuk River fire in 2007 (Mack et al., 2011). The measurements were taken on twenty sites in the burned area and the pre-fire peat layer depth was reconstructed to determine the pre-fire FL. The FL was on average  $165 \pm 15 \text{ t ha}^{-1}$ , and averaged CC and total FC was respectively  $24 \pm 5.0 \%$  and  $40 \pm 9.0 \text{ t ha}^{-1}$  (Table 1k). These measurements represent a thorough effort to document FC, but still represent just one fire that is considered to be a fairly high severity event (Jones et al., 2009). Other measurements of surface FC at fires in the Noatak region of Alaska and a recent burn on the Alaskan North Slope showed minimal organic surface material loss (N. French, unpublished data). These fires may represent more typical fire events with more moderate consumption than was found in the Anaktuvuk River fire. There is no doubt that the lack of sufficient field measurements in the tundra biome means a reasonable estimate of FC in tundra fires is not fully known. While the Anaktuvuk River fire measurements are of value, there should be caution in using these data to generalize since the event represents a more severe event than many fires in the region. They may, however, be indicative of how future fires in the region may impact carbon losses as the region experiences increased fire frequency and severity.

### 3 Discussion

#### 3.1 Spatial representativeness of fuel consumption measured in the field

Due to the spatial heterogeneity in fuels and the limited number of measurements, one important question to ask is how representative the biome-averaged values presented in this review are. Field measurements of FC were spatially well represented in the major biomass-burning regions, like the Brazilian Amazon, boreal North America and the savanna areas in southern Africa. However, several other regions that are important from a fire emissions perspective were lacking any measurements, and these include central Africa (e.g., Congo and Angola, but also regions further north such as Chad and southern Sudan), Southeast Asia and eastern Siberia (Fig. 1). Due to these spatial gaps, it remains uncertain whether measurements of FL, CC, and FC as presented in this study are representative for the whole biome. As mentioned for the “Tundra”, where fire may become increasingly important as the region warms, the one set of field samples included in this review may not be representative of past and future fire.

Within biomes differences were found to be large for certain regions, as shown in Figs. 2–5. For example, we found substantial differences in FL and FC for boreal areas, with

Russian sites having lower values compared to the ones in North America (Fig. 5). This difference might be due to different burning conditions in both regions, with a larger contribution of surface fuels and less high-intensity crown fires occurring in boreal Russia (Wooster et al., 2004). Although available literature data showed that FC for crown fuels was indeed higher than for surface fuels, more data for especially boreal Russia is needed to confirm this line of thought. Moreover, Boby et al. (2010) and Turetsky et al. (2011) showed that the timing of FC measurements (early dry seasons versus late dry season) contribute to different boreal FC values as well. In general, both FC and CC may increase over the course of the dry season as large diameter fuels dry out. This was also suggested by Akagi et al. (2011) for the savanna biome, and is consistent with the seasonal decrease in MCE as proposed by Eck et al. (2013).

Regional differences were also found for the tropical forest biome, where almost all measurements were conducted in the Brazilian Amazon, with a few exceptions for Mexico and Indonesia. Southeast Asia (Myanmar, Vietnam, Laos, and Cambodia) was lacking any FC measurements described in the peer-reviewed literature, but this region is important from a fire emissions perspective. Tropical forests in Mexico had a higher CC than forests in the Amazon and Indonesia (Fig. 3), and had a higher FC as well. Different forest types can likely explain this difference; in Fig. 3 substantial differences are shown for FL, CC, and FC in primary tropical evergreen forest, secondary tropical evergreen forest, and tropical dry forest. Obviously, the number of measurements conducted in a specific forest type will impact the biome-averaged value found for a certain region. Clearly, the definition of a certain biome is not always straightforward, and uncertainty regarding regional discrepancies within the different biomes should be taken into account when averaged values are interpreted and used by the modeling communities.

Coming back to the question posed in the beginning of this section, we think care should be taken with using biome-averaged values. They provide a guideline, but the path forward is to continue developing models or remote sensing options that aim to account for variability within biomes, and to use the database accompanying this paper to constrain these models, rather than simply to use biome-averaged values (further discussed in Sect. 3.2). Use of FC for specific vegetation types within broader biomes (like the different crop types as presented in Fig. 6) or fuel categories offers an interesting alternative and is further discussed in Sect. 3.4.

#### 3.2 Field measurement averages and comparison with GFED3

Although the definition of a certain biome is not always straightforward, the biome-averaged values that we presented in this paper are still valuable for highlighting differences in fire characteristics between regions with

specific vegetation and climate characteristics. We compared our work with estimates from the Global Fire Emissions Database version 3 (GFED3) and several FRE derived studies (Sect. 3.3). GFED3 fire emissions estimates (monthly  $0.5^\circ \times 0.5^\circ$  fields) are based on estimates of burned area (Giglio et al., 2010) and the satellite-driven Carnegie–Ames–Stanford Approach (CASA) biogeochemical model (van der Werf et al., 2010). To calculate FC we divided the GFED3 total biome-specific emissions estimates (g dry matter) in every modeling grid cell by the total burned area observed for every grid cell. Since one grid cell may consist of multiple biomes we followed the GFED3 fractionation of emissions estimates, which represents the contribution of a certain biome to total emissions within one grid cell. Biome-specific information on the area burned within one grid cell was not available, and therefore we assumed that burned area followed the same fractionation as the GFED3 emissions estimates. This assumption may overestimate or underestimate biome-averaged GFED3 FC values: for example, in a deforestation grid cell that consists of savannas and tropical evergreen forests, the contribution of savanna fire emissions to total emissions can be small, even when the contribution of savanna burned area to total burned area observed in a grid cell is actually quite large. In this specific case – when assuming that burned area followed the same fractionation as the emissions – the estimated FC of savannas would be overestimated.

In Table 3 an overview is given for biome-specific FL, CC, and FC that we estimated from data found in the literature. In the fifth column FC per unit burned area of GFED3 is shown for the collocated grid cells, i.e., grid cells in which measurements were taken, and the sixth column presents the difference between GFED3 FC and the field measurements. In general, substantial differences were found between collocated GFED3 FC and the field measurements. Although the average FC agreed reasonably well ( $< 40\%$ ) for crop residue, tropical peat and the boreal peat biome, much larger discrepancies ( $> 59\%$ ) were found for the other biomes. Many field measurements for these biomes had a standard deviation that was close to the measurement average, indicating that the uncertainty is substantial.

Within the savanna biome GFED3 overestimated the FC by 72 % compared to the measurements, and this overestimation was even higher for grassland regions (78 %). A possible cause of these discrepancies is that field campaigns tend to focus on frequently burning areas, so fuels do not have the time to build up and increase their FL (van der Werf et al., 2010). Because of the relatively coarse  $0.5^\circ$  resolution of GFED3, the fire frequency in GFED is the average of more and less frequently burning patches and thus potentially longer than in field sampling sites. On the other hand, only a very small portion of the land's surface burns annually (van der Werf et al., 2013). Improved resolution for the models may help to alleviate this problem and bring model

values closer to the field measurements, although it is very unlikely that this is the only reason for the noted discrepancy.

For tropical forests, an important biome due to large-scale deforestation emissions, substantial differences were found as well: GFED3 overestimated FC by 71 % compared to the field measurement average for collocated grid cells. This discrepancy may be partly explained by the fact that repeated fires in the tropical forest domain (when forest slash that did not burn in a first fire is subject to additional fires during the same dry season) are not always included in the field measurements. Within GFED3, on the other hand, these repeated fires were modeled by the number of active fires observed in the same grid cell (fire persistence), which yields information on the fuel load and type of burning (Morton et al., 2008; van der Werf et al., 2010). Regional differences within the biome, as discussed in Sect. 3.1, will also contribute to the differences found: in our case, the field measurement average was biased towards evergreen tropical forest fires, but when the emphasis is put on fires in secondary or tropical dry forests this average value could change significantly (Fig. 3). It is likely that grid cell heterogeneity in tropical deforestation regions explains the large discrepancy found for the pasture biome, where GFED3 FC overestimated the field measurements by almost 500 %. For these specific pasture grid cells, GFED3 may have been biased towards tropical evergreen deforestation fires, thereby increasing the average FC.

In the temperate forest biome FC was underestimated in GFED3 by 74 % compared to the field measurement average for collocated grid cells. In our averaged field measurement estimate we included all measurements presented in Table 1c. As noticed in Sect. 2.3, it is likely though that studies that provided a total FC (i.e., the FC of ground, surface and/or crown fuels) are more representative of wildfires. Prescribed burns, on the other hand, tend to burn less fuel and therefore the studies that only include ground or surface fuels were probably more representative for this fire type. When focusing on studies that provide information on one specific fuel class only, the field average for the temperate forest would be significantly lower ( $13 \pm 12 \text{ t ha}^{-1}$ ) as well as the discrepancy with GFED3 (+14 %). This FC value of  $13 \text{ t ha}^{-1}$  may be more realistic for prescribed fires, which contribute to roughly 50 % of all temperate forest fire emissions in the contiguous United States (CONUS). Still, it remains very uncertain how well FC measured for specific fuel classes is representative for prescribed fires and wildfires. This issue also counts for boreal forests, where GFED3 overestimated the field measurements by almost 80 %. When only including studies that provided a total FC (i.e., the FC of ground, surface and/or crown fuels), the field average for the boreal forest would increase from  $35 \pm 24$  to  $39 \pm 19 \text{ t ha}^{-1}$  and the discrepancy with GFED3 would decrease (from +79 to +60 %). This value of  $39 \pm 19 \text{ t ha}^{-1}$  may be more representative of boreal wildfires. Note that for temperate and boreal forest measurements sometimes the more restrictive definition of FL (as presented in Sect. 1) was used, and this

**Table 3.** Biome-averaged values for fuel load (FL), combustion completeness (CC), and fuel consumption (FC) field measurements. Column 5 shows the FC per unit burned area as used in GFED3 ( $FC_{GFED3}$ ) and in column 6 the difference (%) in  $FC_{GFED3}$  compared to the average FC of field measurements is given. Standard deviation (SD) is shown in parentheses.

Biome	FL ( $t\ ha^{-1}$ )	CC (%)	FC ( $t\ ha^{-1}$ ) <sup>a</sup>	$FC_{GFED3}$ ( $t\ ha^{-1}$ ) <sup>b</sup>	Difference (%) <sup>c</sup>
Savanna	7.6 (6.5)	71 (26)	4.6 (2.2)	7.9	+72
Grassland savanna	5.3 (2.0)	81 (16)	4.3 (2.2)	7.7	+78
Wooded savanna	11 (9.1)	58 (32)	5.1 (2.2)	8.1	+59
Tropical forest	285 (137)	49 (22)	126 (77)	215	+71
Temperate forest	115 (144)	61 (18)	58 (72)	15	-74
Boreal forest	69 (61)	51 (17)	35 (24)	62	+79
Pasture	74 (34)	47 (27)	28 (9.3)	168	+491
Shifting cultivation	44 (-)	47 (-)	23 (-)	6.5	-72
Crop residue	8.3 (9.9) <sup>d</sup>	75 (21)	6.5 (9.0)	5.6	-13
Chaparral	35 (23) <sup>e</sup>	76 (6.2)	27 (19)	3.5	-87
Tropical peatland	1056 (876) <sup>f</sup>	27 (-)	314 (196)	228	-27
Boreal peatland	-	-	42 (-)	25	-40
Tundra <sup>g</sup>	165 (15)	24 (5.0)	40 (-)	-	-

<sup>a</sup> For biomes where only one or two measurements are available, no uncertainty estimate is given.

<sup>b</sup> FC per unit area burned according to GFED3, averaged over 1997–2009. The number represents the FC rate for the collocated grid cells, i.e., grid cells in which field measurements were taken. Note that for this calculation the assumption was made that GFED burned area is equally divided over different fire types in one grid cell, which may influence average  $FC_{GFED3}$  values.

<sup>c</sup>  $FC_{GFED3}$  compared to the average FC of field measurements for collocated grid cells. Positive numbers indicate that  $FC_{GFED3}$  is higher than the average FC of field measurements.

<sup>d</sup> We assumed an average CC of 88% as reported in McCarty et al. (2011) to estimate FL for the study of Lara et al. (2005).

<sup>e</sup> We assumed a CC of 76% (average CC for the studies of Hardy et al., 1996 and Yokelson et al., 2013) to estimate FL for the study of Cofer III et al. (1988).

<sup>f</sup> We assumed an average CC of 27.2%, as reported in Usup et al. (2004), to estimate FL for the studies of Page et al. (2002) and Ballhorn et al. (2009).

<sup>g</sup> For the measurement location in the tundra biome no area burned was detected by GFED, and therefore no comparison with GFED3 estimates was made.

can have an impact on FC values as well; if one applies a CC calculated with respect to a restrictive pre-fire FL to the total biomass available, the overall FC that was estimated can be too high.

For most biomes, a few field measurements had an FC that was 1 order of magnitude larger than the other values listed in Table 1, which explains the discrepancy between the median and average FC values that was sometimes found (e.g., the “Australia and Tasmania” region in Fig. 4). By neglecting these “outliers” the biome-averaged values may change significantly, but this could lead to erroneously low or high estimates as well. In general, FC shows large variability between biomes, within biomes, and even within a specific fuel type. FC is often hard to measure, and since only a few measurements are available for some biomes, care should be taken when using the biome-averaged values presented in this paper.

### 3.3 Field measurement averages and comparison with FRE derived FC estimates

Besides a comparison with GFED3 data, we performed a comparison of field measurement averages with fire radiative energy (FRE, time-integrated FRP) derived estimates

as well. The basis of the FRE approach for estimating FC is that the heat content of vegetation is more or less constant, and that the FRE released and observed through a sensor can be converted to FC by the use of a constant factor, which was found to be  $0.368 \pm 0.015\ kg\ MJ^{-1}$  across of a range of fuels burned under laboratory conditions (Wooster et al., 2005). More recent field experiments, however, indicated that the conversion factor might be slightly lower for grasslands in North America (Kumar et al., 2011; Schroeder et al., 2014). Smith et al. (2013) investigated the relationship between FC and FRE for pine needles with different fuel moisture contents, and found that FRE released per kilogram biomass consumed decreased with fuel moisture content due to the energy required to evaporate and desorb the water contained in the fuel. Thus, corrections for FRE based FC assessments may be needed for fuels that burn at higher fuel moisture contents. Differences in heat content of fuel may introduce additional variation: for example, a clear relationship between FRE and FC has not yet been demonstrated for fires that burn mostly in the smoldering stage, like organic soils in boreal forests or large woody debris and trunks in tropical deforestation regions. Another potential source of uncertainty in the relation between satellite-derived FRE and

FC is the correction for atmospheric disturbances, which may significantly alter FRP retrievals and hence estimates of FC (Schroeder et al., 2014). Note that, currently, atmospheric correction is not performed for the standard fire products derived from MODIS. Moreover, Schroeder et al. (2014) also indicate that cloud masking in the MODIS FRP product may lead to FRP underestimates as hotspots under thick smoke may erroneously be masked out.

Despite all these uncertainties this approach is promising and there are a number of studies that relate FRE to FC on regional (Roberts et al., 2011; Freeborn et al., 2011) to global scales (Vermote et al., 2009; Ellicott et al., 2009), and Kaiser et al. (2012) used FRE to represent biomass burning in an operational chemical weather forecast framework. However, since such estimates can be derived independently of burned area, only a limited number of studies allow a straightforward comparison to the FC values given in mass units per area burned from the field experiments used in this study.

A common finding of FRE-based estimates is that FC is generally lower than GFED estimates, as shown by Roberts et al. (2011) who estimated FC for Africa through an integration of MODIS burned area and Meteosat Spinning Enhanced Visible and Infrared Imager (SEVIRI) derived FRP and found values that were about 35 % lower than GFED. For the savanna biome a median FC of  $\sim 4 \text{ t ha}^{-1}$  was found for grassland and shrubland. This corresponds relatively well to the mean of  $4.3 \pm 2.2$  and  $5.1 \pm 2.2 \text{ t ha}^{-1}$  found in the grassland savanna and wooded savanna field studies we compiled, respectively. Boschetti and Roy (2009) explored temporal integration and spatial extrapolation strategies for fusing MODIS FRP and MODIS burned area data over a single large fire in a grassland dominated area with sparse eucalypt trees in northern Australia. They estimated an FC range of  $3.97\text{--}4.13 \text{ t ha}^{-1}$ , which is well within the range found in the Australian FC studies summarized in Table 1. Kumar et al. (2011) exploited properties of the power law distribution to estimate FC from FRP for an Australian savanna and a study area in the Brazilian Amazon. While their FC estimate of  $4.6 \text{ t ha}^{-1}$  of the Australian site is similar to the temporal integration results of Boschetti and Roy (2009), the estimate for the Brazilian site is above  $250 \text{ t ha}^{-1}$  and thus substantially higher than the biome-averaged value for Brazilian tropical forest ( $117 \pm 56 \text{ t ha}^{-1}$ ).

In general, realistic values are often obtained for well-observed fires, but unrealistically low or high values can often occur especially for smaller fires due to the sparseness of FRP observations and inaccuracies in the temporal interpolation and the burned area estimates. While FRE seems to provide realistic estimates under a range of conditions, issues of undersampling of FRE and – maybe less importantly – the conversion of FRE to FC still remain to be addressed more completely in order to derive spatially explicit FC estimates using the FRP approach.

### 3.4 Fuel consumption for different fuel categories

As discussed in Sect. 3.1, the interpretation of average FC values for each biome should be done carefully. As an alternative to biome-averaged values, we also provided FC for specific fuel categories, which may be more useful for certain research areas or modeling communities. In Table 2 fuel category information was presented for the savanna, tropical forest, temperate forest and boreal forest biomes. We focused on the main fuel categories found in the literature, and classified these according to the US classification system. Most of these fuel categories were similarly defined in different studies and biomes; the woody debris classes for example were systematically based on their time lag. However, for measurements conducted in boreal forests the definition of woody fuel classes was less consistent, mainly due to differences between Canadian and American sampling methodologies (Keane, 2012). The difference between surface and ground fuels can therefore be especially vague: e.g., litter is classified as surface fuel according to the US fire management standards, while many Canadian studies define litter and organic soils as the forest floor and thus as the ground fuel class. Obviously, this can cause problems when comparing studies, and therefore we recommend a more uniform measurement protocol for this fuel type and biome.

Certain fuel type averages presented in this paper were based on a minimum of three different studies. For these fuel categories specifically, more field measurements are needed to decrease the uncertainty and better understand the variations found, especially within the boreal and tropical forest biomes. Measurements in the boreal and tropical peat biomes deserve specific attention in future measurement campaigns: although peat fires have been studied in several field campaigns, they still remain one of the least understood fire types due to poor knowledge of the depth of the burning and the complex mix of trace gases emitted in these fires as a consequence of the below-ground combustion that is less efficient than during surface or crown fires. Additional studies are needed in order to capture fully the variability and processes occurring in these biomes, especially considering their large FL and FC. Another biome that deserves more attention in future studies is crop residue, since our understanding of FC variability for different crop types is still poor.

## 4 Summary

This study aimed to compile peer-reviewed literature on measured fuel consumption in landscape fires. The field measurements were partitioned into 11 different biomes, and for each biome we have reported biome averages and other statistics. For some biomes, we provided information on different fuel categories as well. The number of study sites varied from 1 for the tundra biome to 39 different measurement sites in the boreal forest biome. In total, we compiled 124 unique

measurement locations. The biome-averages and fuel-type-specific data of fuel load and fuel consumption can be used to constrain models, or can be used as an input parameter in calculating emissions. Care should be taken though with using biome-averaged values because it is unclear whether these are representative and because there is substantial variability within biomes, as indicated by the large standard deviations found.

Modeled values from GFED3 corresponded reasonably well to the co-located measured values for all biomes except for the savanna and tropical forest, where GFED-derived values were over a factor 2 too high. In tropical forests, part of this discrepancy can be explained because field measurements only take one fire into account, while GFED also accounts for consecutive fires that boost fuel consumption.

Although the overall spatial representativeness of the fuel consumption field measurements was reasonable for most fire-prone regions, several important regions from a fire emissions perspective – including Southeast Asia, eastern Siberia, and central Africa – were severely under-represented. When new information on fuel consumption becomes available, the field measurement database will be updated. The most up-to-date version can be retrieved from <http://www.falw.vu/~gwerf/FC>. As a next step, we aim to improve our understanding of the drivers of regional and temporal variability within biomes, as well as for different fuel categories.

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